CHAPTER III - SUMMARY OF TROPICAL CYCLONES

1. WESTERN NORTH PACIFIC TROPICAL CYCLONES

During 1980, the western North Pacific experienced the second consecutive year of below normal tropical cyclone activity. Twenty-eight tropical cyclones occurred during both 1979 and 1980 as compared to the average annual total of about 33. Four significant tropical cyclones failed to develop beyond the tropical depression (TD) stage and nine tropical storms (TS) failed to reach typhoon intensity. Of the 15 tropical cyclones that developed to typhoon (TY) intensity, only 2 reached the 130 kt (67 m/sec) intensity necessary to be classified as super typhoons (ST). Tropical cyclones reaching tropical storm intensity or greater are assigned names in alphabetical order from a list of alternating male/female names found in CINPACINST 3140.1 CH-2. Different lists of alternating male/female names are used for eastern and central North Pacific and North Atlantic cyclones. Each tropical cyclone's maximum surface winds (MAX SFC WND), in knots, and minimum observed sea

TABLE 3-1 WESTERN NORTH PACIFIC

1980 SIGNIFICANT TROPICAL CYCLONES

CYCLONE	TYPE	NAME	PERIOD OF WARNING	CALENDAR DAYS OF WARNING	MAX SFC WIND(KT)	MIN OBS SLP	NUMBER OF WARNINGS	DISTANCE TRAVELLED(NM)
01	TD	TD-01	20 MAR-24 MAR	5	30	1000	17	2439
02	TS	CARMEN	05 APR-08 APR	4	60	980	9	1179
03	TY	DOM	09 MAY-19 MAY	11	90	956	42	1938
04	TY	ELLEN	13 MAY-21 MAY	9	110	931	34	2423
05	TS	FORREST	20 MAY-26 MAY	7	55	990	26	2451
06	TS	GEORGIA	21 MAY-24 MAY	4	55	985	12	993
07	TS	HERBERT	24 JUN-28 JUN	5	50	980	15	2521
08	TS	IDA	06 JUL-11 JUL	6	60	980	23	1527
09	TY	JOE	17 JUL-23 JUL	7	105	940	25	2541
10	TD	TD-10	17 JUL-19 JUL	3	30	1000	7	1007
11	ST	KIM	20 JUL-27 JUL	8	130	908	29	2661
12	TY	LEX	29 JUL-07 AUG	10	80	962	36	1810
13	TY	MARGE	08 AUG-15 AUG	8	110	944	31	1980
14	TD	TD-14	15 AUG-16 AUG	2	20	1003	7	229
15	TY	NORRIS	24 AUG-28 AUG	5	90	950	20	1710
16	TD	TD-16	04 SEP-06 SEP	3	25	1002	8	776
17	TY	ORCHID	07 SEP-11 SEP	5	85	958	19	2043
18	TY	RUTH	14 SEP-16 SEP	3	65	975	13	60
19	TY	PERCY	14 SEP-19 SEP	6	125	919	20	1260
20	TY	SPERRY	15 SEP-20 SEP	6	65	987	22	2624
21	TS	THELMA	26 SEP-30 SEP	5	55	982	16	1681
22	TY	VERNON	27 SEP-03 OCT	7	105	935	25	2141
23	ST	WYNNE	04 OCT-14 OCT	11	150	890	44	3728
24	TS	ALEX	12 OCT-14 OCT	3	35	999	8	1844
25	TY	BETTY	29 OCT-07 NOV	10	120	928	39	3228
26	TS	CARY	29 OCT-01 NOV	4	40	998	14	1068
27	TY	DINAH	21 NOV-25 NOV	5	100	941	17	3530
28	TS	ED	16 DEC-21DEC	6	50	988	20	815
			1980 TOTALS	128*			598	

* OVERLAPPING DAYS INCLUDED ONLY ONCE IN SUM.

level pressure (MIN OBS SLP), in millibars, were obtained from best estimates based on all available data. The distance travelled, in nautical miles, was calculated from the JTWC official best track (see Annex A).

Table 3-2 provides further information on the monthly distribution of tropical cyclones and statistics on Tropical Cyclone Formation Alerts and Warnings. The number of warning days decreased from 149 to 128 from 1979 to 1980.

TABLE 3-2.

1980 SIGNIFICANT TROPICAL CYCLONE STATISTICS

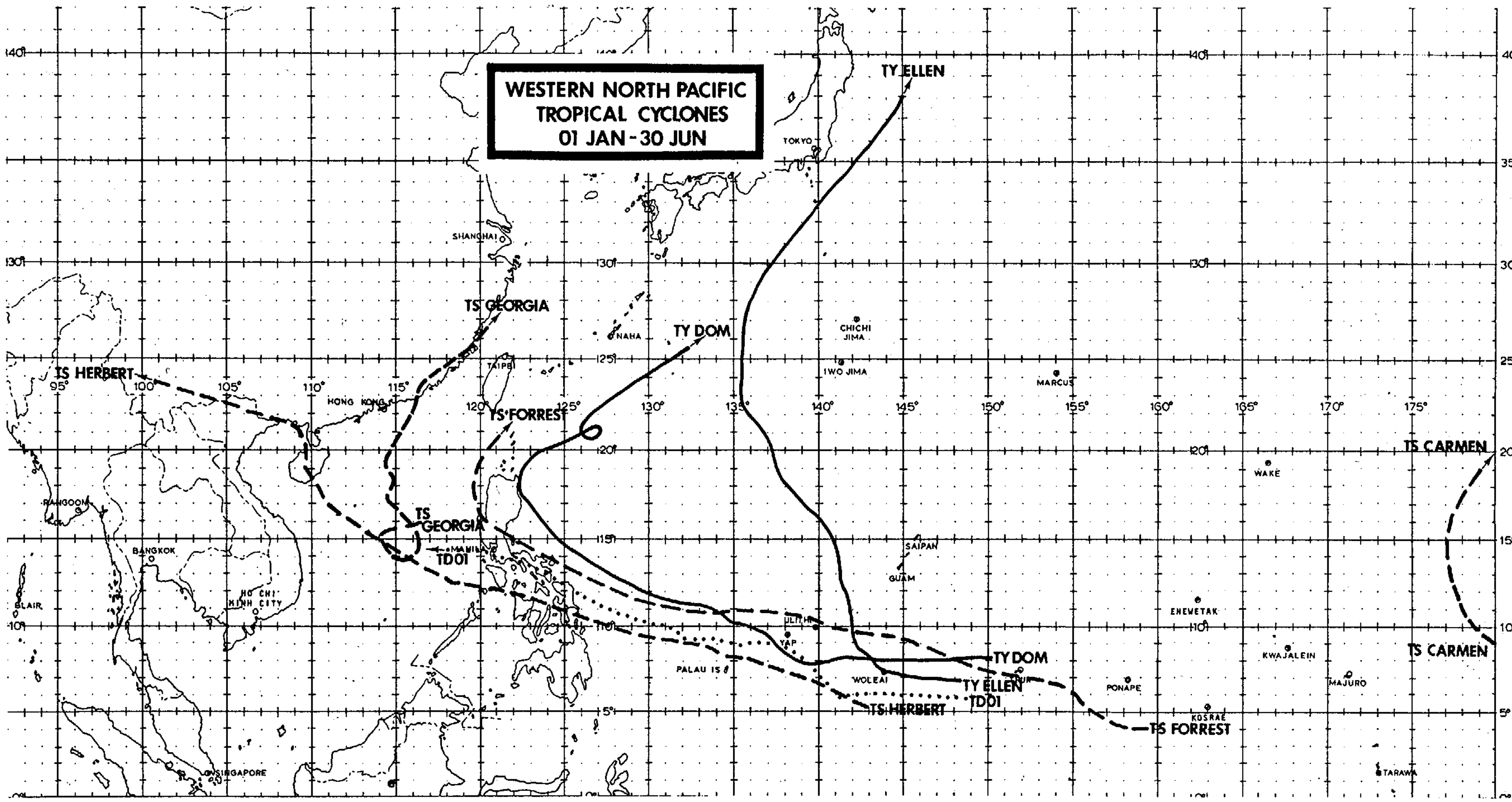
WESTERN NORTH PACIFIC	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTAL	(1959-79) AVERAGE
TROPICAL DEPRESSIONS	0	0	1	0	0	0	1	1	1	0	0	0	4	4.8
TROPICAL STORMS	0	0	0	1	2	1	1	0	1	2	0	1	9	10.0
TYPHOONS	0	0	0	0	2	0	3	2	5	2	1	0	15	17.8
ALL CYCLONES	0	0	1	1	4	1	5	3	7	4	1	1	28	32.6
(1959-79) AVERAGE	.6	.4	.6	.9	1.4	2.0	5.2	6.7	6.0	4.7	2.7	1.4	32.6	

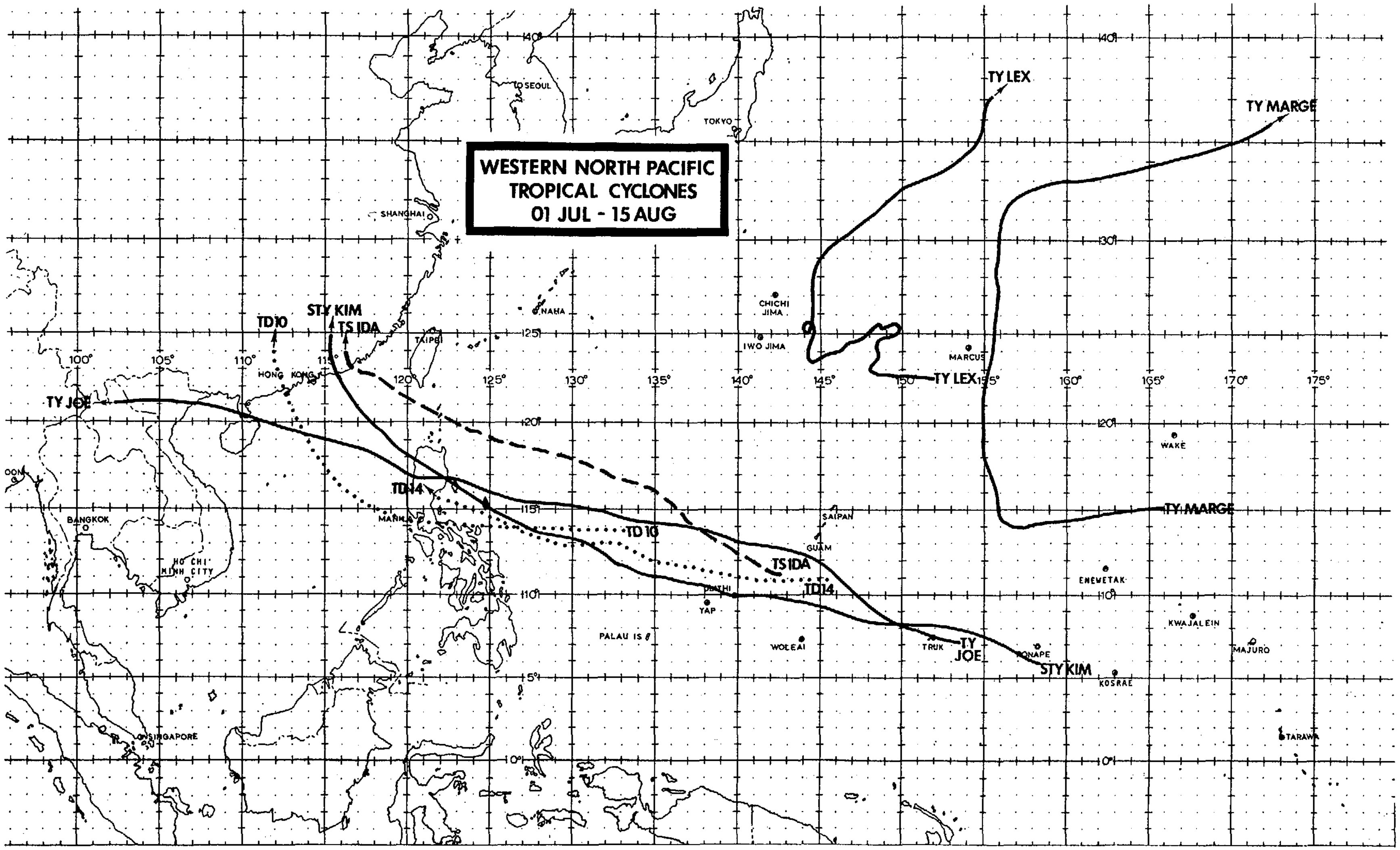
FORMATION ALERTS — 28 of 37 (76%) Formation Alert Events developed into tropical cyclones. Tropical Cyclone Formation Alerts were issued for all significant tropical cyclones which developed during 1980.

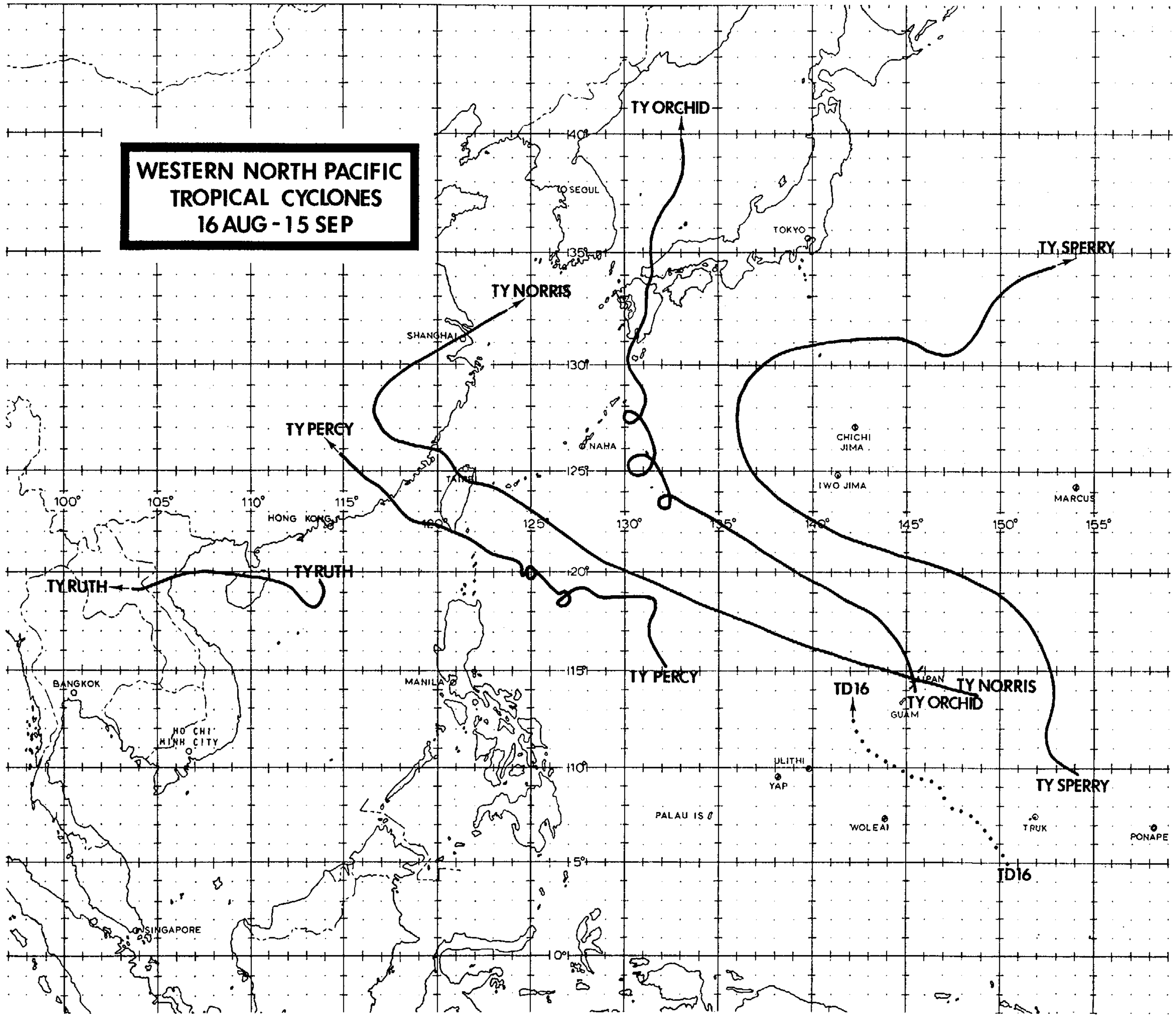
WARNINGS — Number of warning days: 128

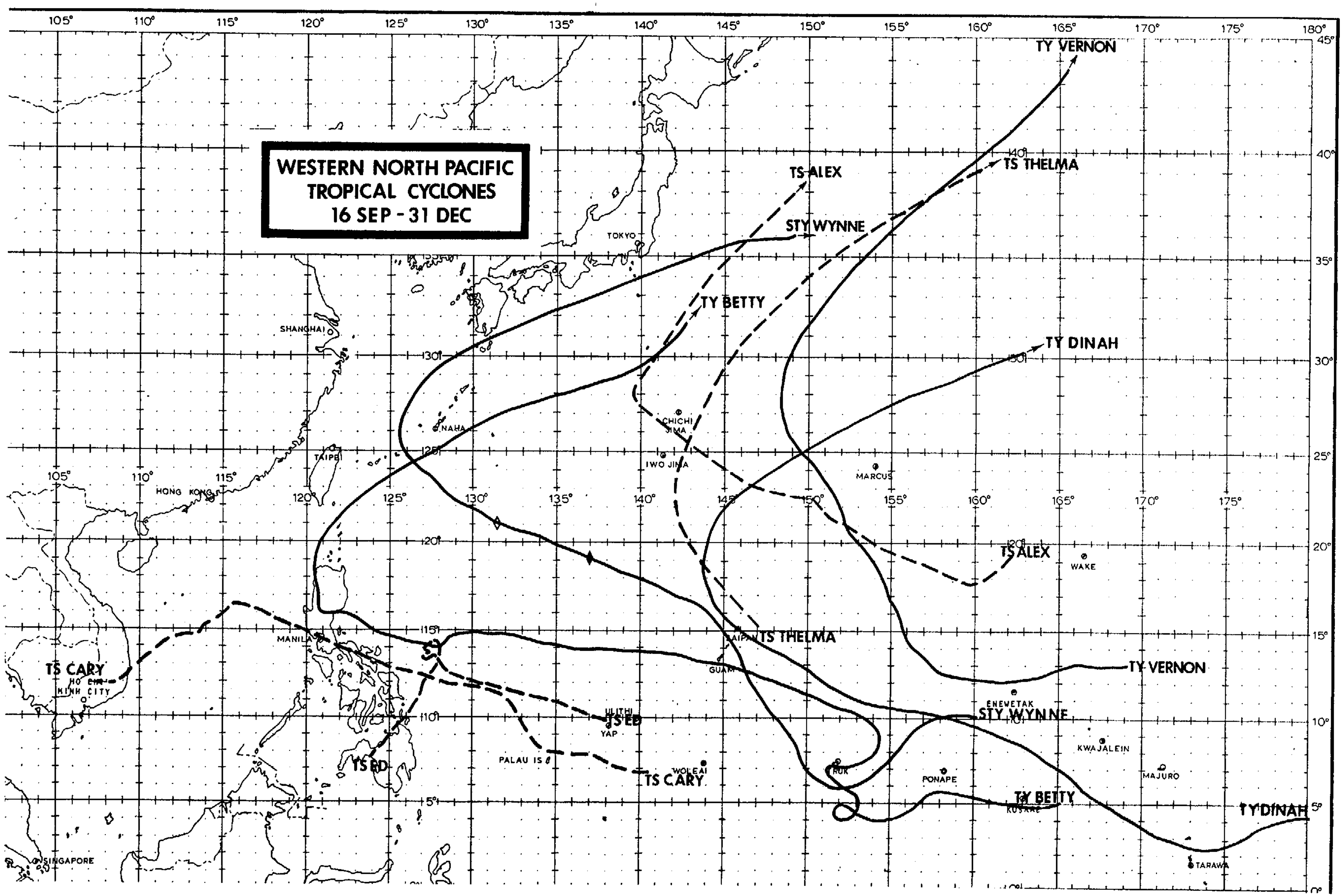
Number of warning days with 2 cyclones: 37

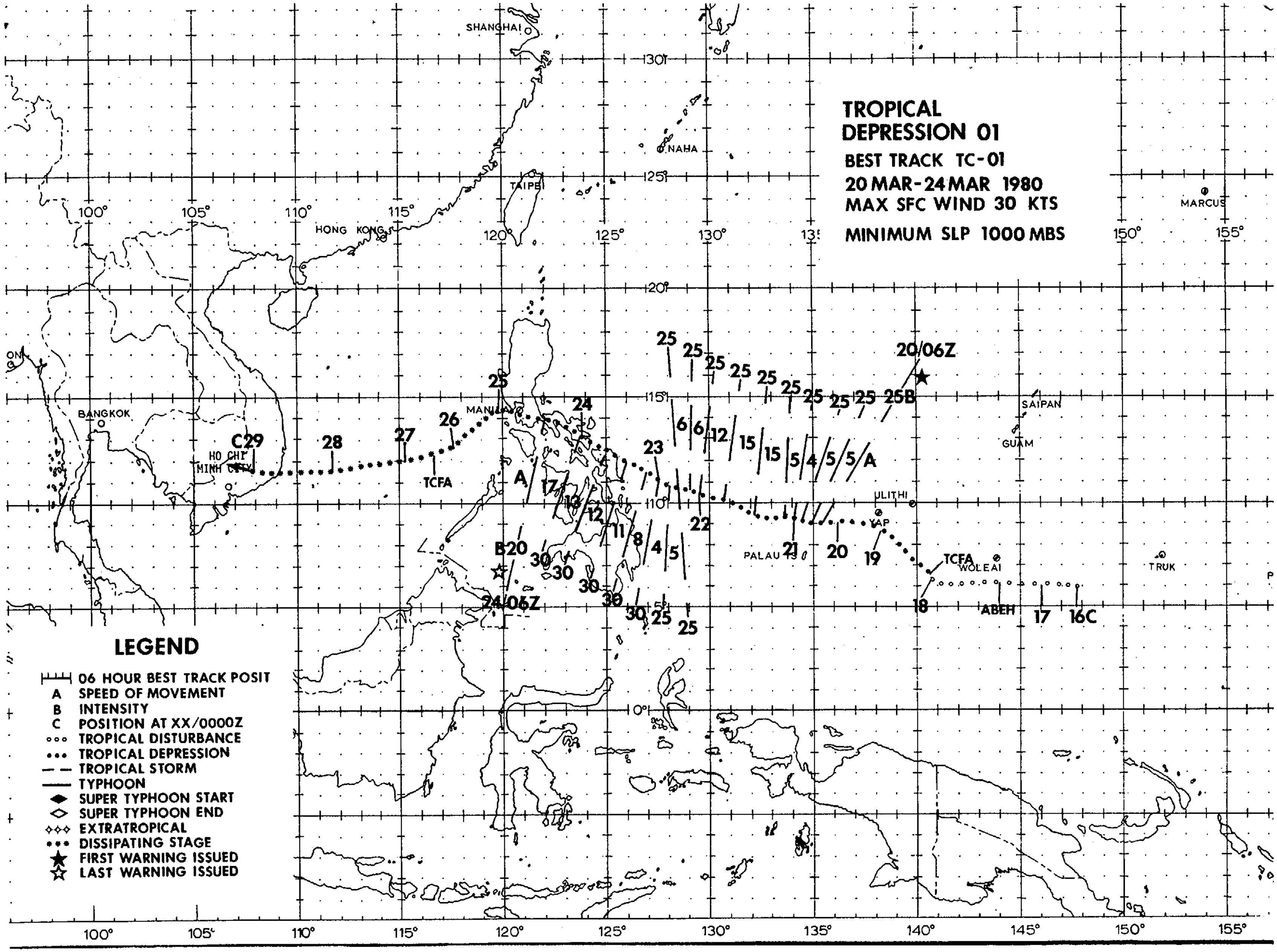
Number of warning days with 3 or more cyclones: 3


WESTERN NORTH PACIFIC
TROPICAL CYCLONES
01 JAN - 30 JUN
TY ELLEN
TY DOM
TS GEORGIA
TS HERBERT
TS FORREST
TS CARMEN
TDOI
TOKYO
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
TAIPEI
HONG KONG
MANILA
RANGOON
BANGKOK
HO CHI MINH CITY
BLAIR
SINGAPORE
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
PONAPE
KOSRAE
ENEWETAK
WAKE
KWAJALEIN
MAJURO
TARAWA


WESTERN NORTH PACIFIC
TROPICAL CYCLONES
01 JUL - 15 AUG
TY LEX
TY MARGE
STY KIM
TS IDA
TD10
TD 10
TD14
TS IDA
TD14
TY JOE
TY JOE
STY KIM
TY LEX
TY MARGE
SEOUL
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
CHICHI JIMA
IWO JIMA
MARCUS
WAKE
BANGKOK
HO CHI MINH CITY
MANILA
SAIPAN
GUAM
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
KOSRAE
ENEWETAK
KWAJALEIN
MAJURO
TARAWA
SINGAPORE
100°
105°
110°
115°
120°
125°
130°
135°
140°
145°
150°
155°
160°
165°
170°
175°
40°
30°
25°
20°
15°
10°
5°


WESTERN NORTH PACIFIC
TROPICAL CYCLONES
16 AUG - 15 SEP
TY ORCHID
TY SPERRY
TY NORRIS
TY PERCY
TY RUTH
TY RUTH
TY PERCY
TD16
TY NORRIS
TY ORCHID
TY SPERRY
TD16
SEOUL
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
BANGKOK
HO CHI MINH CITY
SINGAPORE
CHICHI JIMA
IWO JIMA
MARCUS
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
100°
105°
110°
115°
120°
125°
130°
135°
140°
145°
150°
155°
40°
35°
30°
25°
20°
15°
10°
5°
0°


WESTERN NORTH PACIFIC
TROPICAL CYCLONES
16 SEP - 31 DEC
TY VERNON
TS THELMA
TS ALEX
STY WYNNE
TY BETTY
TY DINAH
TS CARY
TS ED
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
HO CHI MINH CITY
MANILA
CHICHI JIMA
IWO JIMA
MARCUS
WAKE
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
KUSAIE
ENEWETAK
KWAJALEIN
MAJURO
TARAWA
SINGAPORE
105°
110°
115°
120°
125°
130°
135°
140°
145°
150°
155°
160°
165°
170°
175°
180°
45°
40°
35°
30°
25°
20°
15°
10°
5°
0°


TROPICAL DEPRESSION 01
BEST TRACK TC-01
20 MAR-24 MAR 1980
MAX SFC WIND 30 KTS
MINIMUM SLP 1000 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
BANGKOK
HO CHI MINH CITY
SAIPAN
GUAM
MARCUS
ULITHI
YAP
PALAU
WOLEAI
TRUK
20/06Z
24/06Z
TCFA
100° 105° 110° 115° 120° 125° 130° 135° 140° 145° 150° 155°

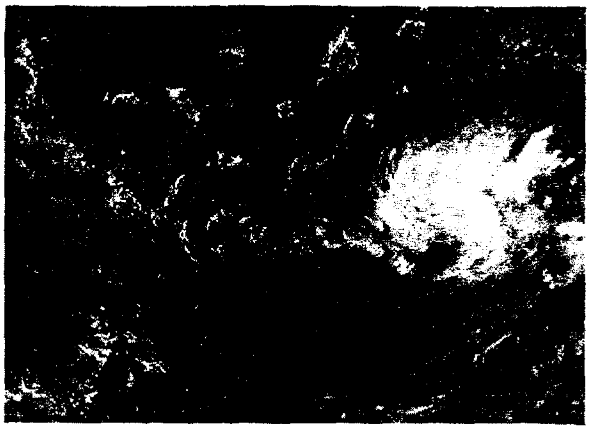


FIGURE 3-01-1. TD 01 at 15-20 kt (8-10 m/sec) intensity about 500 nm (926 km) south-southwest of Guam, 18 March 1980, 0120Z. (DMSP imagery)

TD 01 was first detected as an area of increased convective activity about 500 nm (926 km) south-southeast of Guam on 16 March. During the early part of the year, intense convective activity is usually located south of the equator. March is the start of the transition period when the equatorial trough begins to migrate slowly northward. During this period, the equatorial trough can occasionally extend into the Northern Hemisphere. This extension, however, is normally short-lived because the southwest monsoon has yet to become fully established. Post-analysis indicates that TD 01 developed from a temporary extension of the equatorial trough into the Northern Hemisphere.

The first aircraft reconnaissance mission into TD 01 on the morning of 18 March reported 15-20 kt (8-10 m/sec) surface winds, primarily in the northern semicircle, and a minimum sea-level pressure of 1005 mb. Based on this information and satellite imagery which showed improved upper-level outflow in the southeast quadrant (Fig. 3-01-1), a Tropical Cyclone Formation Alert (TCFA) was issued at 180300Z.

The tropical disturbance was monitored closely for the next 48 hours. The first reconnaissance mission also reported a 60 nm (111 km) displacement between the surface center and the 1500 ft (457 m) center. Subsequent missions discovered a similar displacement between the surface and 700 mb centers. This was consistent with the synoptic data which showed that strong mid- to upper-level southeasterlies were causing TD 01 to tilt with height toward the northwest.

By the 20th, surface winds in the southern semicircle had increased to 20 kt (10 m/sec), while 30 kt (15 m/sec) winds were observed in the northern semicircle. The circulation was better defined on satellite imagery, and the MSLP had decreased to 1000.7 mb. Continued development was expected and the first warning on TD 01 was issued at 200600Z.

Taking into consideration the strong vertical wind shear and the fact that March is historically a month of minimum typhoon development, TD 01 was never forecast to reach more than minimal tropical storm strength of 40 kt (21 m/sec).

From 20 through 24 March, TD 01 followed a climatological west-northwest track toward Luzon, occasionally showing speed changes as it responded to a series of mid-level short-wave troughs moving eastward across the Pacific from the Asian mainland.

As TD 01 approached southeastern Luzon, it began to interact both with a shear line extending toward it from the northeast and a building high pressure ridge between Taiwan and Luzon (Fig. 3-01-2). The net result was a flare-up in the convective activity and an increase in surface wind speed north of the surface center. Although two land stations reported 40 kt (21 m/sec) winds during landfall on Luzon, the sea-level pressures were not observed below 1007 mb. Considering the effects of topography, 30 kt (15 m/sec) appears to be the best estimate of TD 01's intensity at that time. Figure 3-01-2 shows that northeasterly winds of 25-40 kt (13-21 m/sec) were present north of TD 01 to the vicinity of Taiwan. These strong winds were being enhanced by TD 01, but were more the result of the building high pressure ridge off the Asian mainland. Therefore, an extratropical wind warning was issued for the area by NAVOCEANCOMCEN Guam.

After making landfall, TD 01 tracked slowly westward south of Manila into the South China Sea. A TCFA was issued for the remnants of TD 01 at 260615Z when improved organization of the cloud pattern (Fig. 3-01-3) suggested that regeneration might occur. The disturbance was watched for three more days, but ship reports showed nothing more than a weak wave in the east-northeasterly flow, and the system dissipated rapidly after moving ashore on the Vietnam coast near Ho Chi Minh City.

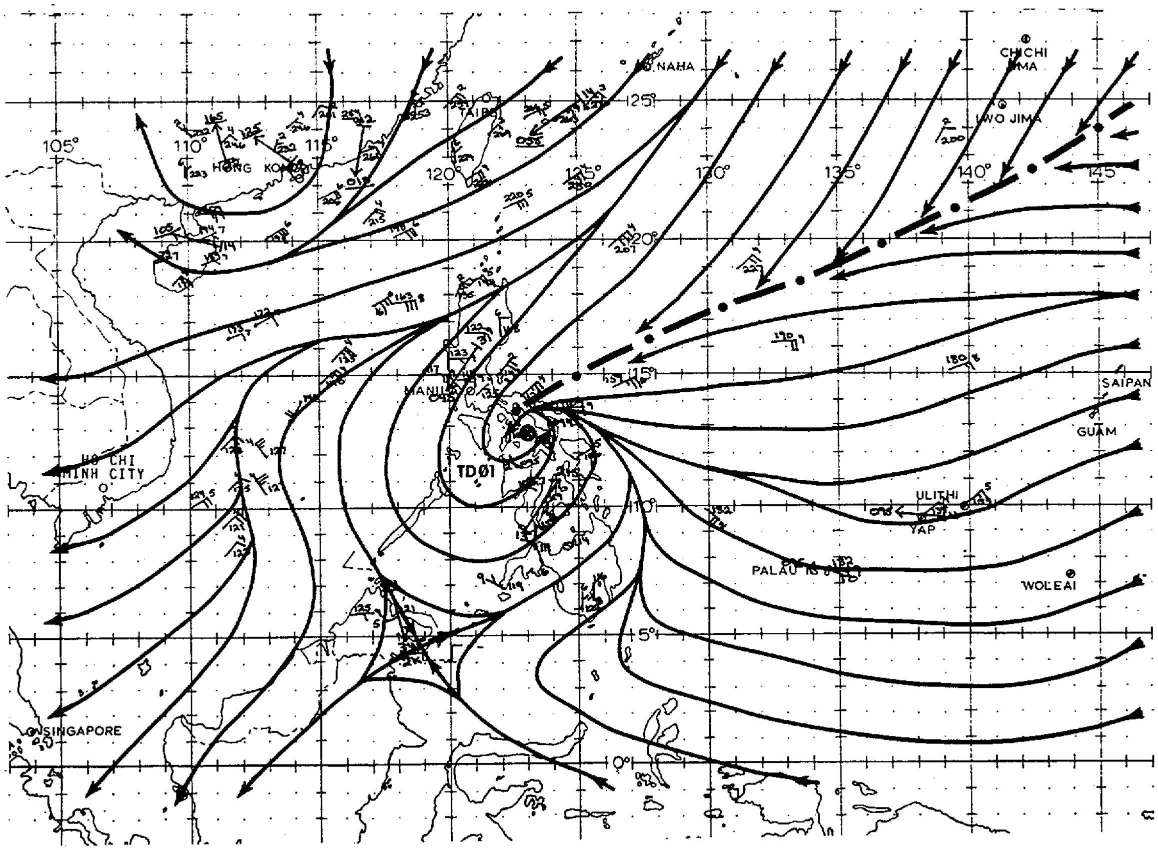


FIGURE 3-01-2. The 240000Z March 1980 surface (———) / gradient-level (ddd←——<ff) wind data and streamline analysis. Wind speeds are in knots.

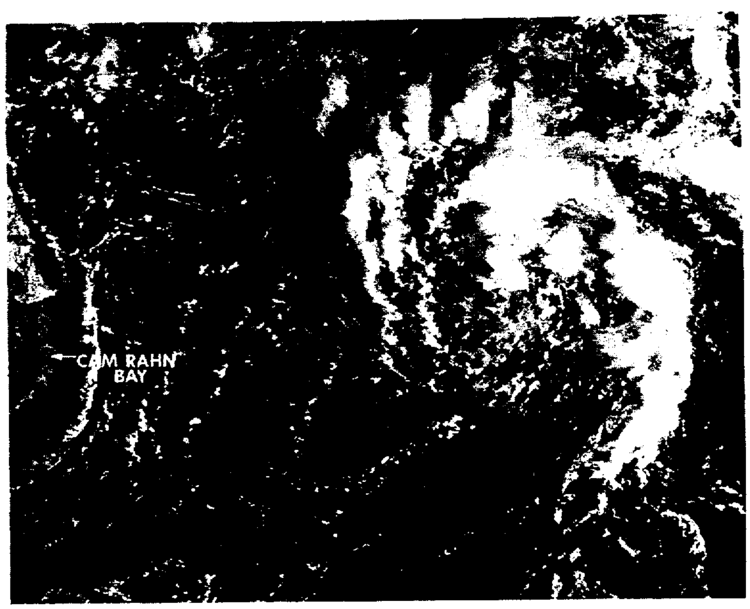
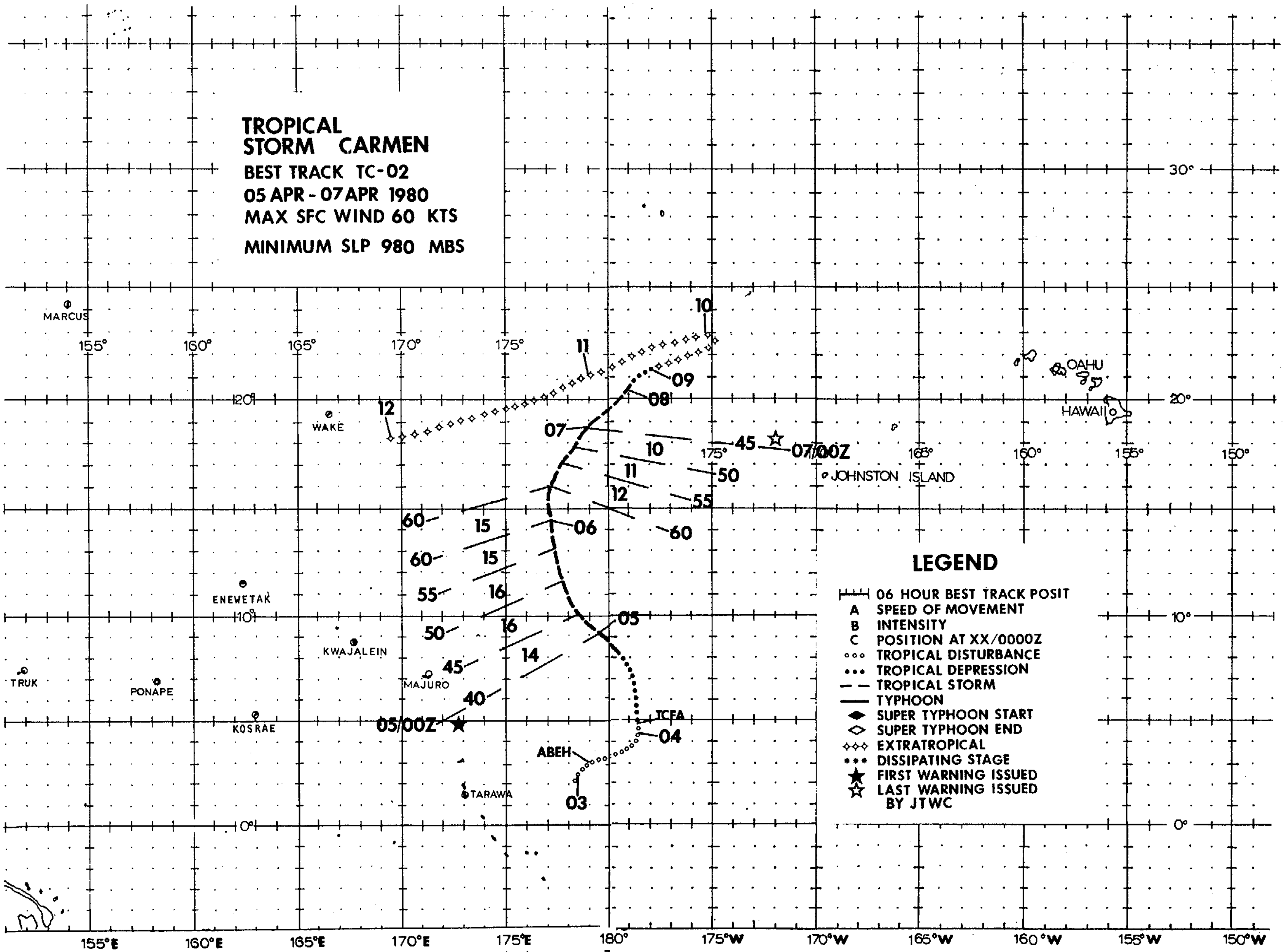


FIGURE 3-01-3. The remnants of TD 01 in the South China Sea showing signs of regeneration, 26 March 1980, 0206Z. (DMSP imagery)


TROPICAL
STORM CARMEN
BEST TRACK TC-02
05 APR - 07 APR 1980
MAX SFC WIND 60 KTS
MINIMUM SLP 980 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
BY JTWC
MARCUS
WAKE
ENEWETAK
KWAJALEIN
TRUK
PONAPE
MAJURO
KOSRAE
TARAWA
JOHNSTON ISLAND
OAHU
HAWAII
ABEH
TCFA
05/00Z
07/00Z
155°E
160°E
165°E
170°E
175°E
180°
175°W
170°W
165°W
160°W
155°W
150°W
30°
20°
10°
0°

Tropical Storm Carmen, the second significant tropical cyclone of the season, might well have gone undetected if it had occurred prior to the advent of meteorological satellite surveillance. Carmen developed in and tracked through a very sparse synoptic data region near the dateline in early April 1980. Once organized, Carmen's closest point of approach to a reporting station (Majuro Atoll, WMO 91376) was 450 nm (833 km). During its entire life, Carmen was closely monitored by the Joint Typhoon Warning Center (JTWC) and the Central Pacific Hurricane Center (CPHC) using polar-orbiting and geostationary satellites to confirm Carmen's existence.

Available satellite imagery and synoptic data indicated that Carmen developed in a relatively active near-equatorial trough (NET) during a period in which a parallel disturbance, TC 20-80 (Wally), was developing in the Southern Hemisphere. (The term parallel disturbances is also referred to as "double vortices".) Similar to many previous cases, most recently Typhoon Kim (1977) and Typhoon Lucy (1977) and their respective Southern Hemisphere cyclones, Carmen and TC 20-80 took nearly mirror-image tracks over open water. In this case, each cyclone moved towards its respective pole in response to a weakness in each hemisphere's sub-tropical ridge. Once organized, Carmen moved north-northwest and then, at the ridge axis, began its recurvature to the northeast. Similarly, TC 20-80 moved south-southwest until it began recurvature to the southeast at the ridge axis. Although TC 20-80 accelerated in its extratropical transition near 26 degrees south latitude, Carmen slowed as she moved eastward across the dateline. Several days later Carmen dissipated in the northeast trade wind flow south of Wake Island.

The disturbance which became Tropical Storm Carmen was first detected in satellite imagery at 0000Z on 2 April. By 021800Z, the area of convection had moved from the equator to near 02N 178E. At 030600Z, the Significant Tropical Weather Advisory (ABEH PGTW) discussed a surface circulation near 03N 179E. The major convection associated with the circulation continued to move northeast at 10 kt (19 km/hr) east of the dateline. The Central Pacific Hurricane Center (CPHC) monitors developing tropical cyclones east of the dateline and the responsibility for issuing tropical cyclone formation alerts (TCFA) in this region belongs to the Naval Western Oceanography Center (NWOC) at Pearl Harbor, Hawaii. By 0200Z on 4 April, the organization of the disturbance had improved significantly and NWOC issued a TCFA for an area that straddled the dateline between 04N and 08N. At 050000Z, the developing cyclone moved west of the dateline, and based on the improved satellite signature, the first warning on TD02 was issued at that time. During the next 48 hours, Carmen intensified, reaching a peak intensity of 60 kt (31 m/sec) at approximately 060000Z. Figure 3-02-1 shows satellite imagery of Carmen at peak intensity. Carmen then gradually weakened as she approached the dateline for a third time (second approach from the west). The last JTWC warning was issued at 070000Z and the CPHC issued its first warning at 070600Z. While east of the dateline, Carmen continued to weaken as her movement slowed to 5 kt (9 km/hr). The final warning was issued by CPHC at 090000Z with TD02 near 21.5N 178W.

Due to Carmen's location (near the dateline) and month of occurrence (April), JTWC forecasters had few viable forecasting aids to develop their warnings. Climatology and analog programs were non-existent for the area and season, and the steering model is unreliable south of 10N. Without the input of these valuable aids, the initial warning was based on sparse mid-level synoptic data and described a north-northwest track with recurvature near 17N. This basic track was maintained in subsequent JTWC warnings. Maintenance of this basic track through recurvature provided JTWC with 72-hour forecast errors (210 nm (389 km)) which were significantly lower than the 10-year average of 348 nm (644 km).

Intensity estimates and forecasts were based entirely on the Dvorak method for estimating tropical cyclone intensity (1975) The first series of Dvorak intensity estimates at 041954Z, 050000Z and 050233Z supported 35 kt (18 m/sec) maximum winds. However, upgrading to tropical storm status did not occur until the 051200Z warning. This delay is not unusual. Initial warnings tend to be conservative because satellite imagery of a developing tropical cyclone often appears more intense for a brief period before returning to a more "normal" signature for the early development stage. Indeed, the Dvorak method has a built in constraint which limits initial estimates to T1.5 (25 kt (13 m/sec)) or less. The initial Dvorak intensities received at JTWC were T2.5 (35 kt (18 m/sec)). In post-analysis, the higher estimates were supported with the trend showing that TD02 (Carmen) actually reached tropical storm strength at 041800Z, 6 hours prior to the first warning.

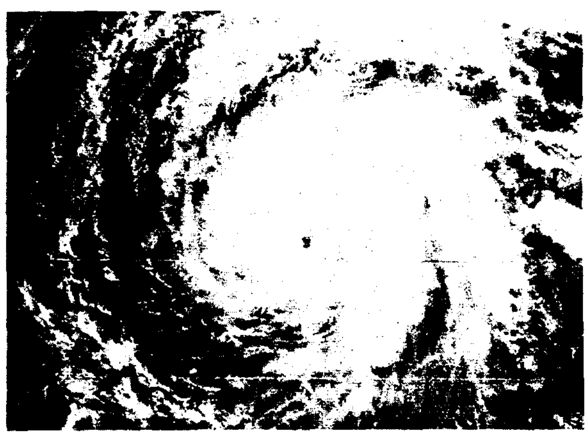
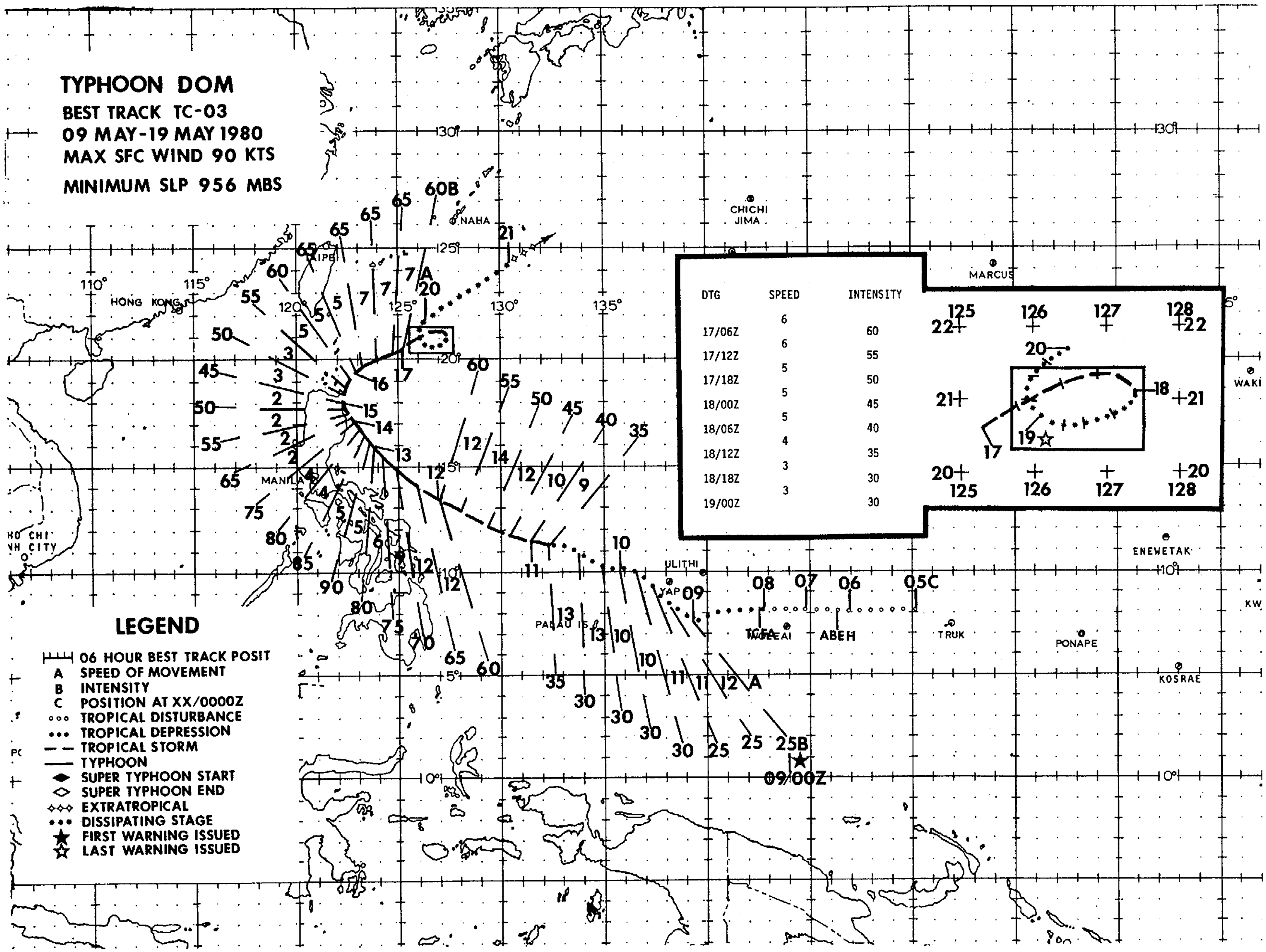


FIGURE 3-02-1. TS Carmen, near maximum intensity of 60 kt (31 m/sec), 05 April 1980, 2231Z. (DMSP imagery)


TYPHOON DOM
BEST TRACK TC-03
09 MAY-19 MAY 1980
MAX SFC WIND 90 KTS
MINIMUM SLP 956 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
DTG SPEED INTENSITY
17/06Z 60
17/12Z 55
17/18Z 50
18/00Z 45
18/06Z 40
18/12Z 35
18/18Z 30
19/00Z 30
6 6 5 5 5 4 3 3
HONG KONG
MANILA
NAHA
CHICHI JIMA
MARCUS
WAKI
ENEWETAK
ULITHI
YAP
PALAU IS.
WOLEAI
TRUK
PONAPE
KOSRAE
HO CHI MINH CITY
09/00Z

Dom was the first tropical cyclone that developed to typhoon intensity in the western North Pacific in 1980. Dom had several features of interest: a pronounced tilt in the vertical axis during the developing stages and the execution of a rare anticyclonic loop in the later stages of his existence.

Satellite imagery showed a weak disturbance which first appeared along the near equatorial trough on 5 May. The disturbance showed no significant development as it tracked across the Caroline Islands during the following three days. The first investigation by reconnaissance aircraft was scheduled on 8 May when a significant increase in convective activity was noted. The weak circulation located by the aircraft prompted JTWC to issue a Tropical Cyclone Formation Alert for an area south of Guam. By the 9th, satellite imagery indicated strong outflow on the west side of the circulation and increased organization of convective cloud elements became evident as the disturbance continued to develop. As the circulation became more organized, reconnaissance aircraft observed an increase in the surrounding surface winds. The first warning on TD03 was issued at 090000Z.

Little change in intensity occurred during the next two days, during which time the 700 mb circulation was displaced as much as 77 nm (143 km) west-southwest of the surface center. This displacement was indicative of marked vertical shear caused by strong mid- to upper-level easterly flow.

Vertical shear remained strong during Dom's early stages of development as he moved westward steered by strong mid- to upper-level easterlies along the southern periphery of the mid-level subtropical ridge axis. On 10 May, a mid-tropospheric low pressure center developed over the Asia Mainland, causing the ridge to recede eastward. This created a weakness in the ridge near the Philippines, allowing Dom to track west-northwestward away from the strong mid- to upper-level easterlies. With the decrease in vertical wind shear, Dom's axis became more vertical and development proceeded. Dom reached tropical storm intensity at 101800Z as an anticyclone with outflow in all quadrants developed at upper-levels.

A large area of low-level convergence formed to the northeast of Dom as evidenced by convective activity shown by satellite

FIGURE 3-03-1. Typhoon Dom at 40 kt (21 m/sec) intensity tracking west-northwestward at 9 kt (17 km/hr), 11 May 1980, 0213Z. (DMSP visual imagery)

imagery on the 11th (Fig. 3-03-1). This are of convection dissipated as an induced ridge formed between Dom and a circulation to the southeast of Guam which would later develop into Typhoon Ellen. Dom attained typhoon intensity at 120600Z. When Dom intensified to 90 kt (46 m/sec), he had a large eye 30 nm (56 nm) in diameter and his speed of movement decreased markedly as he moved away from the strong mid- to upper-level easterly steering flow. Dom became virtually stationary as he drifted slowly toward Luzon with weakening commencing due to the decreased moisture content of the air being drawn into Dom's circulation across the mountainous terrain of Luzon. By 141200Z, Dom had weakened to tropical storm intensity and was tracking northward at 2 kts (1 m/sec) showing indications of impending recurvature.

Dom unexpectedly regained typhoon strength 24 hours after recurvature. Reintensification was made possible by a lessening of the land effect and energy provided by a tongue of warm water extending north of Luzon (Fig. 3-03-2). Dom then tracked northeastward south of the area of maximum sea surface temperature (SST). A later SST analysis (Fig. 3-03-3) showed the decrease in SST which is normally observed after the passage of a tropical cyclone. This decrease in SST is caused primarily by evaporative cooling and the mixing of surface water with cooler sub-surface water and, to a lesser extent, by the addition of rain water and the decrease in solar radiation reaching the surface (Brand, 1970). Dom's final decrease to tropical storm intensity was due to the shearing effect of strong upper-tropospheric westerlies and strong low-level easterlies. The upper-level center continued to track eastward, whereas the surface circulation began a rare anticyclonic loop as it tracked westward under the influence of the low-level easterly flow. At 190000Z, JTWC issued the final warning on Dom, although post-analysis indicated he ceased to exist as a significant tropical cyclone on the 18th.

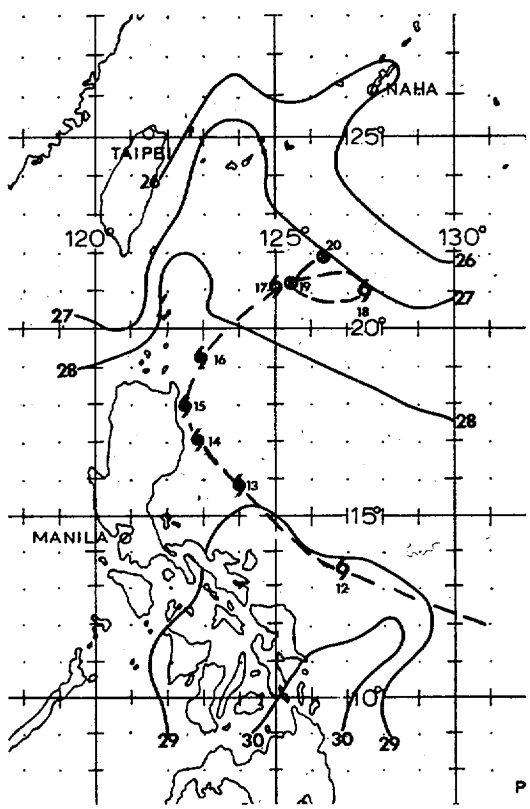


FIGURE 3-03-2. Composite sea surface temperature analysis of data from 10-16 May 1980, produced by the Oceanographic Services Division of Naval Oceanography Command Center, Guam.

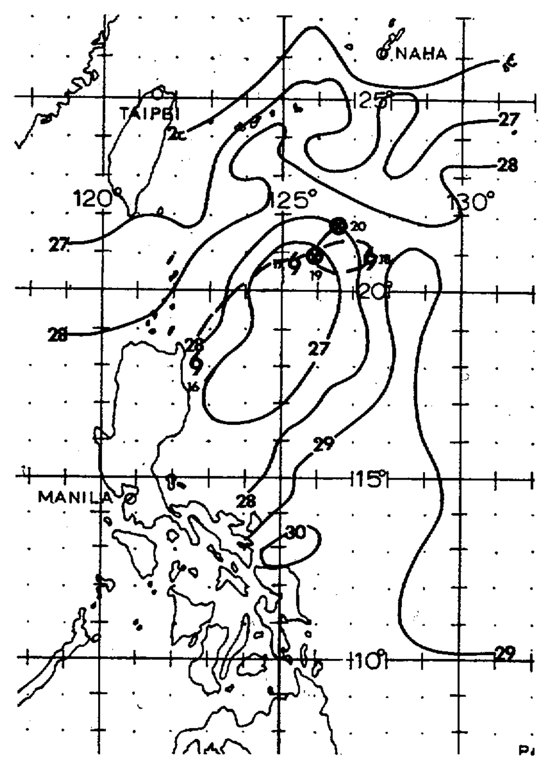
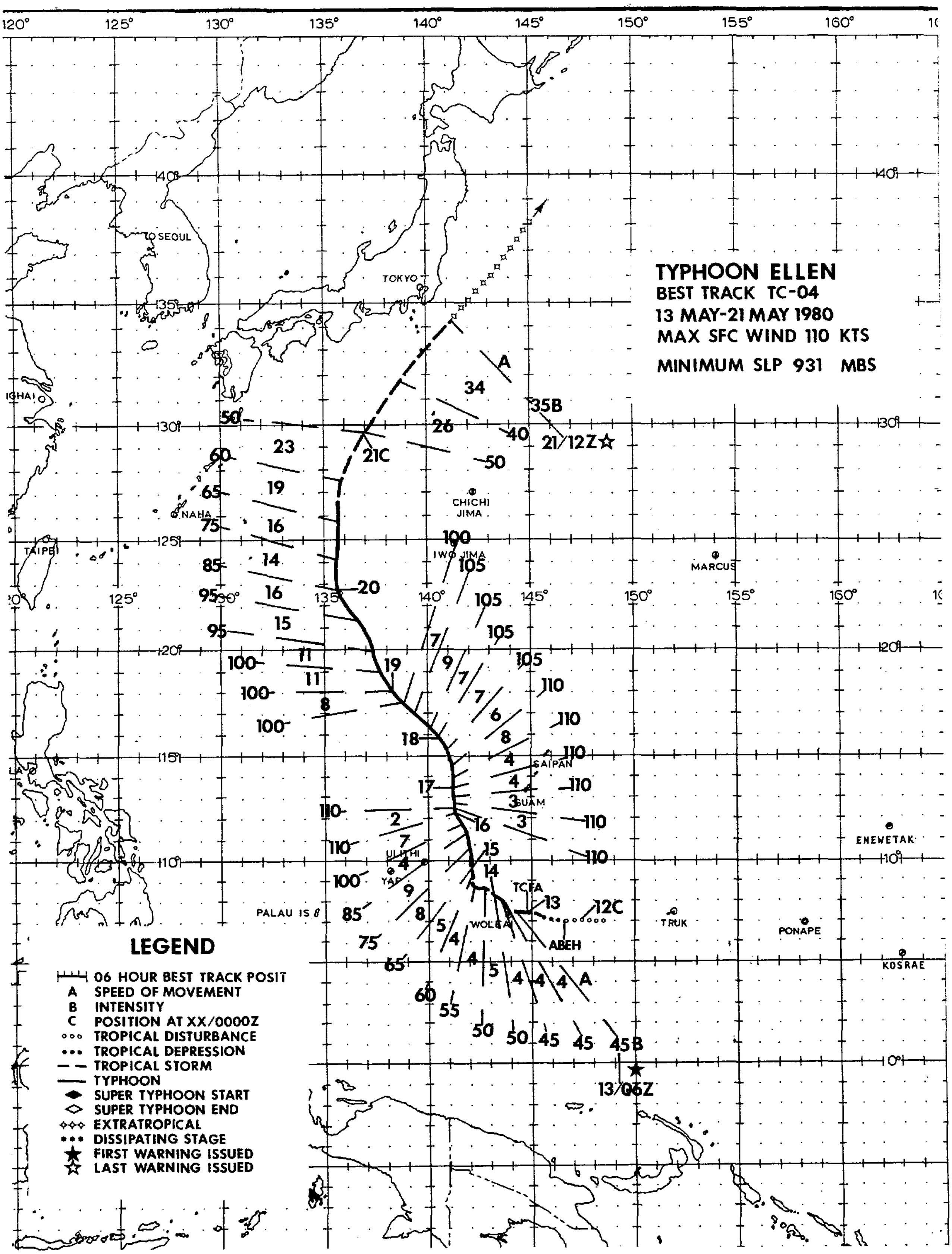


FIGURE 3-03-3. Composite sea surface temperature analysis of data from 17-23 May 1980, produced by the Oceanographic Services Division of Naval Oceanography Command Center, Guam.

Triggered by a mid-tropospheric trough which entered the South China Sea, an extratropical surface low pressure system formed south of Japan at 200000Z with the associated frontal boundary extending to the southwest of Okinawa. At this time, Typhoon Ellen was 600 nm (1113 km) east-northeast of the remnants of Dom. On the 20th, both the remnants of Dom and Typhoon Ellen accelerated toward the extratropical low along the east side of the frontal boundary. By 211200Z, the three systems had merged to form an intensifying mid-latitude storm over the east coast of central Honshu, Japan. This deepening mid-latitude storm tracked northeastward along the northern periphery of the mid-Pacific ridge.


TYPHOON ELLEN
BEST TRACK TC-04
13 MAY-21 MAY 1980
MAX SFC WIND 110 KTS
MINIMUM SLP 931 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SEOUL
TOKYO
NAHA
TAIPEI
CHICHI JIMA
IWO JIMA
MARCUS
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
ENEWETAK
KOSRAE
TCFA
ABEH
21/12Z
13/06Z
21C
12C
35B
45B

Typhoon Ellen developed in an active, near-equatorial trough west of the Truk Islands on 11 May 1980. Strong upper-level divergence over the Caroline Islands and a weak 500 mb steering currents, produced by a northward adjustment of the 500 mb ridge axis to 25N, provided an excellent environment for tropical cyclone development. Ellen was an interesting tropical cyclone from several viewpoints. During her existence, Ellen underwent rapid initial development, abruptly changed track at a low latitude, and followed a slow oscillatory motion for an 18 hour period.

Ellen's initial tropical disturbance became evident on satellite imagery between 111200Z and 120000Z. However, a Tropical Cyclone Formation Alert (TCFA) was not issued at that time because 120000Z synoptic data did not indicate a well-defined surface circulation with lowering surface pressures. A weakening of the satellite signature during the next 12 hours supported this decision. Between 121200Z and 121600Z, Ellen's satellite signature improved markedly and a TCFA was issued. Aircraft reconnaissance at 130422Z confirmed Ellen's rapid development and estimated 45-50 kt (23-26 m/sec) maximum surface winds. The first warning was issued at 130600Z. Post-analysis indicates that Ellen reached tropical storm strength at 121800Z.

Ellen appeared to be following TY Dom's track across the Philippine Sea as she tracked initially west over Woleai Atoll and then west-northwestward toward Ulithi Atoll. On 15 May, Ellen abruptly turned to the north and was headed for Japan. By 150000Z, she was tracking north-northwestward at approximately 8 kt (15 km/hr) and had intensified to 65 kt (33 m/sec). At 170000Z, Ellen passed 220 nm (407 km) west of Guam with maximum sustained surface winds of 110 kt (57 m/sec). Figure 3-04-1 is satellite imagery during this period of Ellen's track.

After her abrupt turn, Typhoon Ellen's surface circulation followed a pronounced

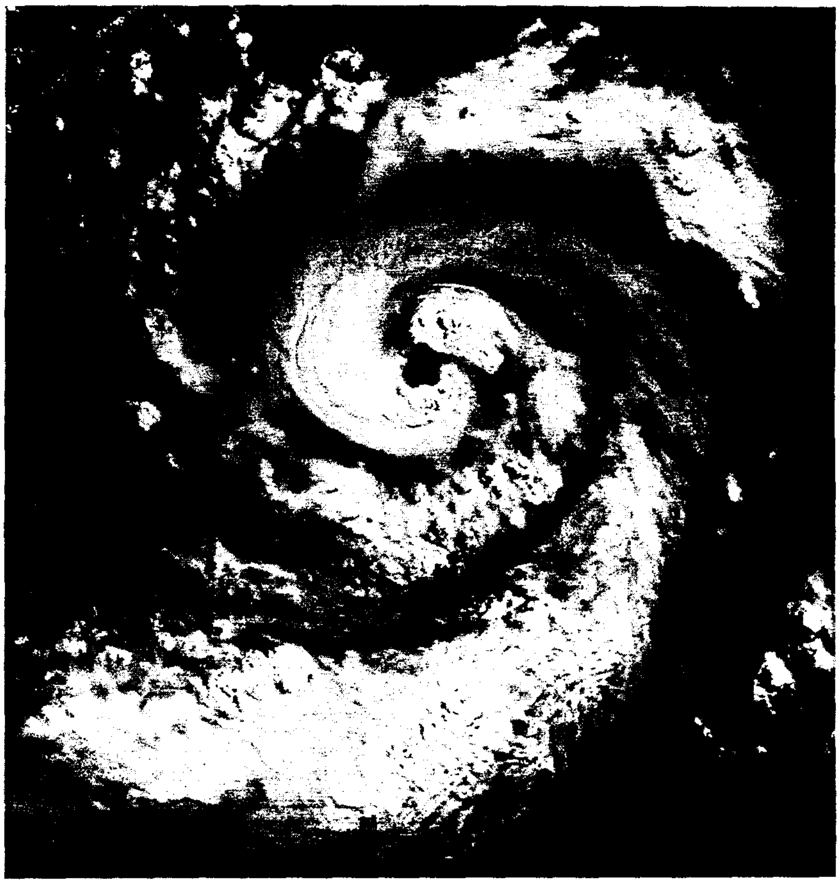


FIGURE 3-04-1. Typhoon Ellen shortly after reaching typhoon intensity, 15 May 1980, 0054Z. (DMSP imagery)

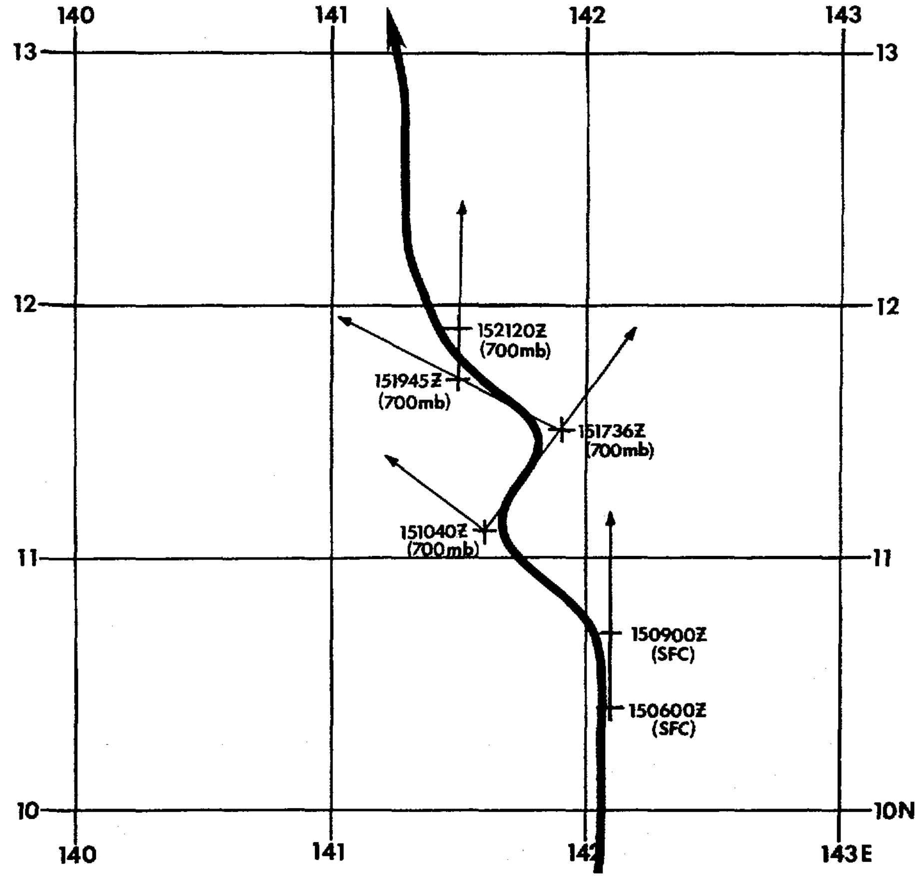


FIGURE 3-04-2. Typhoon Ellen's best track between 141800Z and 160600Z May 1980. Surface and 700 mb positions observed by aircraft reconnaissance, and vectors between successive aircraft fixes, are shown.

oscillatory motion about a basic north-northwest track. Surface positions observed by aircraft reconnaissance and vectors between successive aircraft fixes during this period are illustrated in Figure 3-04-2. These short term oscillations were difficult to interpret and thus made forecasting Ellen's movement very difficult.

As Ellen was undergoing this oscillatory motion, aircraft reconnaissance also observed that the location of minimum sea level pressure appeared to rotate close to the wall cloud in a highly elliptical eye. During the same period, Ellen deepened to her lowest minimum sea level pressure of 931 mb and intensified an additional 45 kt (23 m/sec), reaching her maximum intensity of 110 kt (57 m/sec).

This oscillatory motion and uncertainty in the position and strength of the 500 mb subtropical ridge axis created a significant forecast problem. Forecasts of early recurvature to the northeast did not materialize as Ellen continued on a north-northwest track toward Japan. Once north of the ridge axis, Ellen recurved between 25N and 30N and accelerated northeastward at forward speeds in excess of 30 kt (56 km/hr). Following recurvature, Ellen weakened rapidly and merged with an extratropical low pressure system south of Honshu.

Ellen's actual track passed closer to Japan than originally forecast due to rapid deepening of a mid-latitude trough over Japan and rapid intensification of the subtropical ridge east of Japan. In response, 500 mb winds south of Japan backed in direction and strengthened, causing Ellen to accelerate northeastward and pass 120 nm (222 km) east of Yokosuka Naval Station, Japan (Fig. 3-04-3). Department of Defense resources in Japan reported no major damage, and Yokosuka only reported 20-25 kt (10-13 m/sec) sustained winds during the passage of Ellen. Flooding reported in Kyushu and Shikoku resulted from heavy rain produced by the extratropical low pressure system which eventually merged with Ellen south of Honshu.

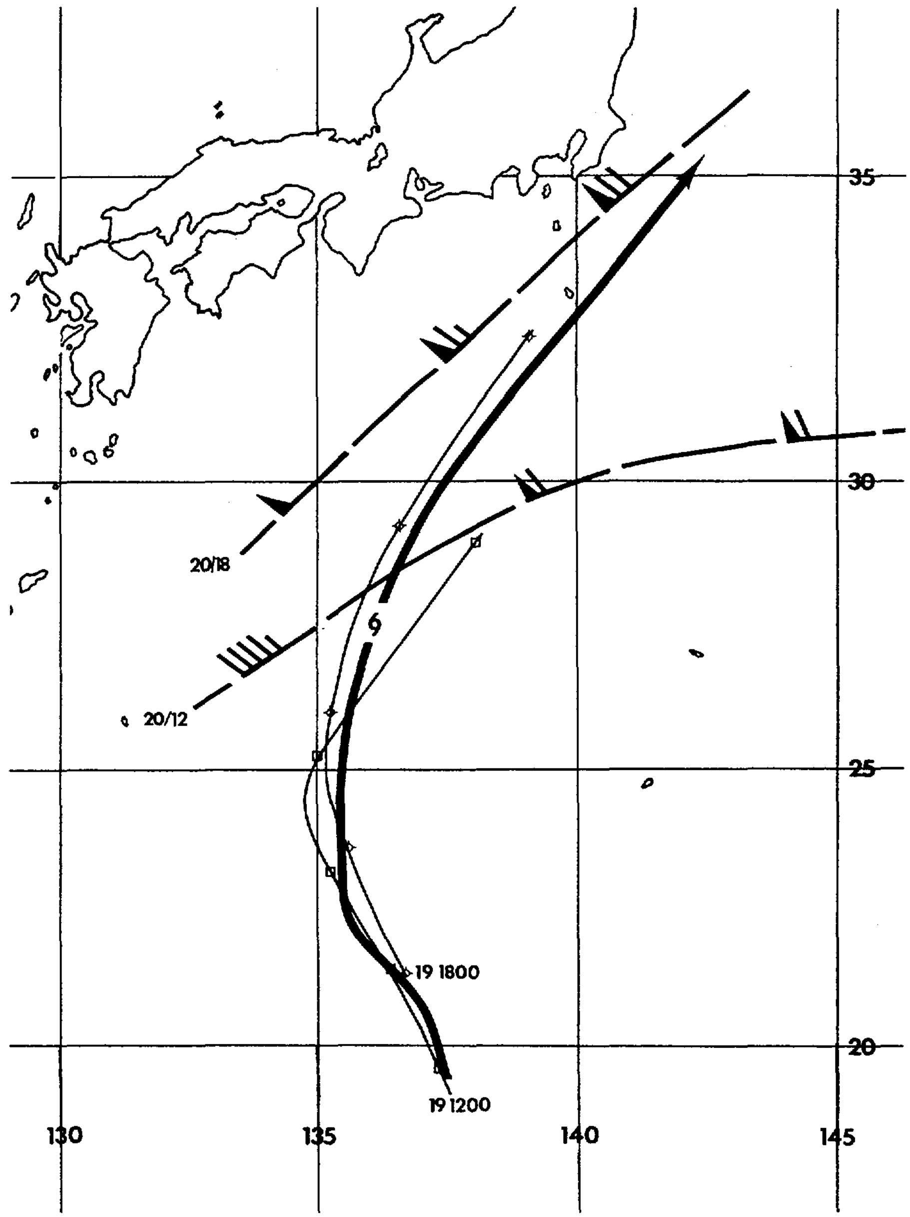
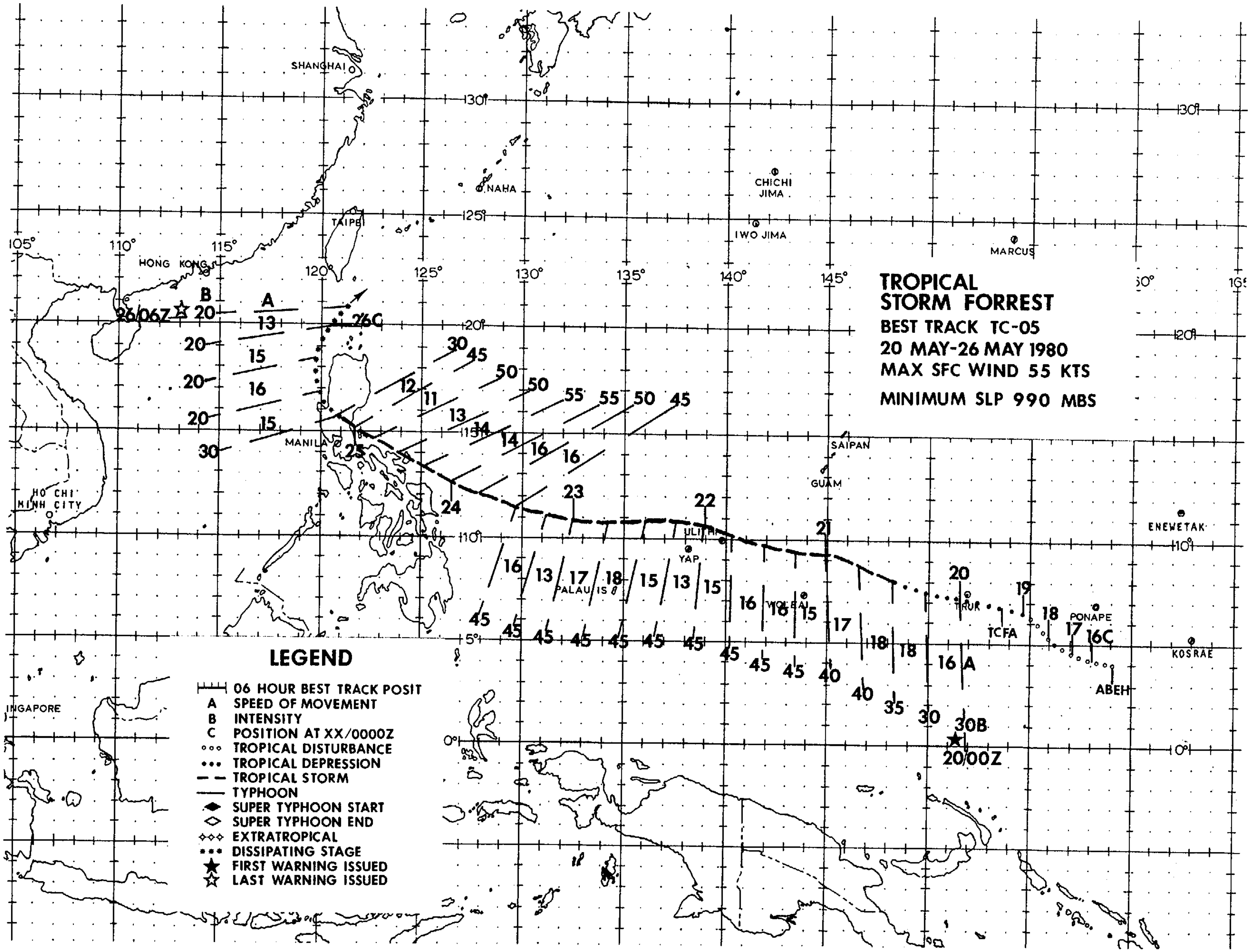


FIGURE 3-04-3. Forecast tracks for Typhoon Ellen from 191200Z and 191800Z data bases. Selected 24-hr forecast wind vectors at 500 mb for each data base are also illustrated, along with the final best track (▬▬) for that period.


TROPICAL
STORM FORREST
BEST TRACK TC-05
20 MAY-26 MAY 1980
MAX SFC WIND 55 KTS
MINIMUM SLP 990 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
TAIPEI
HONG KONG
MANILA
SAIPAN
GUAM
HO CHI MINH CITY
YAP
ULITHI
PALAU IS
WOLEAI
TRUK
PONAPE
KOSRAE
ENEWETAK
TCFA
ABEH
SINGAPORE
26/06Z
20/00Z

Tropical Storm Forrest was influenced by an unusually large and strong mid-tropospheric ridge which built westward across the Philippine Sea following the recurvature of Typhoon Ellen. This ridge dominated the entire northwestern Pacific and affected Forrest's direction of movement, forward speed, and intensity.

The majority of TS Forrest's track was spent skirting the southern periphery of the large subtropical ridge. During the month of May, cyclones typically track northwestward over the Philippine Sea. However, Figure 3-05-1 shows that nearly zonal 500 mb flow prevailed during this period and forced Forrest to track nearly due west. This steady zonal flow also pushed Forrest forward at speeds reaching 18 kt (33 km/hr), which is 3 times the climatological mean speed.

The strength of the subtropical ridge also affected Forrest's intensity. The subtropical ridge raised environmental pressures throughout the northwestern Pacific north of Forrest. Aircraft reconnaissance consistently observed winds in Forrest that were 10 to 15 kt (5 to 8 m/sec) stronger than would be expected from Forrest's minimum sea-level pressure of 990 mb and the Atkinson and Holliday (1977) pressure/wind relationship (Fig. 3-05-2). The Atkinson and Holliday relationship indicates that the 55 kt (28 m/sec) maximum sustained winds observed in Forrest (Fig. 3-05-3) are typically associated with tropical cyclones having a 983 mb minimum sea-level pressure. Aircraft reconnaissance also observed that Forrest tilted south-southwest from the surface to 700 mb. The surface and 700 mb centers were displaced as much as 35 nm (65 km) at times, apparently in response to upper-level northeast flow which existed over Forrest.

Following landfall on Luzon, Forrest weakened rapidly while passing approximately 40 nm (74 km) northeast of Clark AB. Highest observed wind speeds at this location associated with the passage of TS Forrest were in the 10 to 15 kt (5 to 8 m/sec) range.

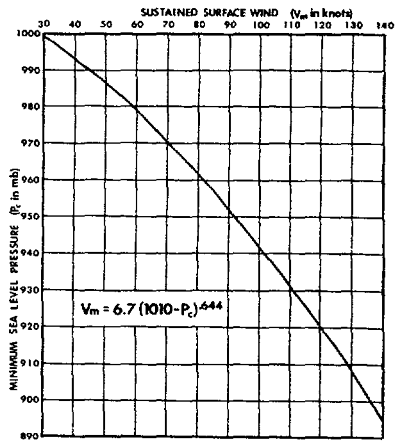


FIGURE 3-05-2. Atkinson and Holliday (1977) maximum sustained surface wind-minimum sea-level pressure relationship.

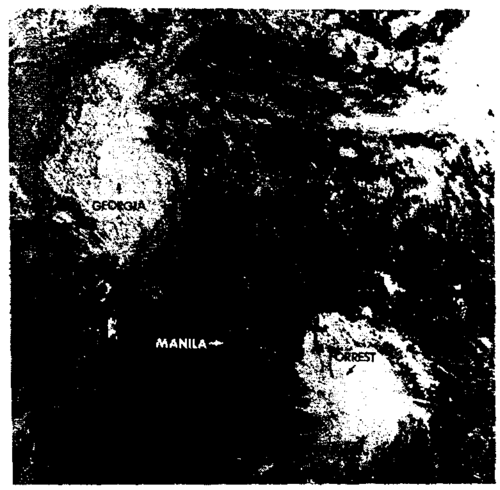


FIGURE 3-05-3. Tropical Storm Forrest at maximum intensity, 23 May 1980, 2344Z. Tropical Storm Georgia is making landfall over southeastern China. (NOAA 6 imagery)

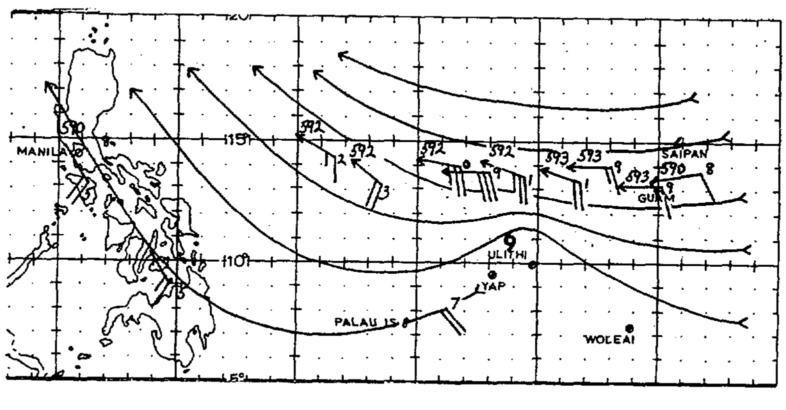
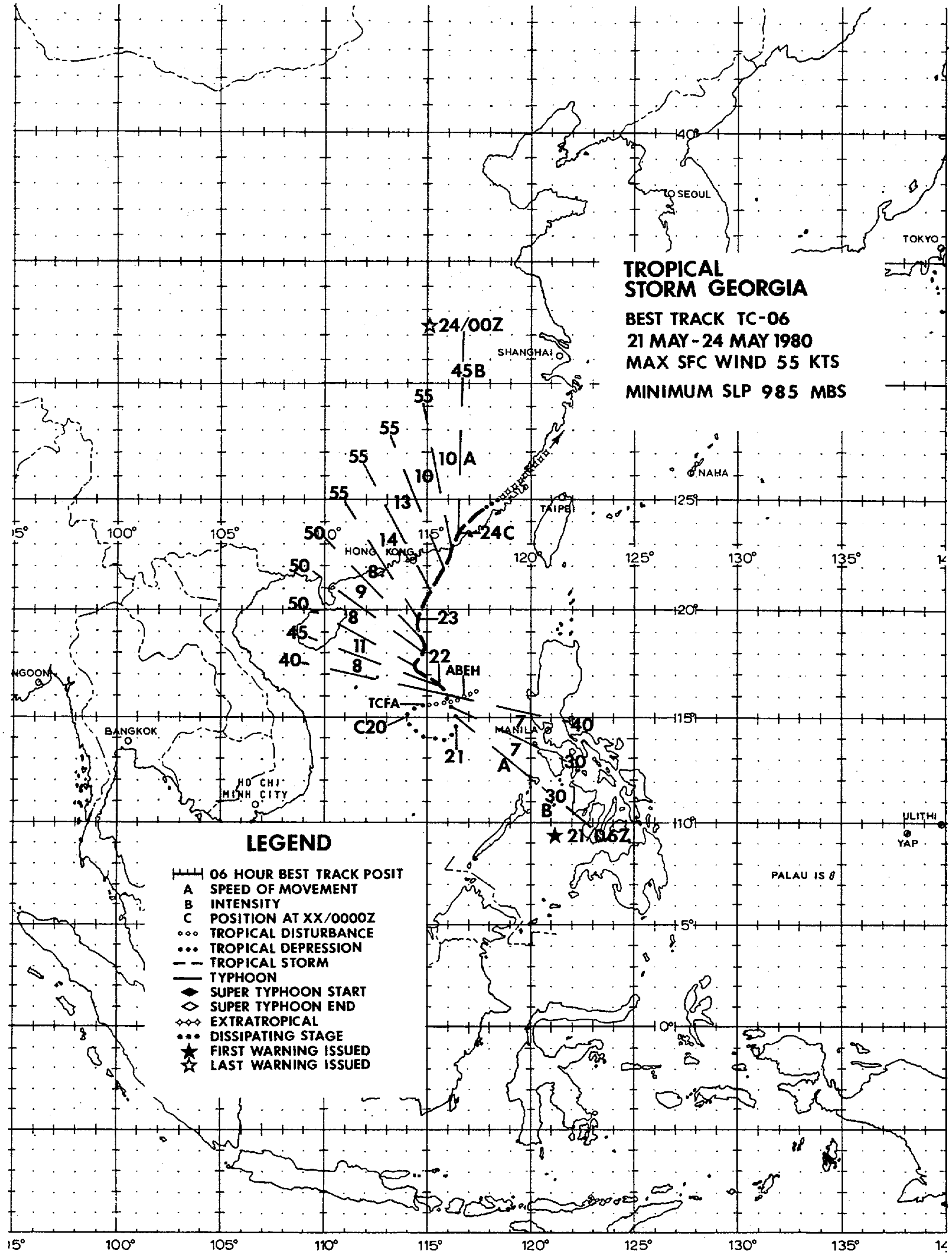


FIGURE 3-05-1. The 220000Z May 1980 streamline analysis of 500 mb rawinsonde (——) and aircraft reconnaissance (←—) data. The 500 mb heights are plotted in decameters and wind speeds are in knots.


TROPICAL STORM GEORGIA
BEST TRACK TC-06
21 MAY - 24 MAY 1980
MAX SFC WIND 55 KTS
MINIMUM SLP 985 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
24/00Z
21/06Z
SEOUL
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
BANGKOK
HO CHI MINH CITY
ULITHI
YAP
PALAU IS
TCFA
100°
105°
110°
115°
120°
125°
130°
135°

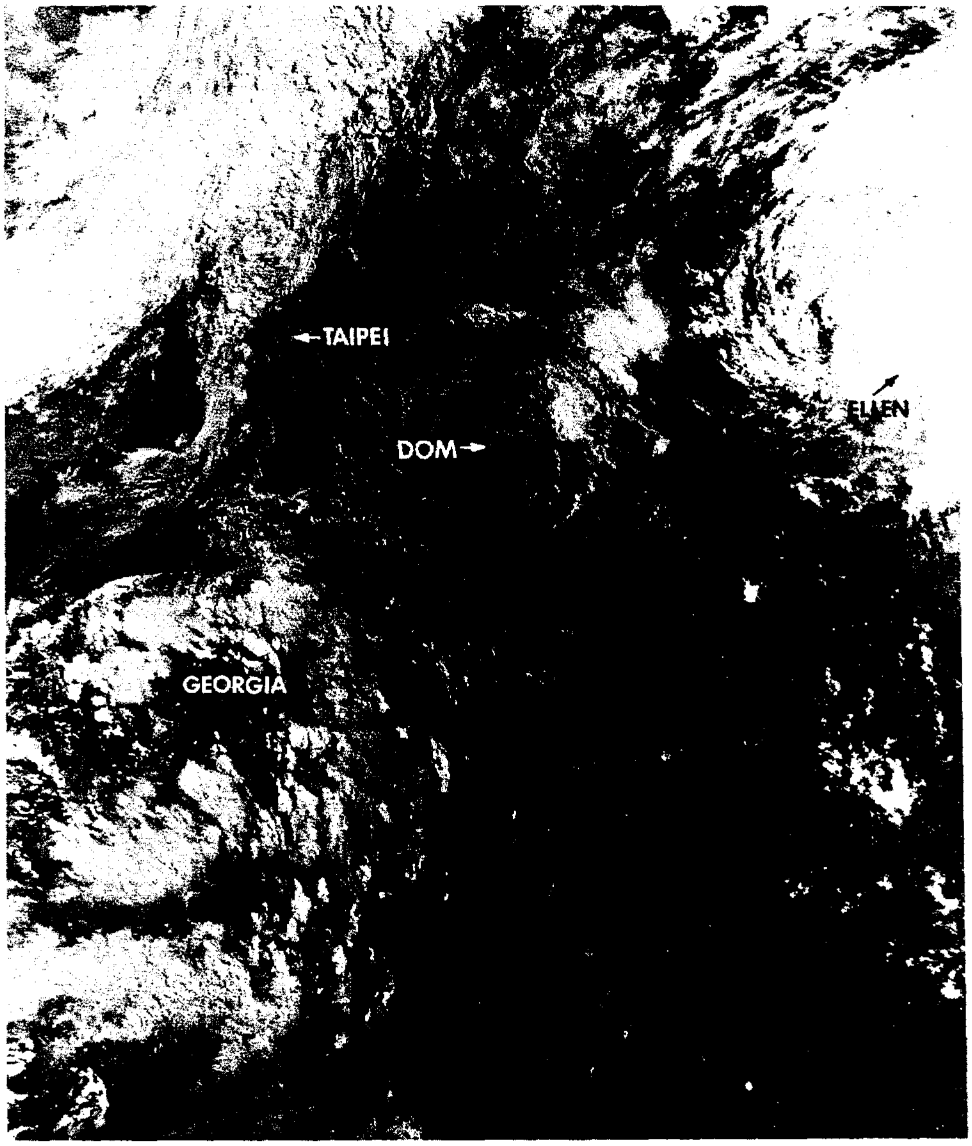


FIGURE 3-06-1. Typhoon Ellen, the remnants of Typhoon Dom, and the initial stage of Tropical Storm Georgia in the South China Sea, 19 May 1980, 2332Z. (NOAA6 imagery)

Tropical Storm Georgia is a classic example of a tropical cyclone which developed in the South China Sea during the transition period between the northeast and southwest monsoon. "Monsoon depressions" are often short-lived, difficult to locate with precision, and usually have broad, but relativily weak, surface circulation patterns. Georgia may well have reached typhoon strength if she had been able to remain over open water.

During the latter part of May, an active surface trough extended from near Iwo Jima southwestward into the South China Sea. Embedded in this trough were Typhoon Ellen, near Iwo Jima, the remnants of Typhoon Dom (an exposed low-level circulation), and the weak tropical disturbance which would become Tropical Storm Georgia (Fig. 3-06-1).

Synoptic data first indicated possible tropical cyclone development in the South

China Sea on the 19th of May. Although the satellite signature was poor, the synoptic data showed a surface circulation with a significant pressure drop near the center. Based on this data, a Tropical Cyclone Formation Alert (TCFA) was issued at 191341Z. Figure 3-06-2 is a surface streamline analysis at 200000Z and illustrates the well-defined surface circulation. The corresponding satellite imagery at about the same time still showed a lack of convective organization (Fig. 3-06-1). The depression finally began to show significant development, and the first warning on TD 06 was issued at 210600Z.

JTWC forecasters relied heavily on the forecast aids which were consistent in indicating northward movement with recurvature between the coast of China and the east coast of Taiwan. Southerly 500 mb steering winds also supported northward movement.

Figures 3-06-3 and 03-06-4 show the increase in surface inflow and the resulting increase in organized convective activity that occurred shortly after Georgia reached tropical storm strength.

Only two aircraft reconnaissance missions were flown on TS Georgia. The first

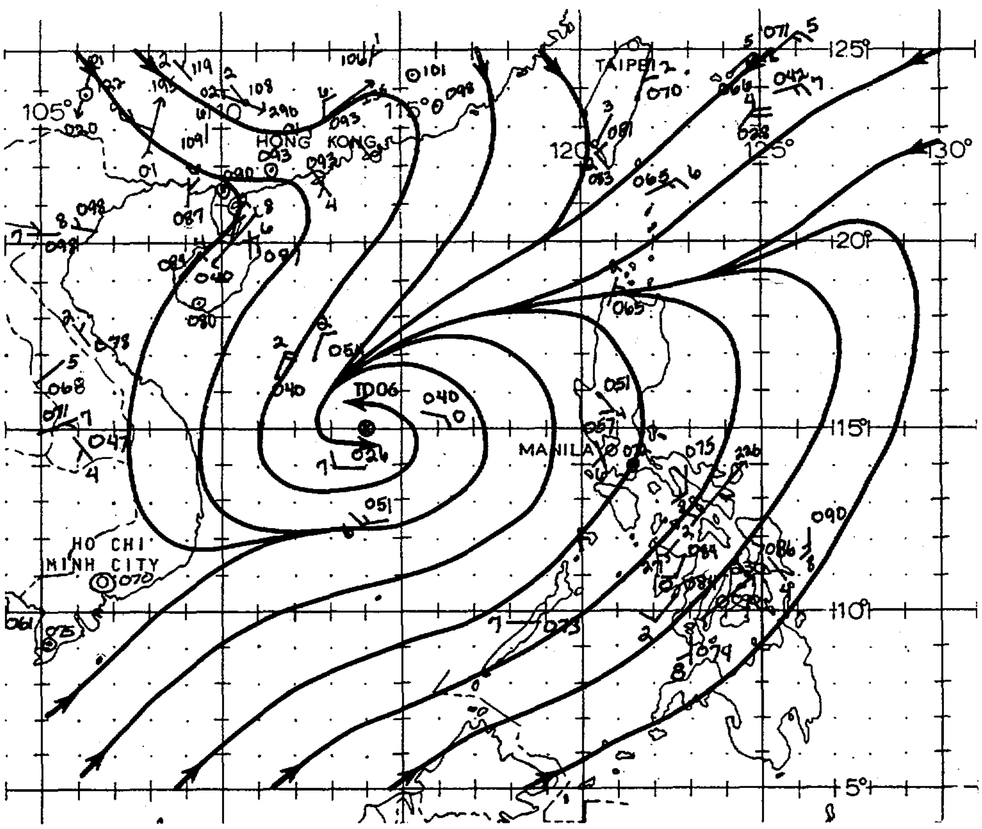


FIGURE 3-06-2. The 200000Z May 1980 surface (⟶)/ gradient (⟵) level streamline analysis in the vicinity of Tropical Storm Georgia. Winds are in knots.

mission observed a minimum sea-level pressure of 986 mb and surface winds of 50 kt (26 m/sec). The second mission could not provide a center fix because of restricted air space due to Georgia's proximity to both China and Hai-nan Island. Two ships, the "Clara Maersk" and the "Chevalier Paul", reported winds of 50 kt (26 m/sec) and 54 kt (28 m/sec), respectively. These observations support the best track estimated maximum intensity of 55 kt (28 m/sec).

After making landfall near Shan-tou, Georgia traveled north-northeastward, about 20 nm (37 km) inland from the coast, eventually passing north of Chin-men-tao, which reported winds of 44 kt (23 m/sec) with gusts to 60 kt (31 m/sec). Rapid weakening occurred thereafter as Georgia was absorbed into an extratropical low pressure system that was moving over the East China Sea from the Asian mainland.

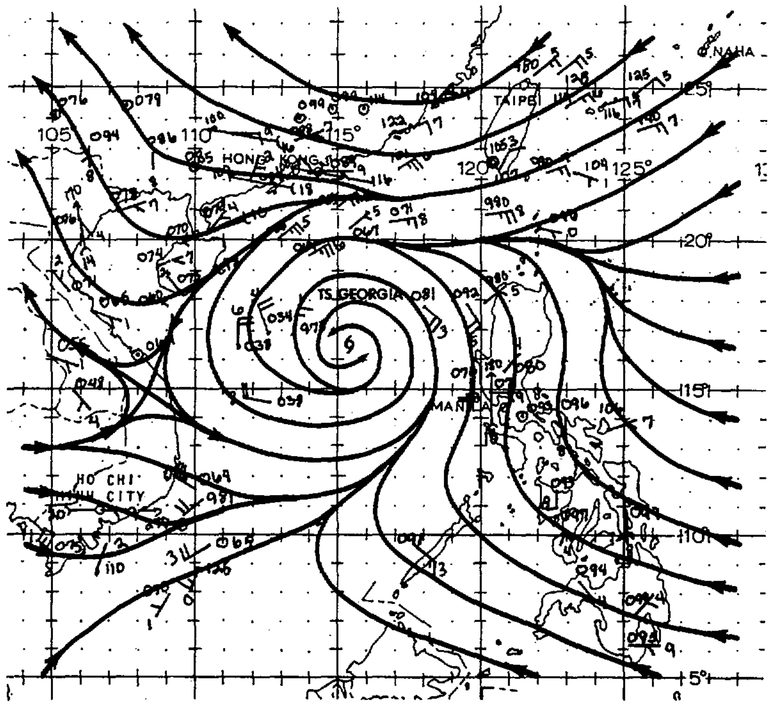


FIGURE 3-06-3. The 220000Z May 1980 surface (⟵⟍)/ gradient (⟵<) level streamline analysis in the vicinity of Tropical Storm Georgia. Winds are in knots.

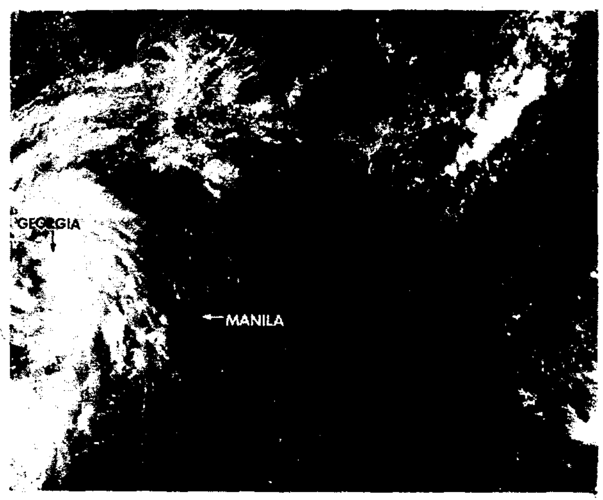
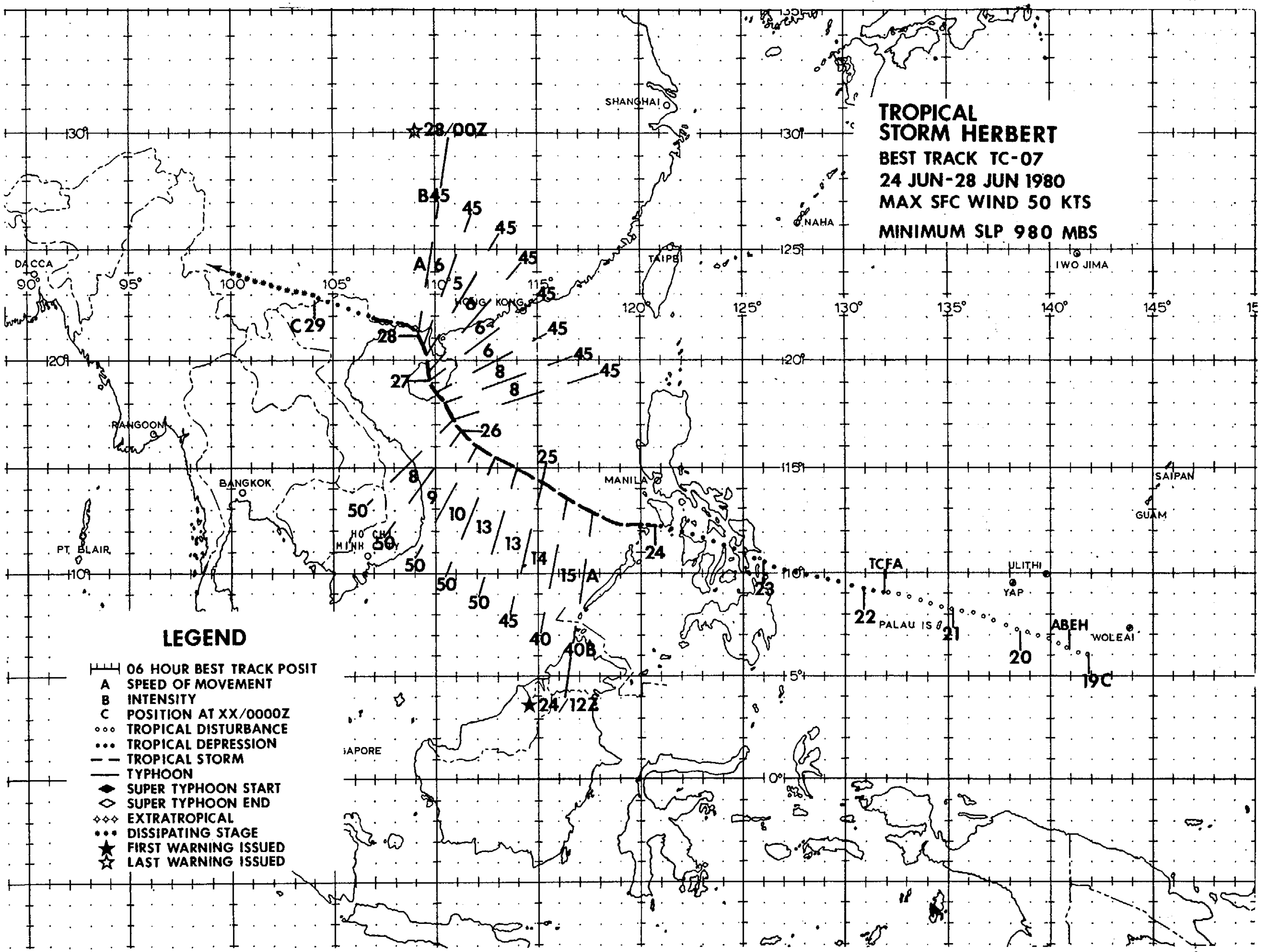


FIGURE 3-06-4. Tropical Storm Georgia at 45 kt (23 m/sec) intensity, 22 May 1980, 0158Z. (DMSP imagery)


TROPICAL
STORM HERBERT
BEST TRACK TC-07
24 JUN-28 JUN 1980
MAX SFC WIND 50 KTS
MINIMUM SLP 980 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
NAHA
TAIPEI
IWO JIMA
DACCA
HONG KONG
RANGOON
BANGKOK
MANILA
SAIPAN
GUAM
PT. BLAIR
ULITHI
YAP
PALAU IS
WOLEAI
TCFA
28/00Z
24/12Z
C29
28
27
26
25
24
23
22
21
20
19C
90°
95°
100°
105°
110°
115°
120°
125°
130°
135°
140°
145°

A broad equatorial trough existed on 18 June stretching from the Philippine Islands to the eastern Caroline Islands along 5N. Although synoptic data suggested several circulations along the trough axis, satellite imagery during the following 48 hours indicated increased convective activity around the eastern periphery of the trough as a result of convergent easterly flow.

At 201200Z, increased convection was noted near the primary surface circulation east of the Palau Islands. By 211200Z, satellite imagery indicated improved organization with synoptic data revealing increased southwest gradient level inflow and 20-25 kt (11-13 m/sec) wind reports from ships northeast of the depression. As a result, a tropical cyclone formation alert (TCFA) was issued at 211800Z.

The depression moved west-northwestward toward Leyte in the Philippine Islands on 22 June. The mountainous island chain was expected to prevent further development and the TCFA was cancelled at 221800Z. However, the potential for significant tropical cyclone development was expected to improve once again as the depression entered the South China Sea.

Thus, with the depression located south of Mindoro and moving west-northwestward, a formation alert was reissued at 240000Z. Aircraft reconnaissance at 240717Z located a circulation center just west of Busuanga Island with surface winds estimated at 40 kt (21 m/sec) and a minimum sea level pressure of 996 mb. Based on the aircraft data and evidence of increased convective activity on satellite imagery, the first warning on Tropical Storm Herbert was issued at 241200Z.

In the South China Sea, Herbert tracked northwestward toward Hai-nan Island while intensifying slowly. Maximum intensity of 50 kt (26 m/sec) was attained at 250600Z and was sustained for the next 24 hours as Herbert passed 15 nm (28 km) southwest of the Paracel Islands. Peak winds of 46 kt (24 m/sec) were reported by the islands at 260000Z. Landfall on Hai-nan occurred near 261800Z with maximum sustained winds of 45 kt (23 m/sec). Over Hai-nan, Herbert tracked around the western face of Wu Chih Sham Mountain and exited due north into the Gulf of Tonkin. A north-northwest track over the Gulf of Tonkin ended with landfall south of Chin-hsien, China at 280300Z with 45 kt (23 m/sec) intensity, as verified by land station reports. Once over southern China, Herbert weakened quickly and dissipated as a significant tropical cyclone by 290000Z (Fig. 3-07-1).

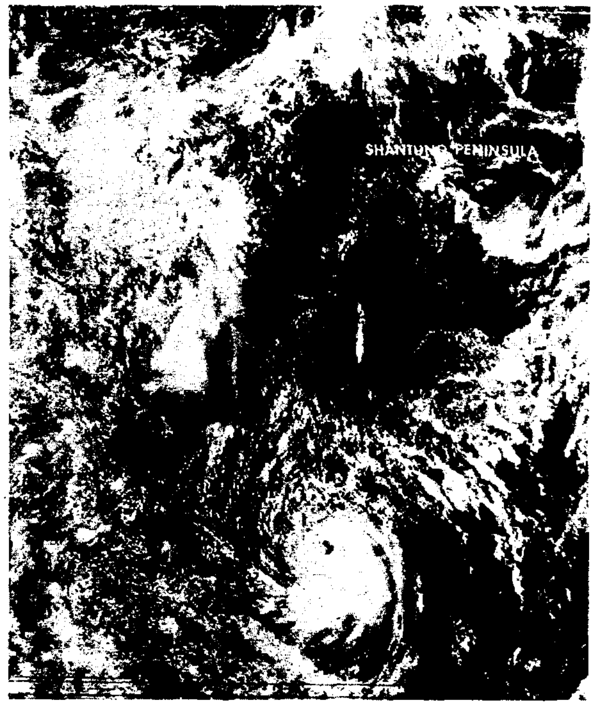


FIGURE 3-07-1. *Tropical Storm Herbert at 45 kt (23 m/sec) intensity making landfall over southern China, 28 June 1980, 0316Z. (DMSP imagery)*

A strong mid-level ridge extending from southern China eastward across the Pacific along 24N provided the steering flow as Herbert tracked steadily along the southern periphery of the ridge. The 500 mb analyses on 25 and 26 June showed that the ridge extended westward to near 108E just west of Hai-nan Island. Thus, a turn toward the northeast was expected following landfall over southern China. However, the 270000Z 500 mb analysis revealed that the ridge actually built westward across southern China, resulting in Herbert's westward track during his dissipation stage following landfall.

The definitive mid-level synoptic pattern and steering flow provided JTWC with good warning continuity and resulted in excellent forecast vector errors of 77 nm (143 km), 128 nm (237 km), and 57 nm (106 km) for 24, 48, and 72 hours, respectively.

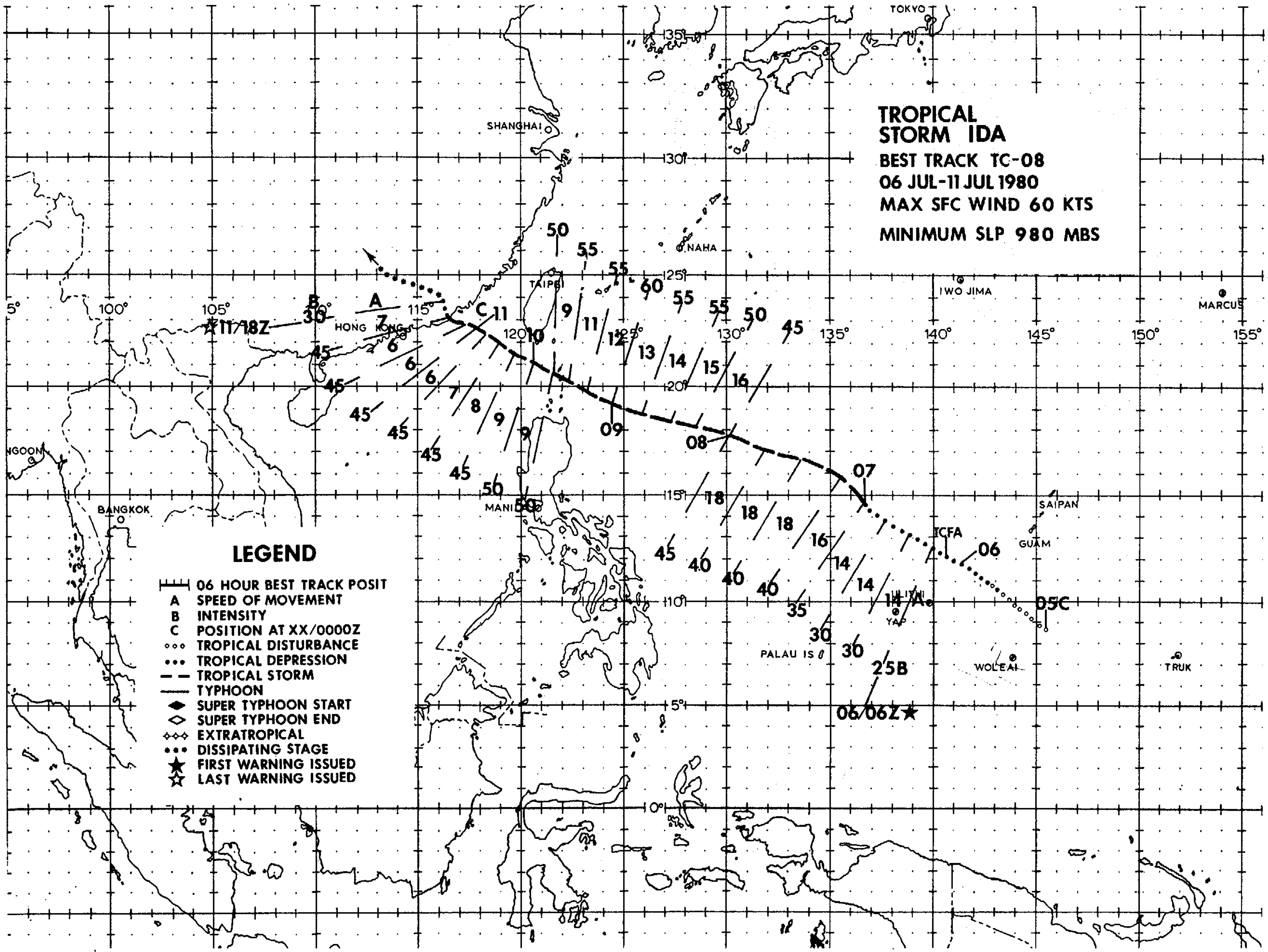

TROPICAL
STORM IDA
BEST TRACK TC-08
06 JUL-11 JUL 1980
MAX SFC WIND 60 KTS
MINIMUM SLP 980 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
TOKYO
SHANGHAI
NAHA
TAIPEI
IWO JIMA
MARCUS
HONG KONG
BANGKOK
MANILA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
TCFA
11/18Z
06/06Z
05C
06
07
08
09
25B

FIGURE 3-08-1. Tropical Storm Ida intensifying in the Philippine Sea. The discontinuous outer rainband (clockwise from the northwest to the east) was evident from the cyclone's initial development until weakening near the Bashi Channel, 6 July 1980, 2236Z. (NOAA visual imagery)

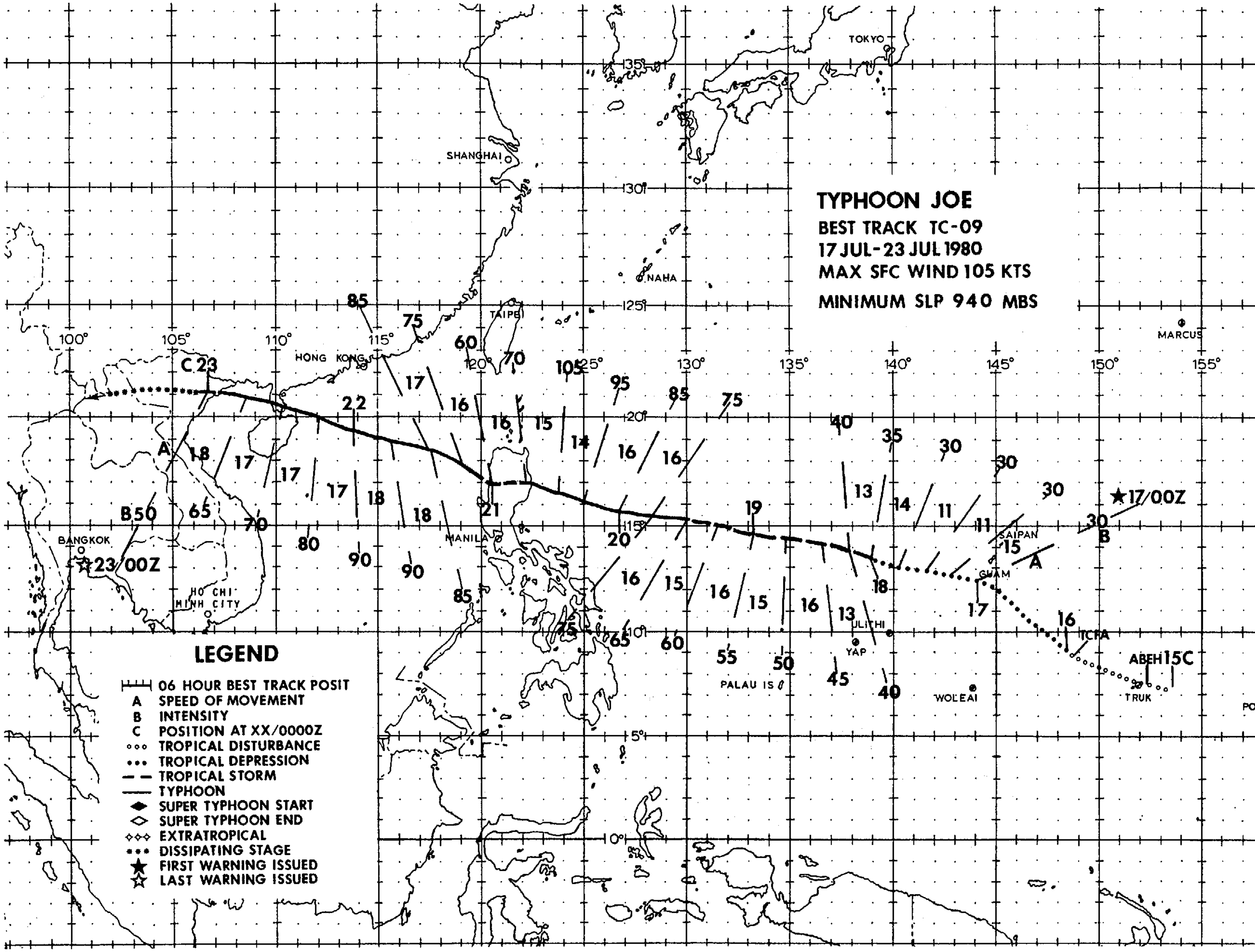
During the first three weeks of July, the monsoon trough extended eastward from the South China Sea to near 160E. Areas of active convection were common during that period with several disturbances eventually developing into significant tropical cyclones (Ida, Joe and Kim).

TD 08, the first of these disturbances which became organized, formed in the vicinity of two areas of active convection which JTWC had been tracking within the trough for several days. At 051800Z, satellite imagery indicated improved convective organization about a surface circulation near 12N 142E. Aerial reconnaissance at 060145Z located a surface center and observed maximum winds of 25 kt (13 m/sec). At 060300Z, a Tropical Cyclone Formation Alert (TCFA) was issued and the initial warning followed at 060600Z based on continued organization as indicated by satellite imagery.

Ida's track never posed a dilemma for JTWC forecasters. She initially tracked northwest before interacting with a persistent ridge, whose axis was along 28N. Maintenance of the ridge throughout Ida's lifespan was responsible for the cyclone's overall west-northwest track and the ability of JTWC to predict landfall within 35 nm (65 km) of the actual point as early as 77 hours prior to the occurrence. As Ida approached the western Philippine Sea, her forward movement slowed from a maximum of 18 kt (33 km/hr) to less than 10 kt (18 km/hr) in the Bashi Channel. During this period, Ida reached her maximum intensity of 60 kt (31 m/sec) and lowest sea level pressure of 980 mb. As Ida moved through the Bashi Channel, she weakened to 45 kt (23 m/sec) and then maintained this intensity until making landfall on the southeastern coast of mainland China, just south of Shan-t'ou (WMO 59316) at 1300Z on 11 July.

A predominate feature during all but the later stages of Ida's track was a strong and persistent rain or feeder band. Figure 3-08-1 shows this feature on NOAA satellite imagery. On 6 July, while Ida was organizing southwest of Guam, the Naval Air Station at Agana, Guam recorded 1.15 inches (29 mm) of rain in 3 hours and a peak gust of 31 kt (16 m/sec) as the band passed over the island. Subsequent aircraft reconnaissance in support of the 072130Z and 090735Z position fixes noted increases in the flight level winds and temperatures while transiting this outer band. These higher flight-level temperatures and winds were an apparent response to the release of latent heat of condensation from extensive convection within the band. The mission ARWO[1] on the 090735Z fix stated, "The 700 mb center was very weak with no strong wind band in any quadrant close to the center. It took us quite a while to locate a fairly broad area of light and variable winds...The northern end of this area was open, and to the northeast we observed a broad band of surface winds peaking at 85 kt (44 m/sec) situated about 25 nm (46 km) out." This was not an uncommon feature from aircraft reconnaissance data on Ida. Generally these maximum winds were located in the northeast quadrant, close to the rainband. However, the 85 kt (44 m/sec) wind was considered to be a transistory feature because all other indicators showed no reason for a sudden and short-lived intensification of the cyclone.

[1]CHARLES B. STANFIELD, Capt, USAF: Mission Aerial Reconnaissance Weather Officer (ARWO).


TYPHOON JOE
BEST TRACK TC-09
17 JUL-23 JUL 1980
MAX SFC WIND 105 KTS
MINIMUM SLP 940 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
BANGKOK
HO CHI MINH CITY
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
MARCUS
TCFA
ABEH
17/00Z
23/00Z

Typhoon Joe, the ninth tropical cyclone in the Western Pacific region, proved to be very predictable. A near static synoptic pattern prevailed in the mid- to upper-troposphere within the subtropics throughout most of Joe's existence. Figures 3-09-1 and 3-09-2 show the distinguishable traits in the structure of the mid- and upper-troposphere. As a result of this pattern, Joe followed a nearly straight track from genesis to dissipation with few exceptions.

Joe's genesis from a tropical disturbance into a mature tropical cyclone was slow. Satellite imagery first indicated a disturbance along the equatorial trough on 14 July over the Caroline Islands. Later satellite data revealed a gradual increase of convective activity with an apparent increase in organization. As a result of the information received from this series of satellite imagery, a tropical cyclone formation alert (TCFA) was issued at 2153Z on the 15th.

The first aircraft reconnaissance of the disturbance on the 16th found a weak surface circulation which did not extend up to the 700 mb level. At that time the minimum sea level pressure was 1006 mb and the disturbance was tracking northwestward at 14 kt (26 km/hr). Defense Meteorological Satellite Program (DMSP) imagery at 0021Z on the 17th suggested that the disturbance was developing a circulation center that extended at least to mid-tropospheric levels (Fig. 3-09-3) with strong convective activity located west of the exposed surface circulation. As discussed by Huntley and Diercks (1980), weak developing tropical cyclones often have the 700 mb center displaced from the surface circulation in the direction of strongest convective activity. As the tropical cyclone develops and intensifies, the surface circulation moves under the 700 mb center and becomes vertically aligned. Later satellite data did show that the surface center had moved closer to the area of strong convection. This sequence of events prompted JTWC to issue the first warning at 0000Z on 17 July for Tropical Depression 09. Aircraft reconnaissance on the 17th substantiated that the disturbance had indeed developed significantly since the 16th and that TD 09's circulation center had extended up to the 700 mb level with no significant displacement of the surface and 700 mb center noted at that time.

FIGURE 3-09-3. Typhoon Joe during his early stage of development, 17 July 1980, 0021Z. Arrow shows location of exposed low-level circulation center. (DMSP imagery)

FIGURES 3-09-1 and 3-09-2 are on the following page.

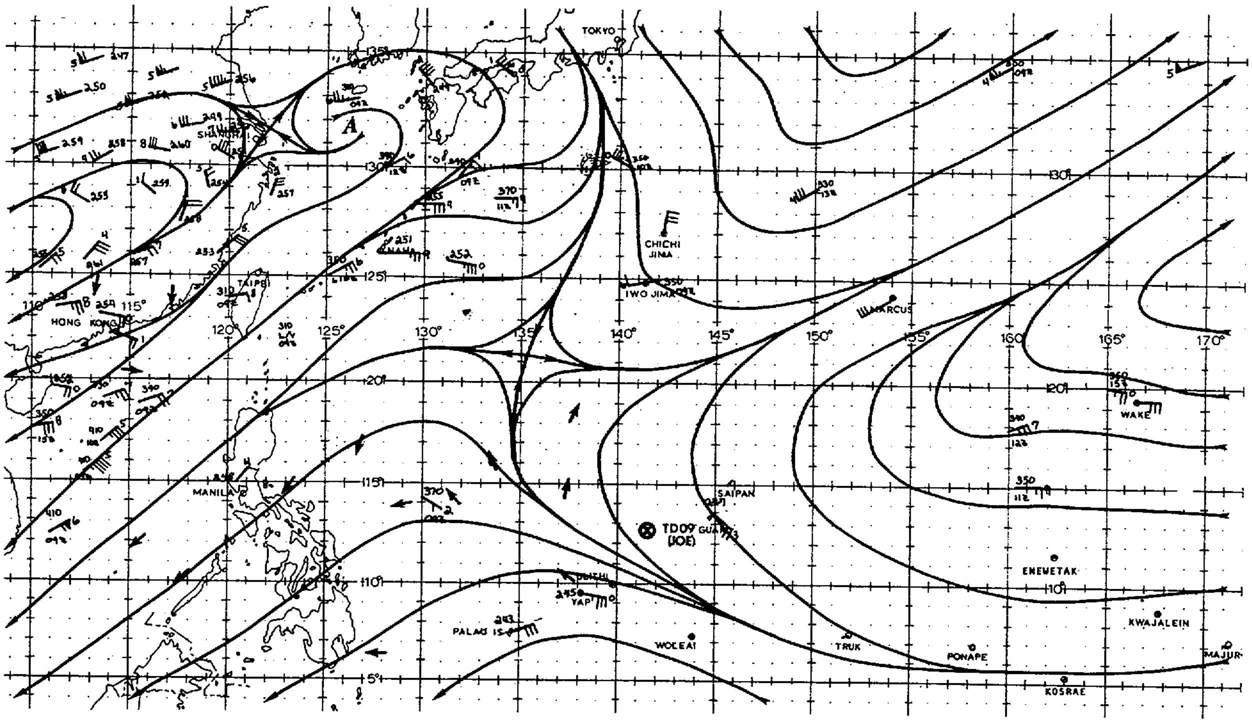


FIGURE 3-09-1. 200 mb streamline analysis at 171200Z July 1980. The analysis depicts the synoptic pattern which prevailed during much of Typhoon Joe's existence. Wind data are a combination of RAOBS, AIREPS, and satellite-derived (+) winds for the 250 mb to 150 mb levels. Wind speeds are in knots.

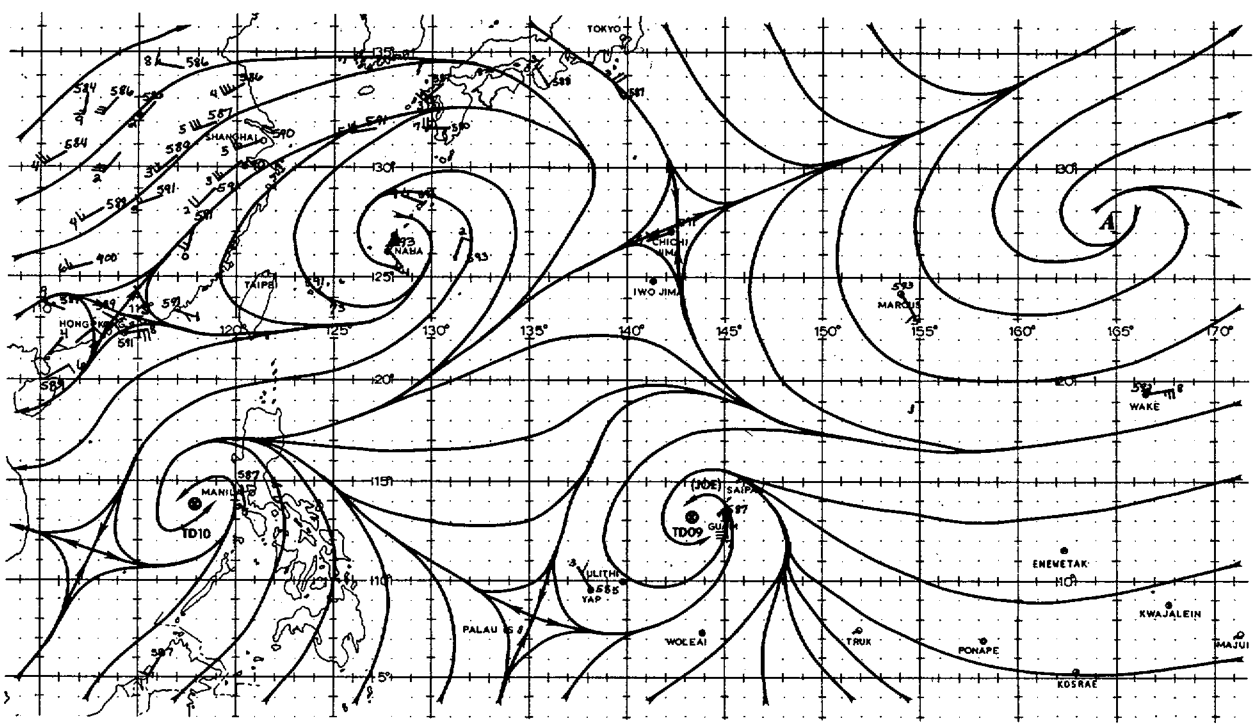


Figure 3-09-2. 500 mb streamline analysis at 170000Z July 1980. The analysis depicts the synoptic pattern which prevailed during much of Typhoon Joe's existence. Wind speeds are in knots.

TD 09 developed rapidly from that point and was upgraded to Tropical Storm Joe on the 18th. Typhoon strength was attained on the 19th.

As mentioned earlier, Joe tracked along a nearly straight course through much of his existence. His forward speed of movement was rapid and nearly constant, even while passing over Luzon. This unusually persistent track and high speed of movement was correlated with an abnormally strong mid- and upper-tropospheric subtropical ridge. The subtropical ridge at both levels deviated significantly from the climatological norm. The 200 mb anticyclone, normally located over the China mainland, extended further east to the north of Joe's track and south of Japan (Fig. 3-09-1). Similarly, the mid-tropospheric ridge was to the north of Joe's track and was much stronger than normal (Fig. 3-09-2).

This pattern did not significantly change during Joe's lifetime, except briefly while he was first developing into a tropical depression. Joe's track took a slight northwestward jog in response to a short wave trough which weakened the mid- to upper-tropospheric ridge. This short wave trough quickly passed eastward and the ridge built back north of Joe.

Six hours prior to landfall over Luzon, Joe attained an intensity of 105 kts (54 m/sec) with a minimum sea level pressure of 940 mb at 1200Z on 20 July. Joe weakened rapidly to tropical storm strength while crossing Luzon, but still remained very destructive. As he tracked across the mountainous terrain of Luzon, where peaks approach 10,000 ft, the track deviated slightly, becoming more westward. It took just over 6 hours for Joe to cross Luzon, but in that short time, the Philippine Islands were inundated by heavy rains which produced massive flooding and resulted in extensive crop and property damage.

Approximately 177,000 people were left homeless and 19 deaths were reported. Exact figures could not be compiled in time due to Typhoon Kim which hit the Philippines within a week of Joe, compounding destruction that the Philippines had already suffered. No significant damage was reported to U. S. military installations in the Philippines.

Upon entry into the South China Sea, Joe reintensified to typhoon strength. Before this time, JTWC expected Joe to track northwest onto the Asian mainland about 100 nm (185 km) west of Hong Kong and dissipate. The mid- and upper-tropospheric ridge, however, extended westward, causing Joe to continue on a west-northwest track toward Vietnam. Also, from the time Joe entered the South China Sea through dissipation, he maintained a rapid speed of movement due to the strong ridge to the north. Typhoon Joe attained a second maximum intensity of 90 kt (46 m/sec) as determined by Dvorak analysis of satellite data (Fig. 3-09-4). At the time of maximum intensity, the radius of winds greater than 30 kt (15 m/sec) extended 450 nm to the east of Joe's center, covering most of the South China Sea north of 10N. While transiting across the South China Sea, Joe devastated the coastal regions which paralleled his track. Much damage to crops and property occurred in southern China due to flooding caused by torrential rains. Joe also left many homeless and claimed more lives while tracking toward Vietnam.

Satellite imagery showed that Joe had an eye as he made landfall near Haiphong, Vietnam. During this period, winds were reported in excess of 70 kt (36 m/sec) by the Vietnam News Agency. After landfall, Joe dissipated rapidly due to land and vertical wind shear effects. The final warning was issued by JTWC at 0000Z on 23 July as the remnants of Joe began to dissipate over the mountains of Laos.

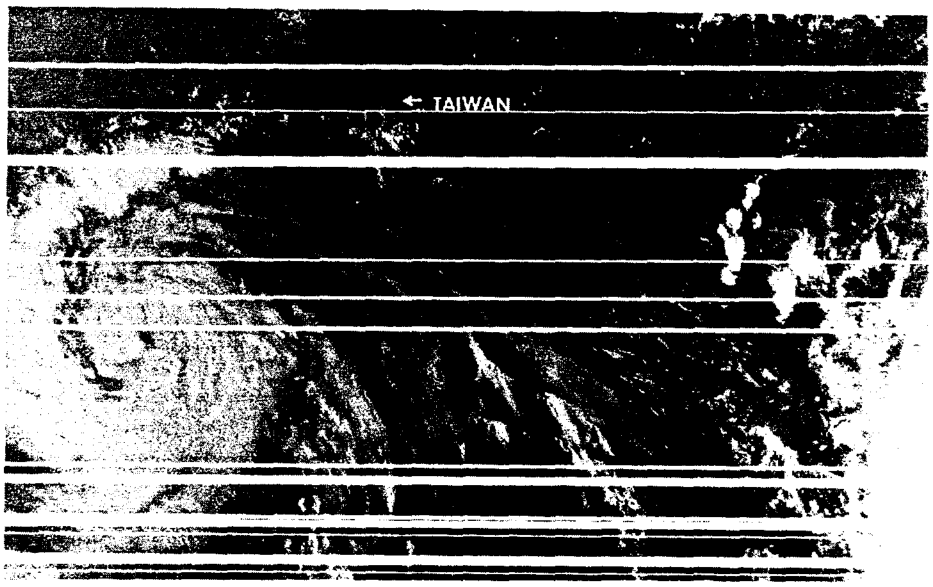
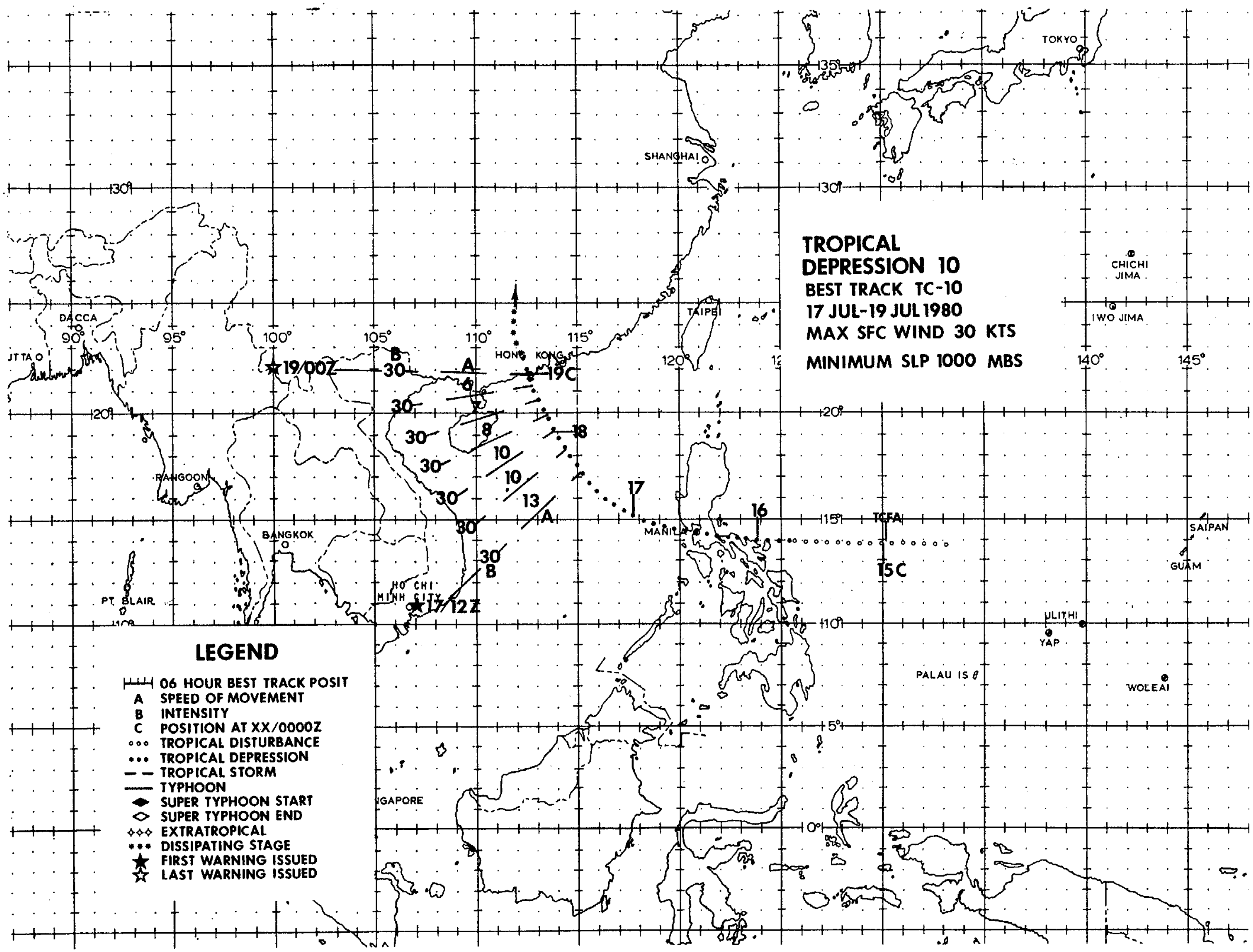


FIGURE 3-09-4. *Typhoon Joe after reintensifying to 90 kt (46 m/sec) in the South China Sea, 21 July 1980, 2347Z. (DMSP imagery)*


TROPICAL
DEPRESSION 10
BEST TRACK TC-10
17 JUL-19 JUL 1980
MAX SFC WIND 30 KTS
MINIMUM SLP 1000 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
19/00Z
17/12Z
19C
18
17
16
15C
TCFA
TOKYO
SHANGHAI
TAIPEI
HONG KONG
MANILA
BANGKOK
HO CHI MINH CITY
RANGOON
DACCA
PT. BLAIR
CHICHI JIMA
IWO JIMA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
90°
95°
100°
105°
110°
115°
120°
140°
145°
135°
130°
20°
15°
10°
5°

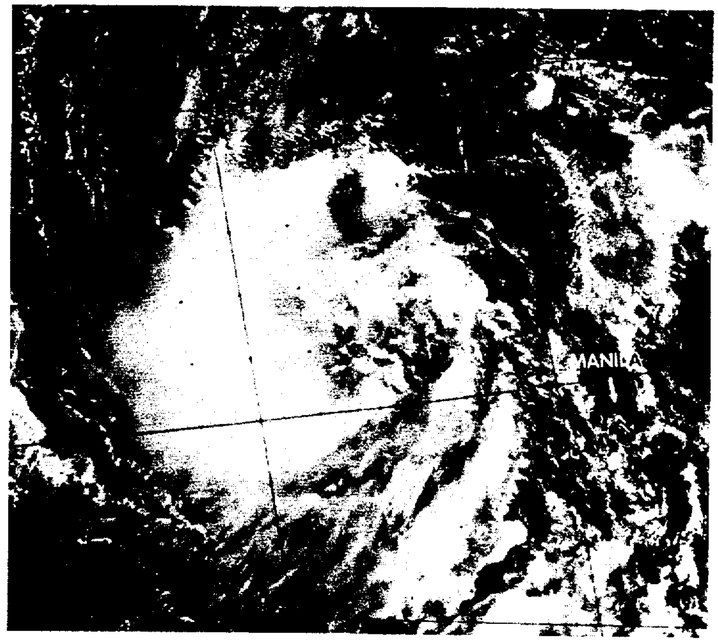


FIGURE 3-10-1. Tropical Depression 10 220 nm (407 km) west-northwest of Manila 13 hours prior to issuance of first warning. Heavy convection is located two degrees west of the surface circulation, 17 July 1980, 0203Z. (DMSP imagery)

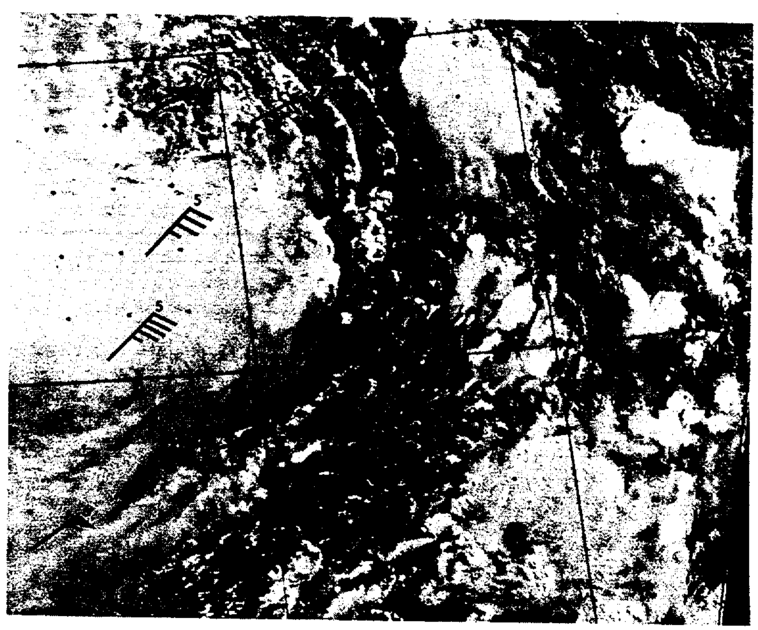
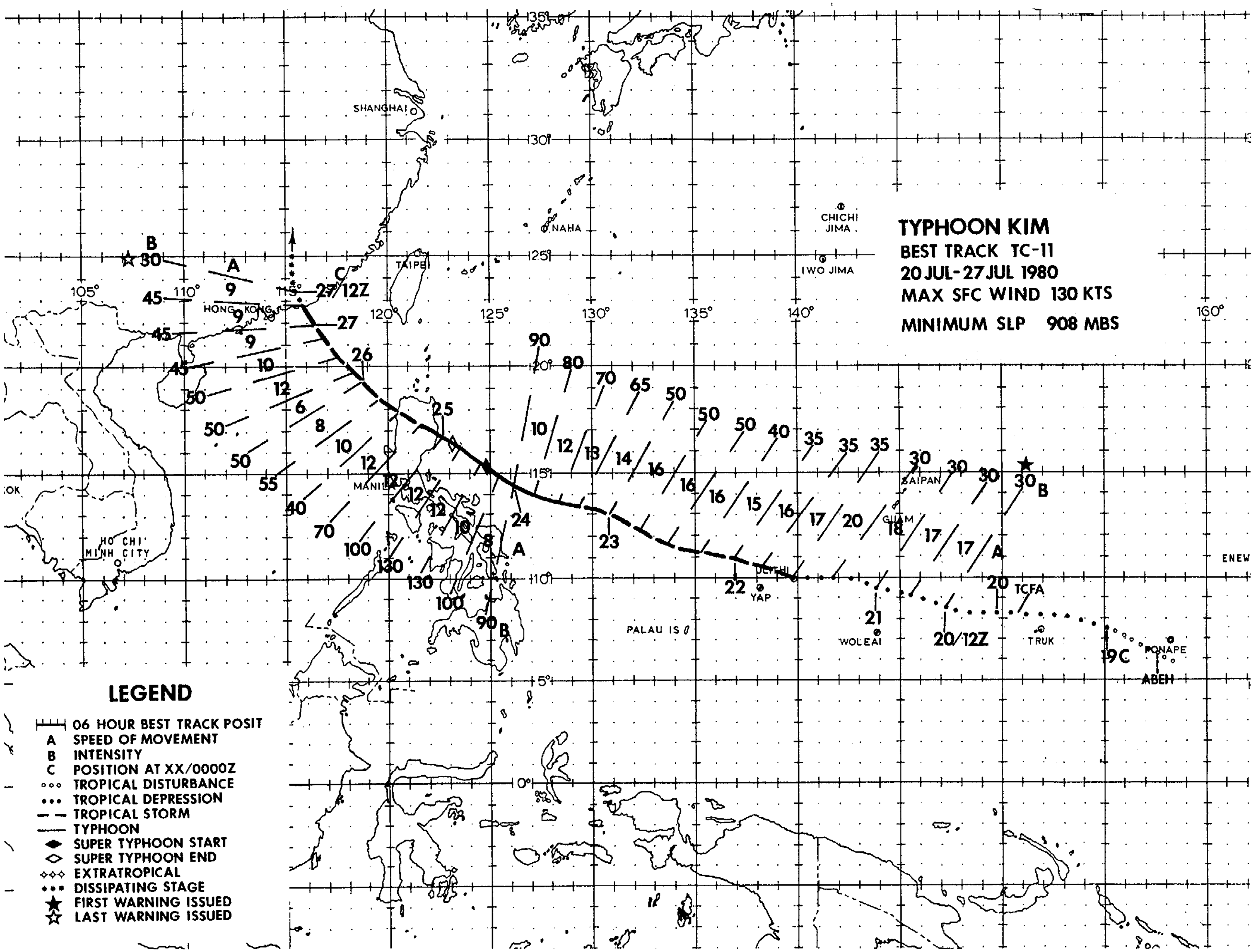


FIGURE 3-10-2. Tropical Depression 10 215 nm (398 km) south of Hong Kong with low-level circulation partially exposed due to strong vertical wind shear, 17 July 1980, 2334Z (NOAA6 imagery). Wind barbs represent aircraft and radiosonde reports near the 200 mb level at 171200Z.


TYPHOON KIM
BEST TRACK TC-11
20 JUL-27 JUL 1980
MAX SFC WIND 130 KTS
MINIMUM SLP 908 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
TAIPEI
NAHA
CHICHI JIMA
IWO JIMA
HONG KONG
MANILA
HO CHI MINH CITY
SAIPAN
GUAM
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
TCFA
27/12Z
20/12Z
19C
105°
110°
115°
120°
125°
130°
135°
140°
160°

Super Typhoon Kim, one of the most intense typhoons of the 1980 season, slammed onto the eastern coast of Luzon four days after Typhoon Joe had practically immobilized the area. Accounts of the aftermath of Typhoon Kim indicated that an estimated 15 people were killed and 167,000 residents of the Philippines were displaced. Torrential rains caused massive flooding over Luzon as far south as Manila

Kim, the first super typhoon of the 1980 season, was first detected on satellite imagery on 19 July. The disturbance appeared as an area of enhanced convection embedded in the near-equatorial trough. Further intensification appeared likely as the tropical upper-tropospheric trough (TUTT) was positioned to the northwest of the convective area. Because the disturbance was in a favorable position for continued development, a Tropical Cyclone Formation Alert (TCFA) was issued at 192040Z. Aircraft reconnaissance data at 200800Z indicated a well-defined closed surface circulation with wind speeds of 25 to 30 kt (12 to 15 m/sec) and a central pressure of 1001 mb approximately 360 nm (667 km) southeast of Guam. Based on this data, the first warning on TD 11 was issued at 201200Z.

TD 11 initially moved west-northwestward passing approximately 240 nm (444 km) south of Guam before heading directly towards the island of Ulithi. At 211200Z, TD 11 passed directly over Ulithi, which reported a wind maximum of 35 kt (18 m/sec). This information, plus a subsequent aircraft report of a central surface pressure of 997 mb, prompted JTWC to upgrade TD 11 to Tropical Storm Kim at 211800Z. Aircraft data at that time, however, indicated that Kim was poorly alig-

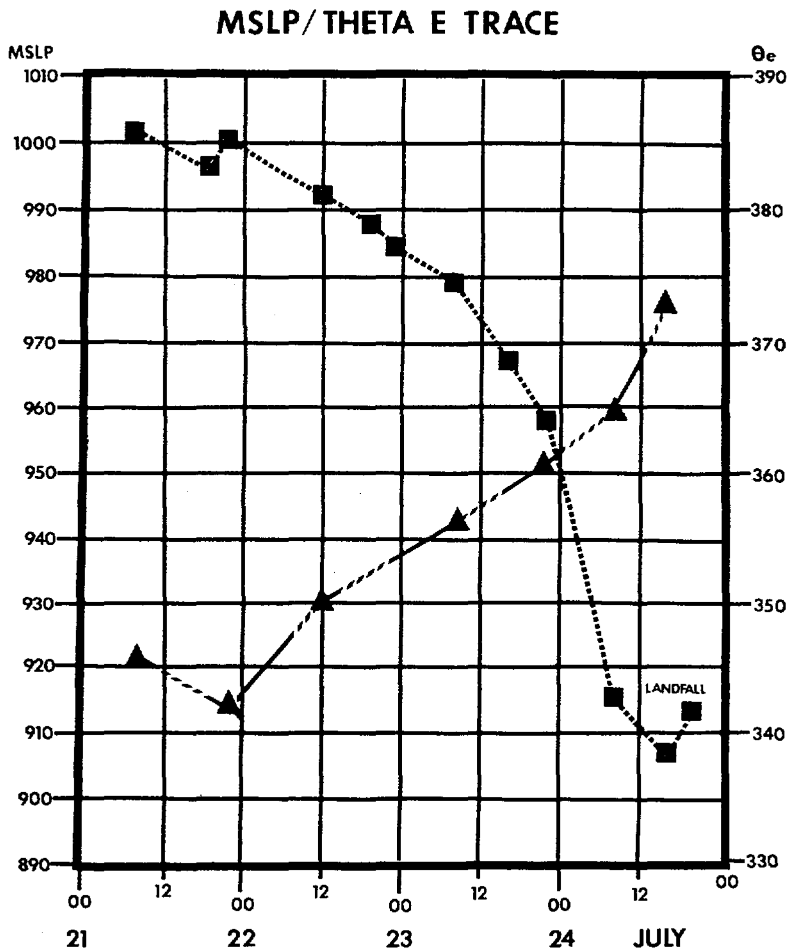


FIGURE 3-11-1. Time cross-section of Kim's minimum sea-level pressure versus 700 mb equivalent potential temperature (THETA E (θ_e)) as derived from aircraft reconnaissance data.

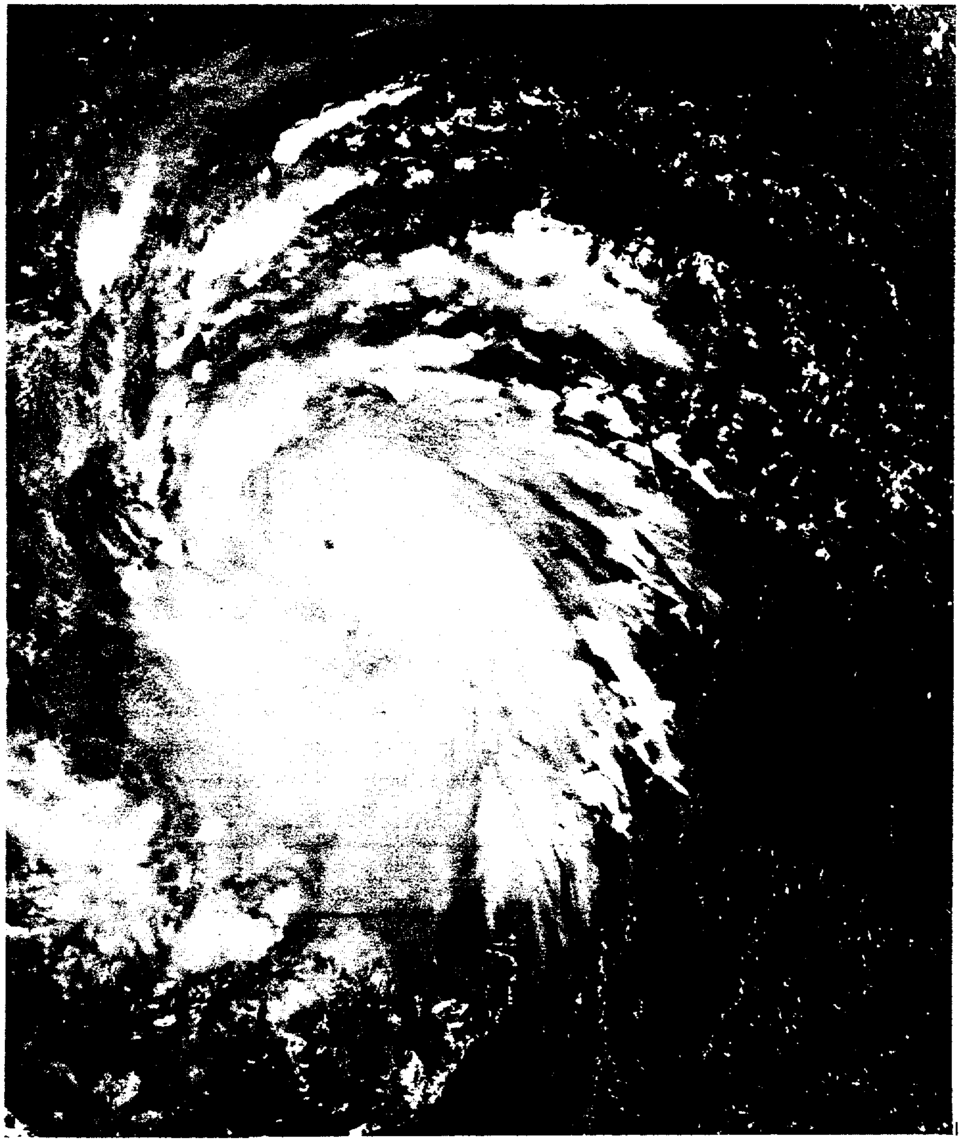


FIGURE 3-11-2. Typhoon Kim at approximately 110 kt (57 m/sec) intensity before she reached super typhoon strength, 24 July 1980, 0125Z. (DMSP imagery)

ned in the vertical, with the 700 mb center well to the southeast of the surface center and the 700 mb wind flow largely disorganized. Therefore, further intensification was slow during the 22nd and 23rd. During this period, Kim followed a path similar to Typhoon Joe across the Philippine Sea, tracking west-northwestward along the southern periphery of the subtropical mid-tropospheric ridge.

At 230600Z, aircraft reconnaissance observed a fairly substantial drop in surface pressure to 979 mb and indications that an eyewall was partially forming. Upon receipt of the data, which signalled the beginnings of a period of more rapid intensification, Kim was upgraded to a typhoon.

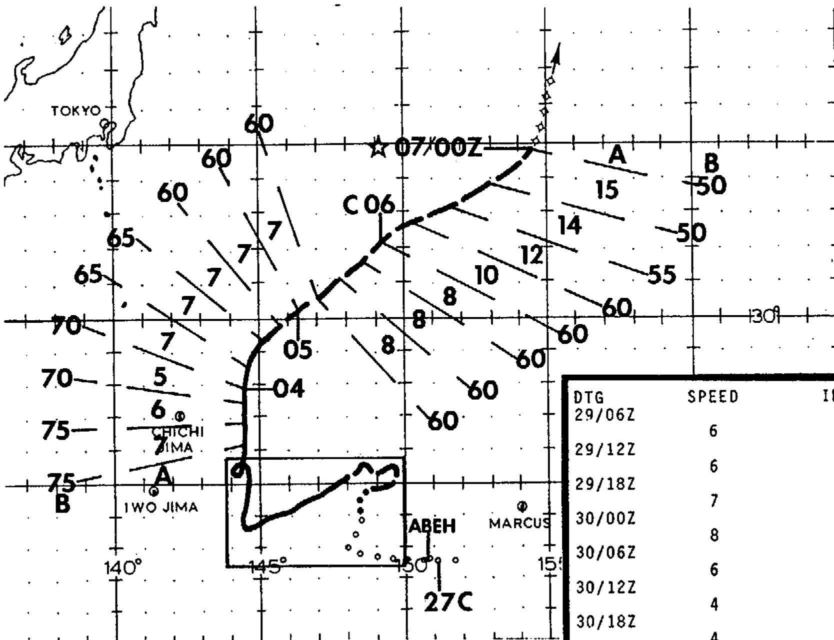
During this period of falling pressures and corresponding intensification, an empirically derived forecasting aid (Fig. 3-11-1) proved very valuable to JTWC. This forecasting aid relates surface pressure and 700 mb equivalent potential temperature (θ_e) to future intensification. The hypothesis is that rapid intensification is likely to take place in a tropical cyclone within the next 12 to 36 hours after these two traces intersect. Typhoon Kim's intensification trend verified this study.

At 241603Z, a minimum sea level pressure of 908 mb was measured by dropsonde. This pressure was sufficiently low to qualify Kim as a super typhoon (Fig. 3-11-2). By the next aircraft penetration, however, Kim's central pressure had risen to 918 mb. A possible reason for this rise in pressure was that Kim was now only eight hours from landfall on the coast of Luzon and the mountainous terrain had begun to disturb Kim's low-level inflow. Shortly afterwards, at about 250000Z, Typhoon Kim moved onto the coast of Luzon (Fig. 3-11-3) with accompanying maximum sustained winds of 100 kt (52 m/sec) and reported wind gusts as high as 125 kt (64 m/sec).

Terrain further weakened Kim as she moved slowly across Luzon before emerging in the South China Sea as an ill-defined tropical storm. JTWC forecasters expected Kim to reintensify as a typhoon over the South China Sea similar to Joe only several days earlier. Aircraft reconnaissance, however, continued to report that Kim lacked significant organization and that her associated convective tops were significantly lower than previously observed.

A weakness in the mid-tropospheric ridge, thought to have been induced by Typhoon Joe's passage several days earlier, allowed Kim to track more northwest towards Hong Kong. Changing little in direction or intensity as she tracked across the South China Sea, Kim finally made landfall on the coast of China 90 nm (167 km) northeast of Hong Kong at about 270600Z. Maximum sustained winds of 45 kt (23 m/sec) and wind gusts to 60 kt (31 m/sec) were reported as Kim moved inland.

FIGURE 3-11-3. Typhoon Kim over the east coast of Luzon, Philippines and the remnants of Typhoon Joe in the vicinity of northern Laos, 25 July 1980, 0246Z. (NOAA 6 imagery)



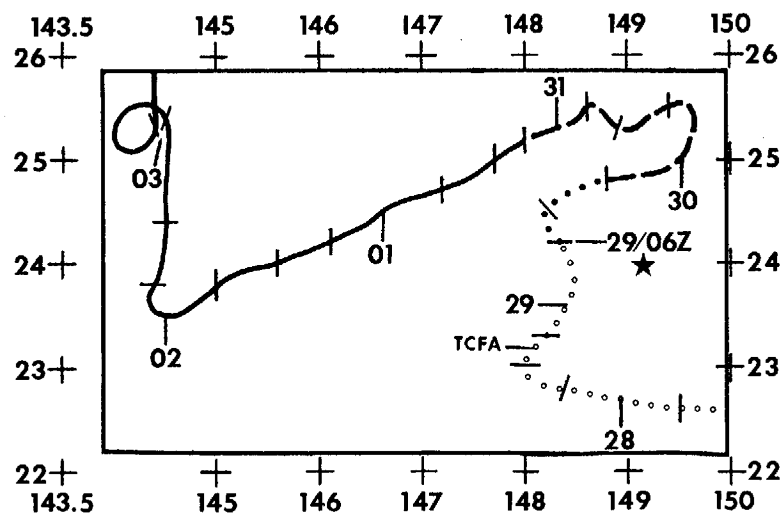
TYPHOON LEX

BEST TRACK TC-12
29 JUL-07 AUG 1980
MAX SFC WIND 80 KTS
MINIMUM SLP 962 MBS

LEGEND

- 06 HOUR BEST TRACK POSIT
- A SPEED OF MOVEMENT
- B INTENSITY
- C POSITION AT XX/0000Z
- ∘∘∘ TROPICAL DISTURBANCE
- ••• TROPICAL DEPRESSION
- — — TROPICAL STORM
- —— TYPHOON
- ◆ SUPER TYPHOON START
- ◇ SUPER TYPHOON END
- ◇◇◇ EXTRATROPICAL
- ••• DISSIPATING STAGE
- ★ FIRST WARNING ISSUED
- ☆ LAST WARNING ISSUED

DTG	SPEED	INTENSITY
29/06Z	6	30
29/12Z	6	30
29/18Z	7	35
30/00Z	8	40
30/06Z	6	45
30/12Z	4	50
30/18Z	4	55
31/00Z	3	60
31/06Z	4	65
31/12Z	5	65
31/18Z	6	65
01/00Z	6	65
01/06Z	6	65
01/12Z	6	70
01/18Z	6	70
02/00Z	7	75
02/06Z	7	75
02/12Z	9	75
02/18Z	7	75
03/00Z	9	80



Typhoon Lex was the most difficult tropical cyclone to forecast during the entire 1980 season. This typhoon developed from a Tropical Upper Tropospheric Trough (TUTT) near 22N 152E and initially moved westward. From this point, Lex made five right angle or greater turns and executed one tight cyclonic loop before finally heading northeastward into the western Pacific east of Japan. The only saving grace was that Lex remained well away from major landmasses and did not affect any military installations ashore or afloat.

Lex was first observed as a small disturbed area of convection on 24 July. The first satellite position fix at 260600Z placed the disturbance approximately 125 nm (230 km) south-southwest of Marcus Island. The disturbance moved almost due west (Fig. 3-12-1), and a Tropical Cyclone Formation Alert (TCFA) was issued at 281500Z when the satellite signature improved. The first warning was issued for Tropical Depression (TD) 12 at 290600Z after aircraft reconnaissance located a surface circulation center with a central pressure of 1002 mb and estimated maximum surface winds of 35 kt (18 m/sec). Twelve hours later, as the satellite signature continued to improve, the cyclone was upgraded to Tropical Storm Lex.

During the early development stage, a deep steering current was not evident above Lex. However, a broad 200 mb trough to the north-northeast seemed to have the strongest influence and turned Lex from a westward to a northeastward track. As the upper trough moved eastward, a middle-and upper-level ridge built northwest of Lex. The steering currents veered from southwesterly to northeasterly in response to the intensifying subtropical ridge, and Lex turned to a southwestward track.

Lex continued to intensify slowly during his southwestward movement, reaching typhoon strength of 65 kt (33 m/sec) at 310600Z. Shortly after 020000Z, Lex again changed direction and headed on a northward track through a break in the subtropical ridge. The break developed as a trough deepened to the north over the Sea of Japan. At the same time, anticyclonic cells intensified at all levels to the southeast and west-southwest of Lex.

Lex executed a cyclonic loop while accelerating northward and, before completing the loop, reached his maximum intensity of 80 kt (41 m/sec). The satellite signature for Lex at maximum intensity is illustrated in Figure 3-12-2. Upon exiting the loop, Lex continued tracking north until a deep surface low and associated cold front began moving eastward across Japan. As the frontal system approached from the west, Lex commenced recurvature to the northeast and accelerated slightly. The slow entrainment of cold air caused Lex to weaken and transition into an extratropical system. The last warning was issued for Lex at 070000Z August 1980.

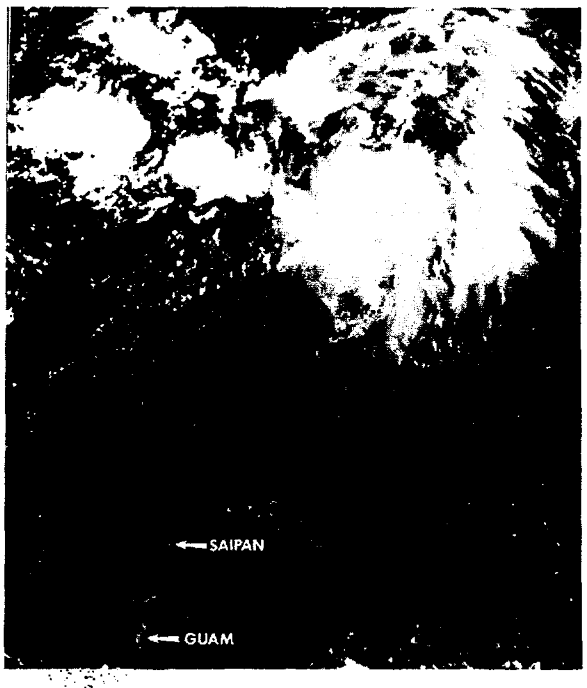
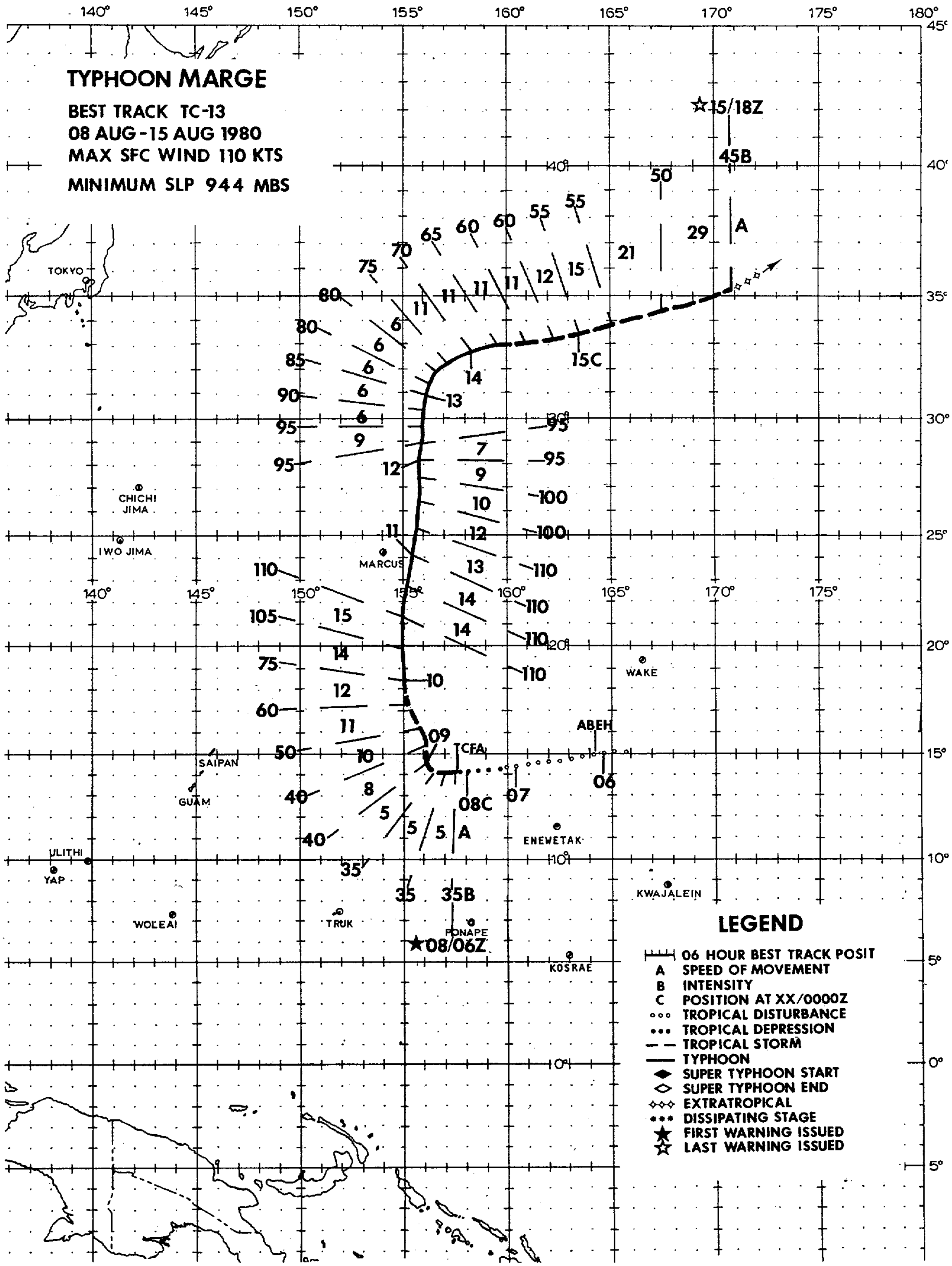


FIGURE 3-12-1. Lex as a tropical disturbance prior to issuance of a TCFA, 27 July 1980, 0026Z. (DMSP imagery)

FIGURE 3-12-2. Typhoon Lex at maximum intensity of 80 kt (41 m/sec), 2 August 1980, 2242Z. Iwo Jima is about 155 nm (285 km) west of Lex. (NOAA6 imagery)


TYPHOON MARGE
BEST TRACK TC-13
08 AUG-15 AUG 1980
MAX SFC WIND 110 KTS
MINIMUM SLP 944 MBS
TOKYO
CHICHI JIMA
IWO JIMA
MARCUS
SAIPAN
GUAM
ULITHI
YAP
WOLEAI
TRUK
PONAPE
ENEWETAK
KWAJALEIN
KOSRAE
WAKE
ABEH
TCFA
08C
15C
35B
45B
15/18Z
08/06Z
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED

Marge was the sixth tropical cyclone to reach typhoon strength during 1980. She developed west of the Marshall Islands in an area that had shown considerable instability since Typhoon Lex passed through the area in late July.

The convection that signaled Marge's formation first appeared on satellite imagery at 051200Z August 1980. Because of continual intensity variations in the convection, the tropical disturbance was not considered suspect until 060600Z when it was first mentioned in the Significant Tropical Weather Advisory Bulletin (ABEH PGTW). Through most of this period, the convection was embedded in a broad easterly flow. Typhoon Lex still displayed considerable influence over the region, causing the usual easterly current to be diverted northward over the Mariana Islands (Fig. 3-13-1). By 060000Z, Lex had moved far enough to the north that his influence over the easterly flow had weakened and the surface flow had split. One current was still drawn northward toward Lex, while the other current curved southward between the Marshall Islands and the Northern Mariana Islands. The southern current was drawn back into a broad low-level circulation between the eastern Caroline Islands and the Marshall Islands (Fig. 3-13-2). Satellite imagery showed an increase in convection corresponding to this change in the flow pattern.

Convective activity appeared to consolidate near 15N 159E by 061600Z. The convective area continued to expand and by 070000Z covered an area nearly 5-degrees square, with the most intense activity remaining near 15N 159E. Post-analysis shows that Marge formed during the period between 070000Z and 071200Z. An evaluation of the satellite imagery for this time period indicates that tropical depression stage was attained at 070600Z. By 071200Z, a north-south trough oriented along 160E was analyzed at the surface/gradient level. The circulation associated with Marge appeared to be part of this trough (Fig. 3-13-3).

The first reconnaissance into Marge, at 080533Z, observed surface winds of 35 kt (18 m/sec) and a central pressure of 998 mb. Based on these data, the initial warning on

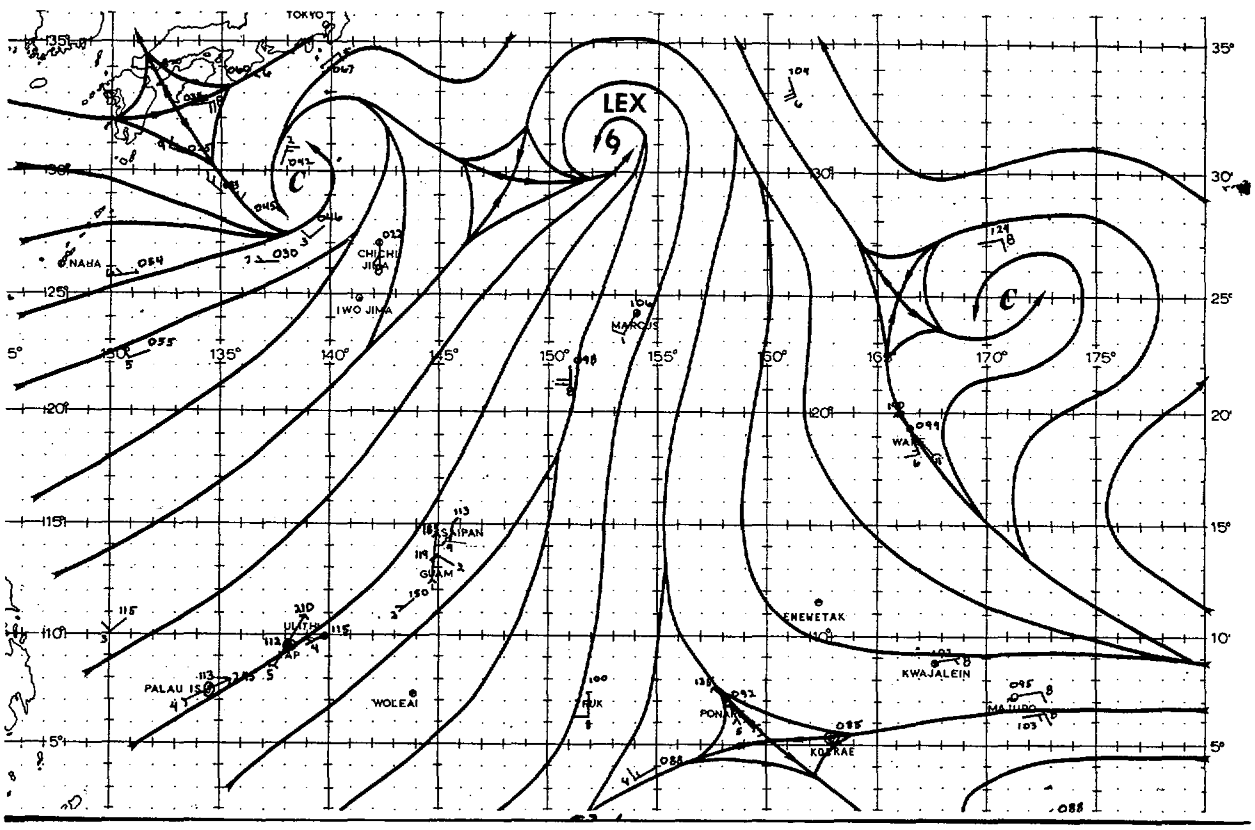


FIGURE 3-13-1. The 051200Z August 1980 surface (⟵⟍)/gradient-level (ddd⟵⟨ff) wind data and streamline analysis. Wind speeds are in knots.

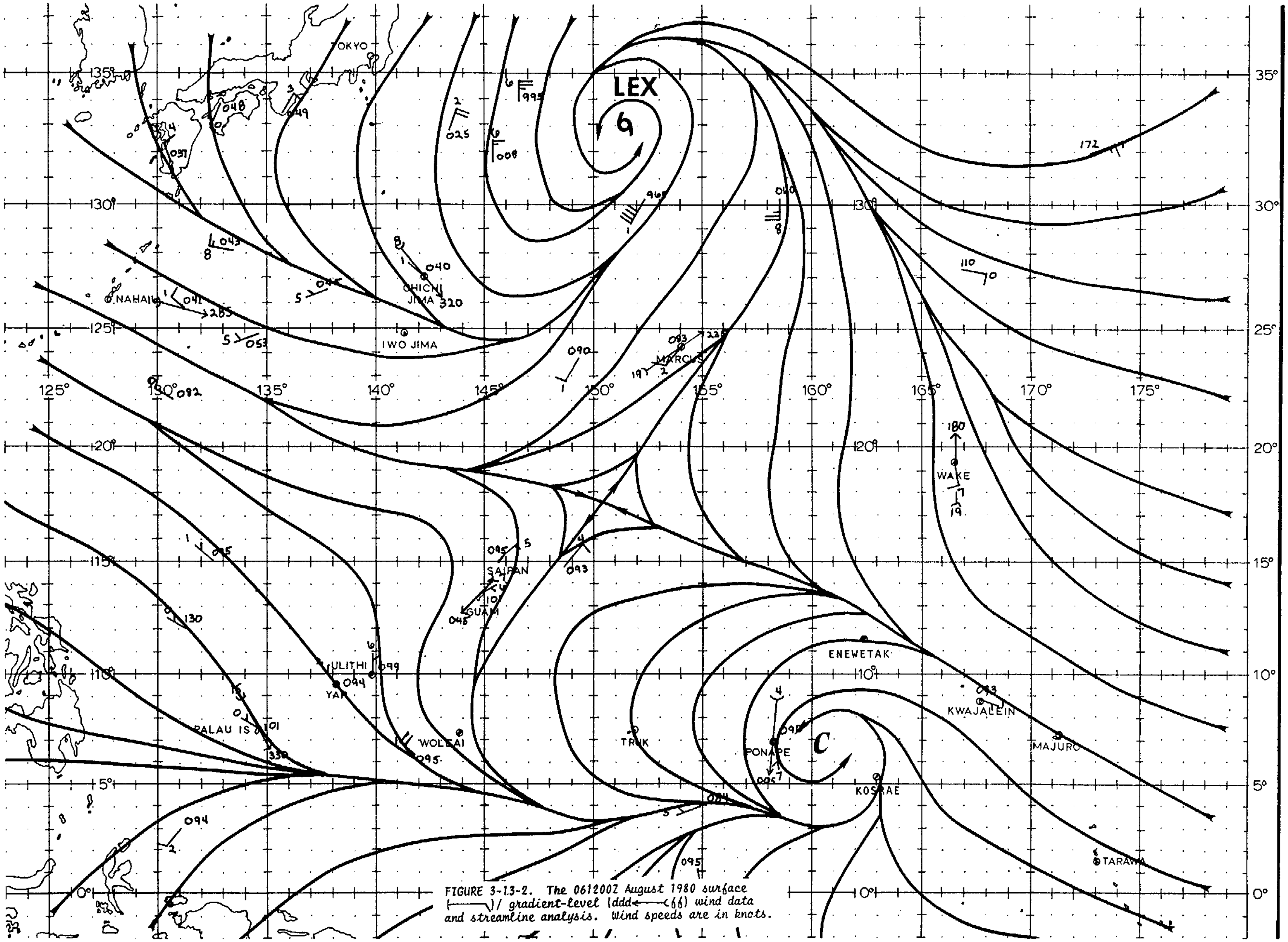


FIGURE 3-1.3-2. The 061200Z August 1980 surface (wind barb) / gradient-level (ddd←ff) wind data and streamline analysis. Wind speeds are in knots.

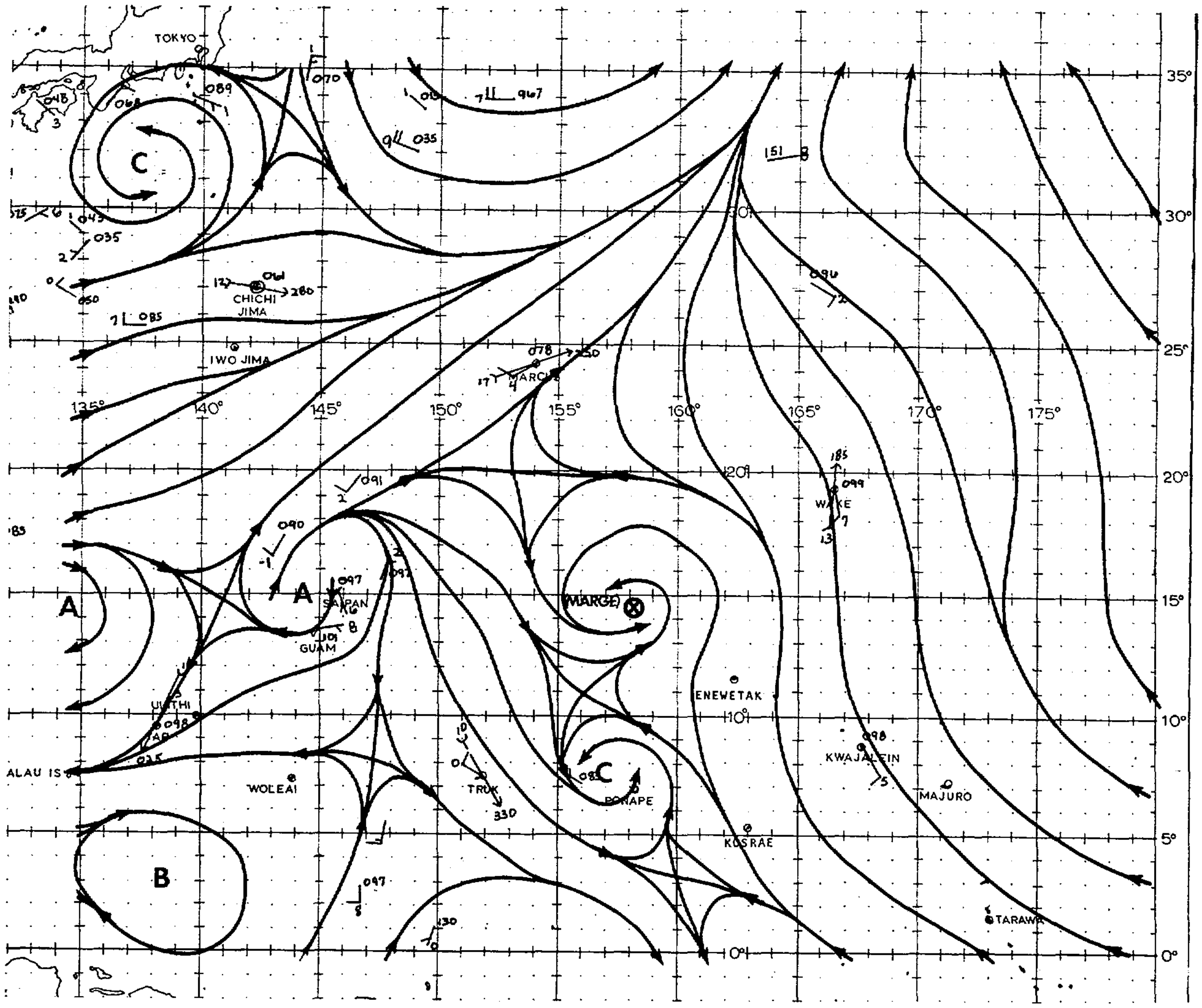


FIGURE 3-13-3. The 071200Z August 1980 surface (⟵) /gradient-level (ddd⟵(ff) wind data and streamline analysis. Wind speeds are in knots.

Tropical Storm Marge was issued at 080800Z.

Marge initially followed a generally west-southwestward track. Objective forecast aids showed considerable scatter at this stage, a common occurrence during the formative stages of a tropical cyclone. A mid-tropospheric ridge was analyzed to the north of Marge. The key questions at that time concerned the status of this ridge, i.e., was it strong enough to keep Marge on a west-southwestward track, or was there a weakness which could allow her to recurve to the north-northwest. The 080000Z 500 mb streamline analysis (Fig. 3-13-4) indicated that Marge was located in a col, thus providing a channel for a more northerly track than predicted by climatology. The 500 mb reconnaissance data provided by the 54th Weather Reconnaissance Squadron north of Marge proved very valuable in locating this col. A sequence of satellite fixes between 081600Z and 082330Z was the first indication that Marge was reacting to the weakness in the ridge and had started a northward turn. A 090300Z satellite position fix, combined with aircraft fixes at 090615Z and 090839Z, confirmed the northward track.

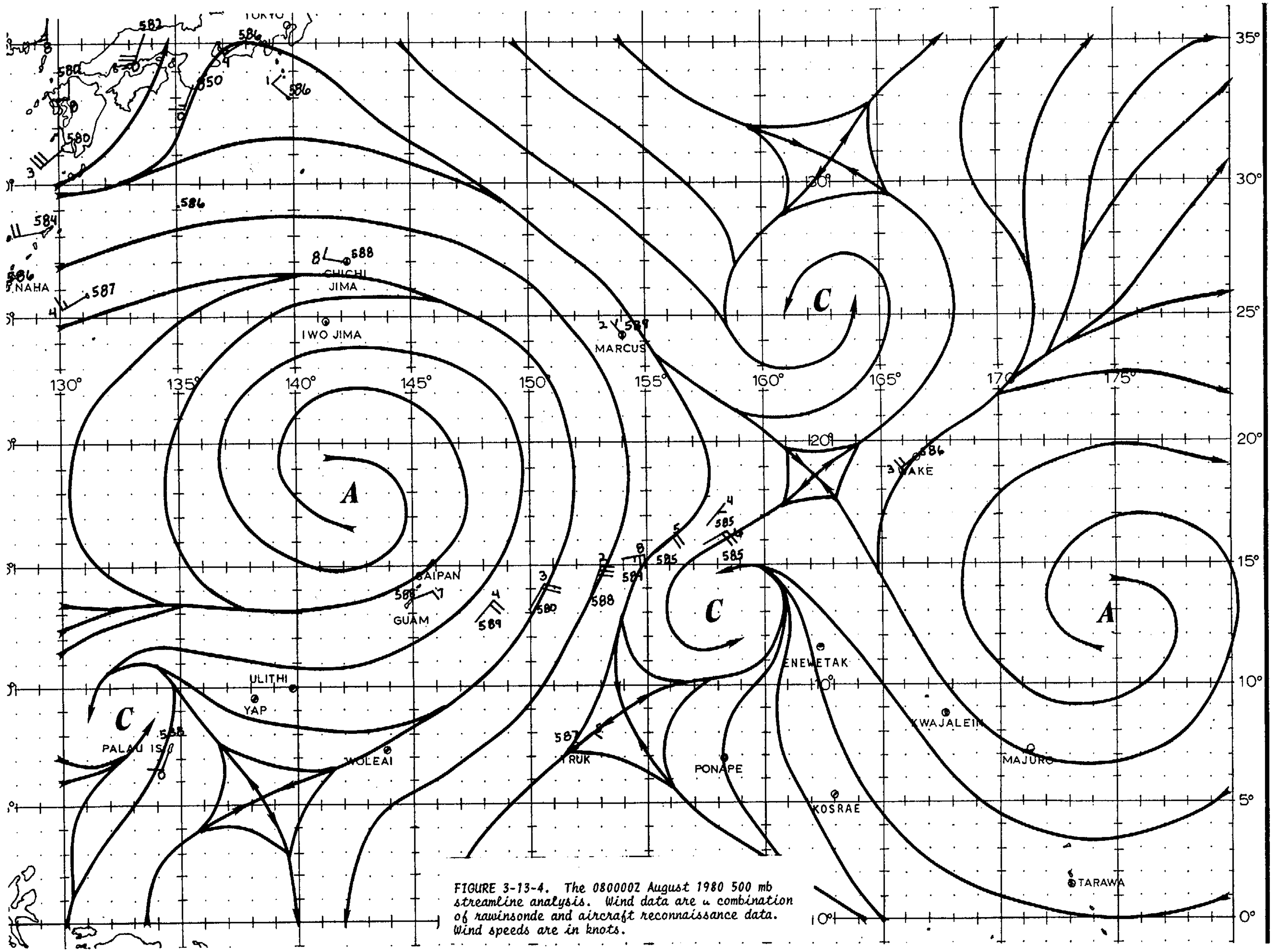


FIGURE 3-13-4. The 080000Z August 1980 500 mb streamline analysis. Wind data are a combination of rawinsonde and aircraft reconnaissance data. Wind speeds are in knots.

Marge continued northward for 17 degrees of latitude on a track between two centers in the subtropical ridge. During the northward trek, Marge intensified to typhoon strength which she maintained for nearly 5 days. A minimum sea-level pressure of 944 mb supported a maximum intensity of 110 kt (56 m/sec) (Atkinson and Holliday, 1977) for 18 hours (Fig. 3-13-5).

By 131200Z, Marge began to encounter strong upper-level westerlies. A second course change accompanied by gradual acceleration and weakening began at that time. Marge tracked east-northeastward and continued to accelerate under the influence of the strong mid-latitude westerlies. The final warning on Marge was issued at 151800Z as she transitioned into an extratropical cyclone and merged with a mid-latitude low pressure system.

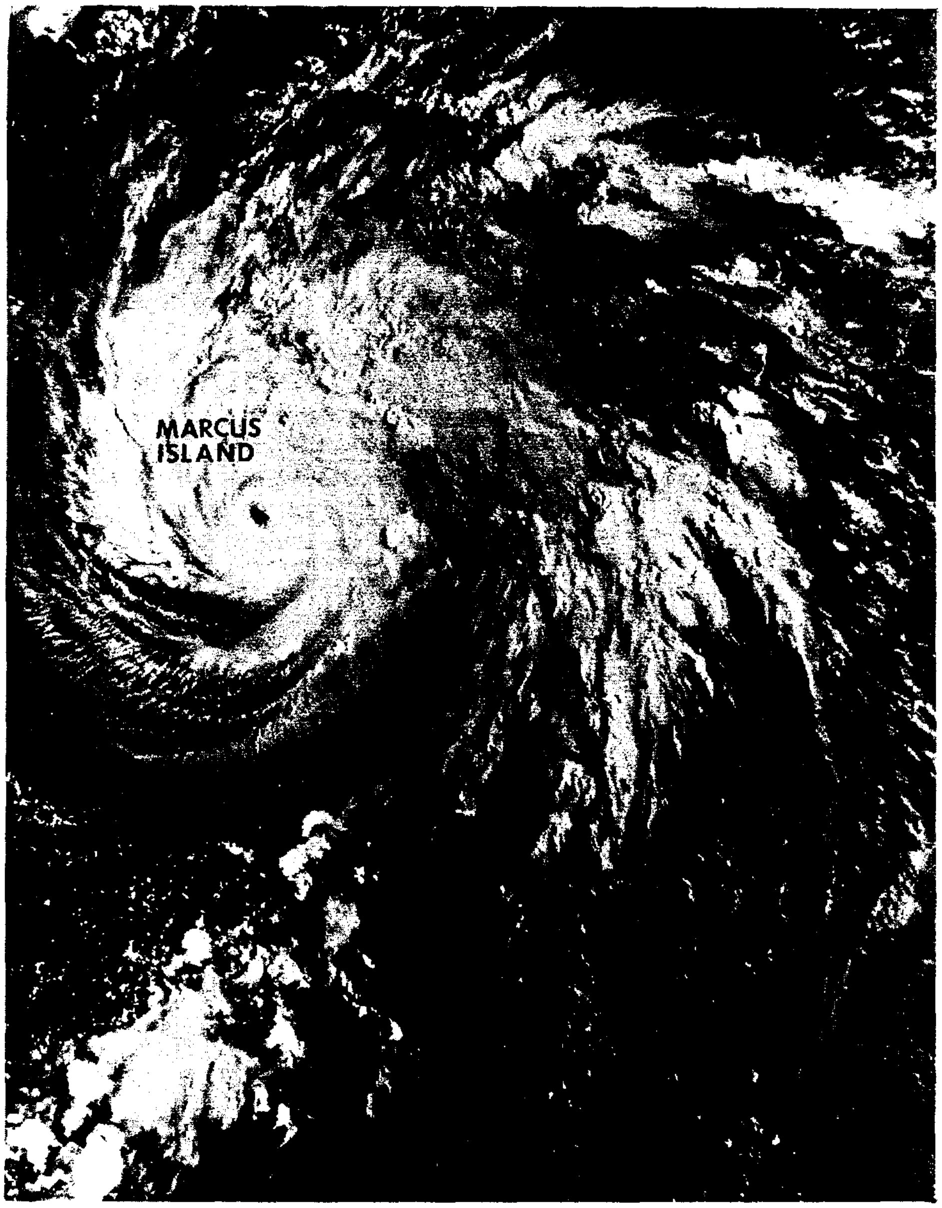


FIGURE 3-13-5. *Visual satellite imagery of Typhoon Marge just after reaching maximum intensity and minimum sea-level pressure, 10 August 1980, 2124Z. (NOAA6 imagery)*

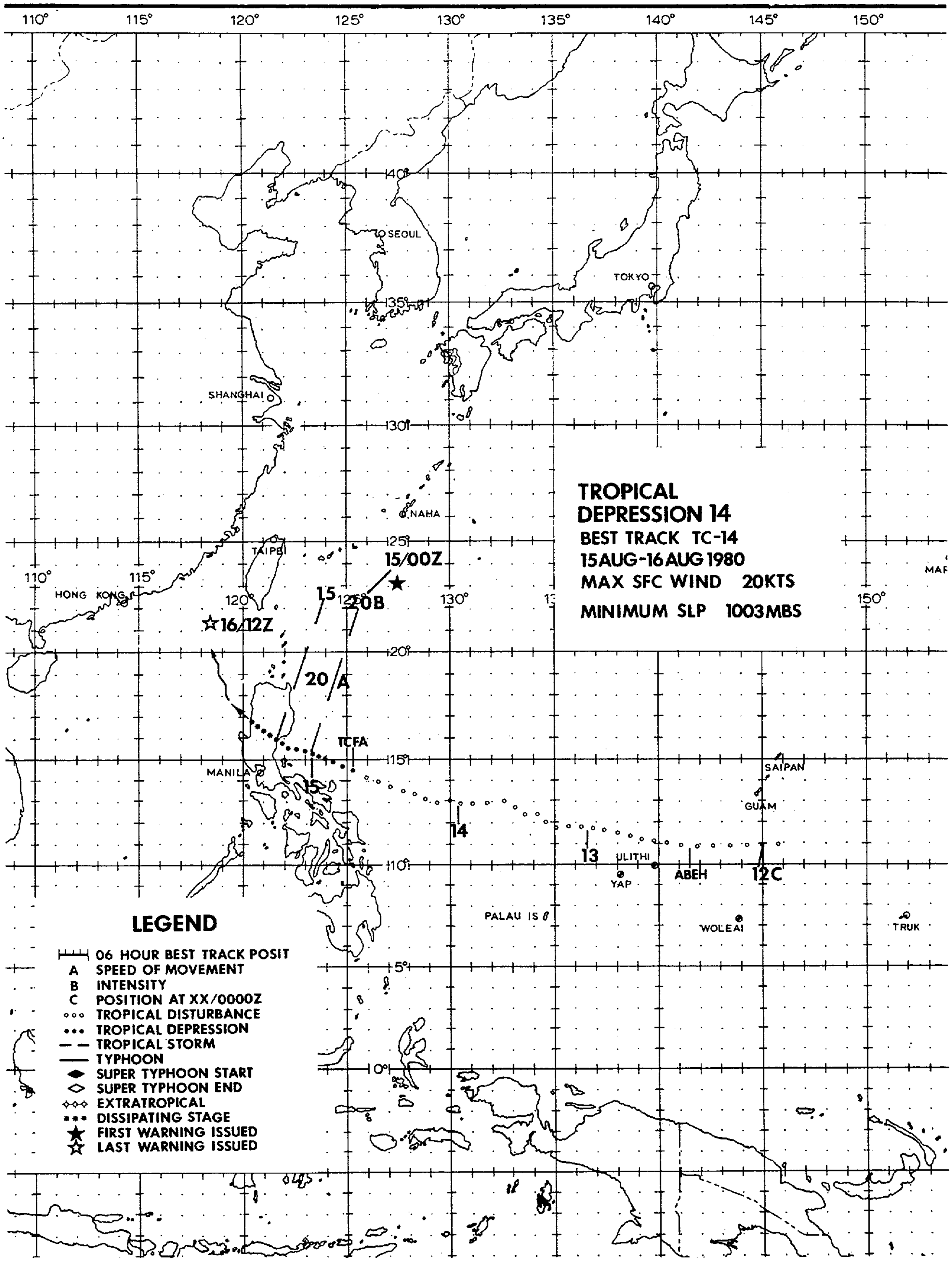
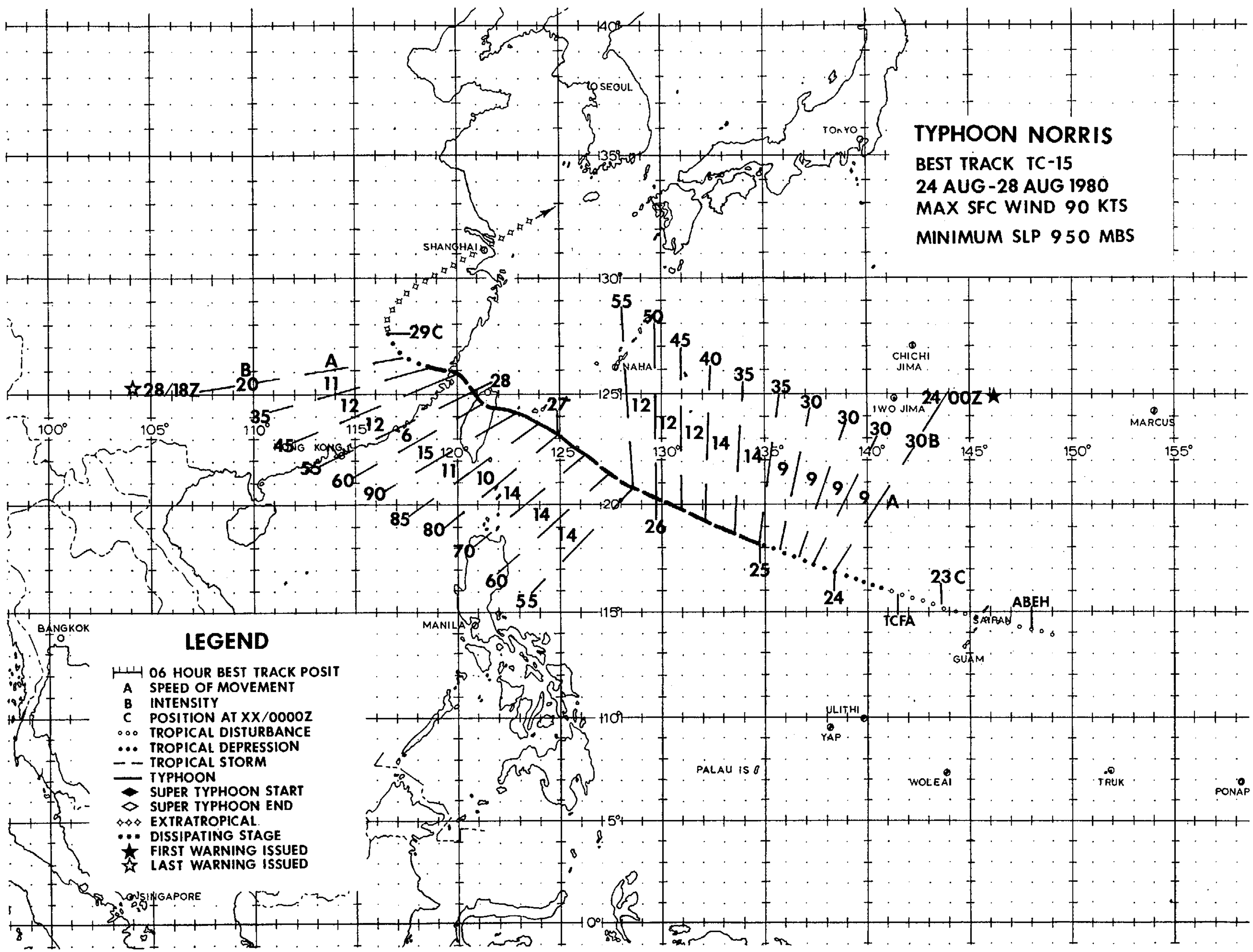

TROPICAL DEPRESSION 14
BEST TRACK TC-14
15AUG-16AUG 1980
MAX SFC WIND 20KTS
MINIMUM SLP 1003MBS
15/00Z
☆16/12Z
20B
20/A
TCFA
15
14
13
12C
SEOUL
TOKYO
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
SAIPAN
GUAM
ULITHI
YAP
ABEH
PALAU IS
WOLEAI
TRUK
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED

FIGURE 3-14-01. Tropical Depression 14 at 20 kt (10 m/sec) intensity just prior to landfall on the east coast of Luzon, 14 August 1980, 2317Z (NOAA6 imagery). TD 14 weakened rapidly over the mountains of Luzon and never reintensified over the South China Sea as was expected.


TYPHOON NORRIS
BEST TRACK TC-15
24 AUG-28 AUG 1980
MAX SFC WIND 90 KTS
MINIMUM SLP 950 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SEOUL
TOKYO
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
HONG KONG
MANILA
BANGKOK
SINGAPORE
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
SAIPAN
GUAM
ABEH
TCFA
24/00Z
23C
24
25
26
27
28
28/18Z
29C
100°
105°
110°
115°
120°
125°
130°
135°
140°
145°
150°
155°
40°
35°
30°
25°
20°
15°
10°
5°
0°

The near equatorial trough was reestablished between Guam and Ponape as Typhoon Marge moved northward toward Marcus Island on 10 August. A weak surface circulation developed along the trough axis south of Guam and slowly drifted toward the Philippine Islands over the next two weeks. Although this disturbance never developed into a significant tropical cyclone, it played a major role in delaying the intensification of a disturbance that tracked westward from Wake Island and eventually became Typhoon Norris.

A deep Tropical Upper-Tropospheric Trough (TUTT) was first analyzed on the 151200Z 200 mb analysis over the Marshall Islands from Wake Island southwestward to Truk. Sparse surface data gave no indications of a perturbation in the low-level tradewind flow at that time.

For the next seven days, the TUTT and associated convective activity to the southeast migrated slowly westward. Figure 3-15-1 depicts the position of the TUTT in relation to the area of enhanced convection that eventually developed into Typhoon Norris. It was not until 211200Z that the upper-level disturbance was reflected at the surface as a weak circulation.

The increased convection and resultant heating of the upper troposphere, as the disturbance approached a position north of Guam, is graphically illustrated in Figure 3-15-2. The streamline analysis reveals the development of a sharp ridge which built northeastward toward Marcus Island and eventually split the TUTT into two cells. By 230000Z, an upper-level anticyclone had formed over the surface disturbance, and, as the disturbance continued to organize, a Tropical Cyclone Formation Alert was issued at 230900Z. The preceding discussion illustrates the initiation of a tropical cyclone induced by upper-level divergence and enhanced convection southeast of the TUTT cell (Sadler, 1978). Norris tracked virtually straight west-northwestward at an average speed of 12 kt (22 km/hr) from the time of first warning as a tropical depression at 240200Z until landfall on northern Taiwan at 271600Z. This straight track was due to the strong mid-level subtropical ridge which extended along 27N from southern China eastward to the International Dateline during the latter part of August.

The circulation mentioned earlier near the Philippine Islands prevented Norris from developing and intensifying more rapidly. The surface flow pattern was split between the two circulations until 260000Z when the other circulation finally went ashore over Luzon and dissipated. With all the low-level inflow now available, Norris intensi-

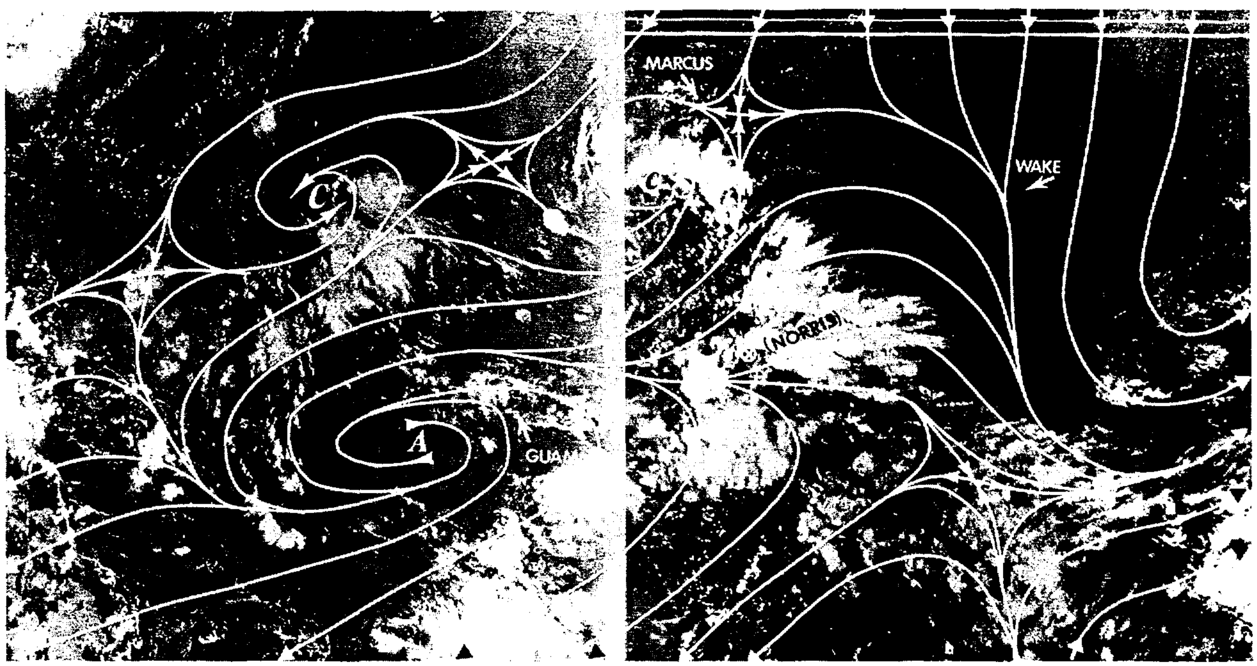


FIGURE 3-15-1. 210000Z August 1980 200 mb streamline analysis superimposed on satellite imagery at 202200Z. This figure depicts convective activity associated with upper-level cyclonic circulations and the enhanced convection southeast of the TUTT that eventually developed into Typhoon Norris. (NOAA6 imagery)

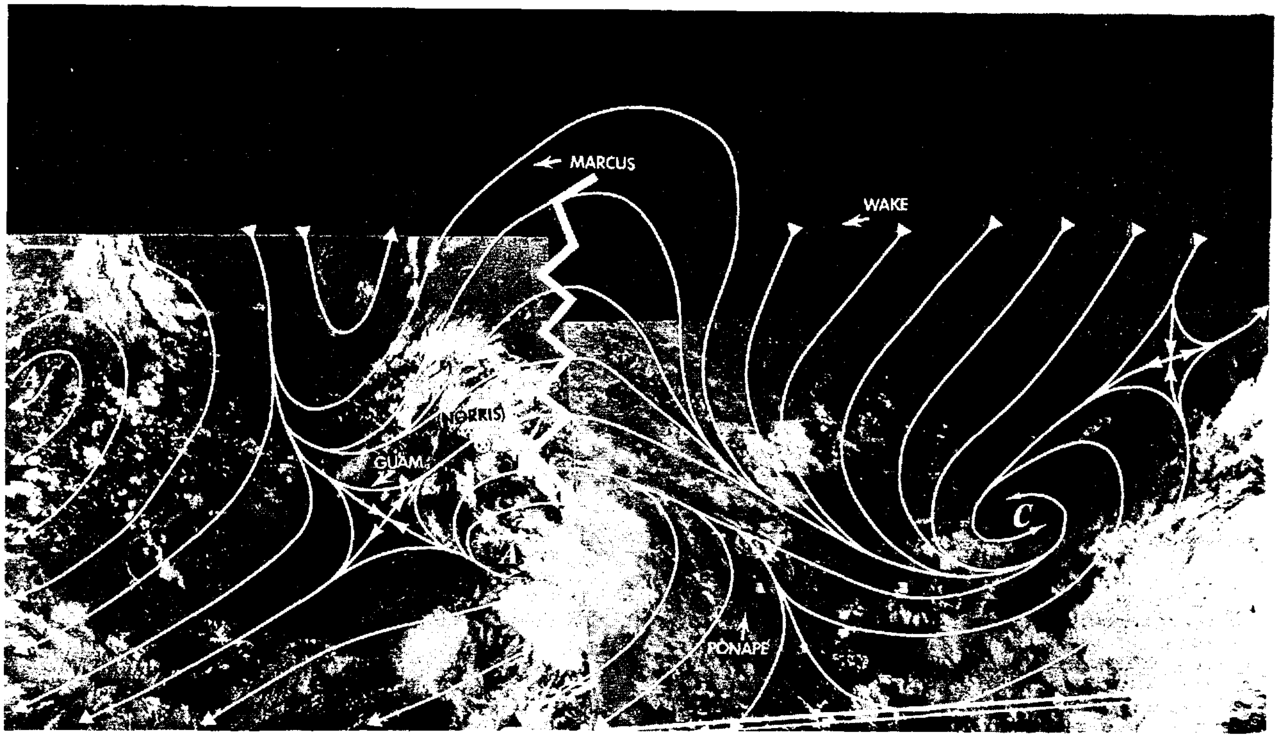


FIGURE 3-15-2. 220000Z August 1980 200 mb streamline analysis superimposed on satellite imagery at 212100Z. This figure illustrates convective heating which built a sharp ridge to the east of the TUTT cell. This ridge pushed the TUTT northward and eventually split it into two cells. (NOAA6 imagery)

fied quickly from 50 kt (26 m/sec) and 985 mb at 260000Z to a peak of 90 kt (46 m/sec) and 950 mb about 36 hours later.

Norris' equivalent potential temperature (θ_e) and minimum sea-level pressure (MLSP) curves intersected at 260000Z. Using JTWC's θ_e/MSLP study (see discussion on Super Typhoon Kim), Norris' sea level pressure was expected to fall 44 mb and maximum surface winds to increase 55 kt (28 m/sec) beyond that point. It seems very likely that this intensification would have occurred if landfall on Taiwan had not taken place within 42 hours.

The well-established mid-level ridge north of Norris, with a strong high pressure cell located between Taiwan and Okinawa, was responsible for the climatological west-northwestward track with Norris skirting the northern tip of Taiwan. However, with Norris 500 nm (926 km) southeast of Taiwan, the 260000Z 500 mb analysis indicated that the ridge was beginning to weaken between Taiwan and Okinawa with the high pressure cell retreating northeastward to a position east of Okinawa. By 261200Z, a definite break in the ridge was evident with the high cell now over the Bonin Islands and a secondary cell located near 25N 112E over southern China. The numerical forecast series during this period also supported the persistence of this break. Thus, JTWC's warnings after 260000Z were consistent in forecasting recurvature

FIGURE 3-15-3. Typhoon Norris 20 nm (37 km) southeast of Yonagunijima (WMO station 47912) and 95 nm (176 km) southeast of Taipei near peak intensity of 90 kt (46 m/sec), 27 August 1980, 1051Z (NOAA6 imagery). Yonagunijima reported sustained winds of 80 kt (41 m/sec) one hour later.

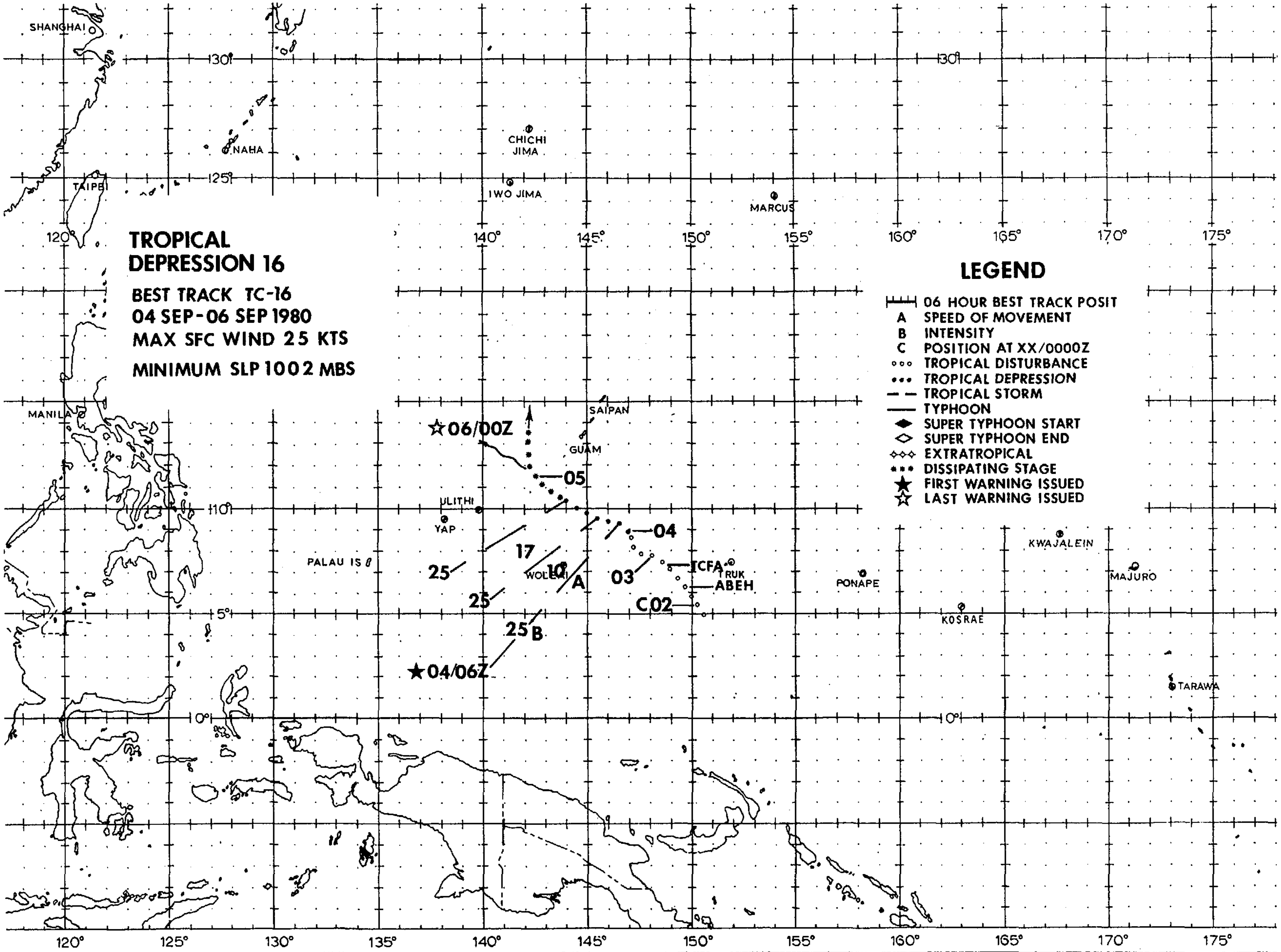
north of Taiwan and along the coast of mainland China into the Korea Strait.

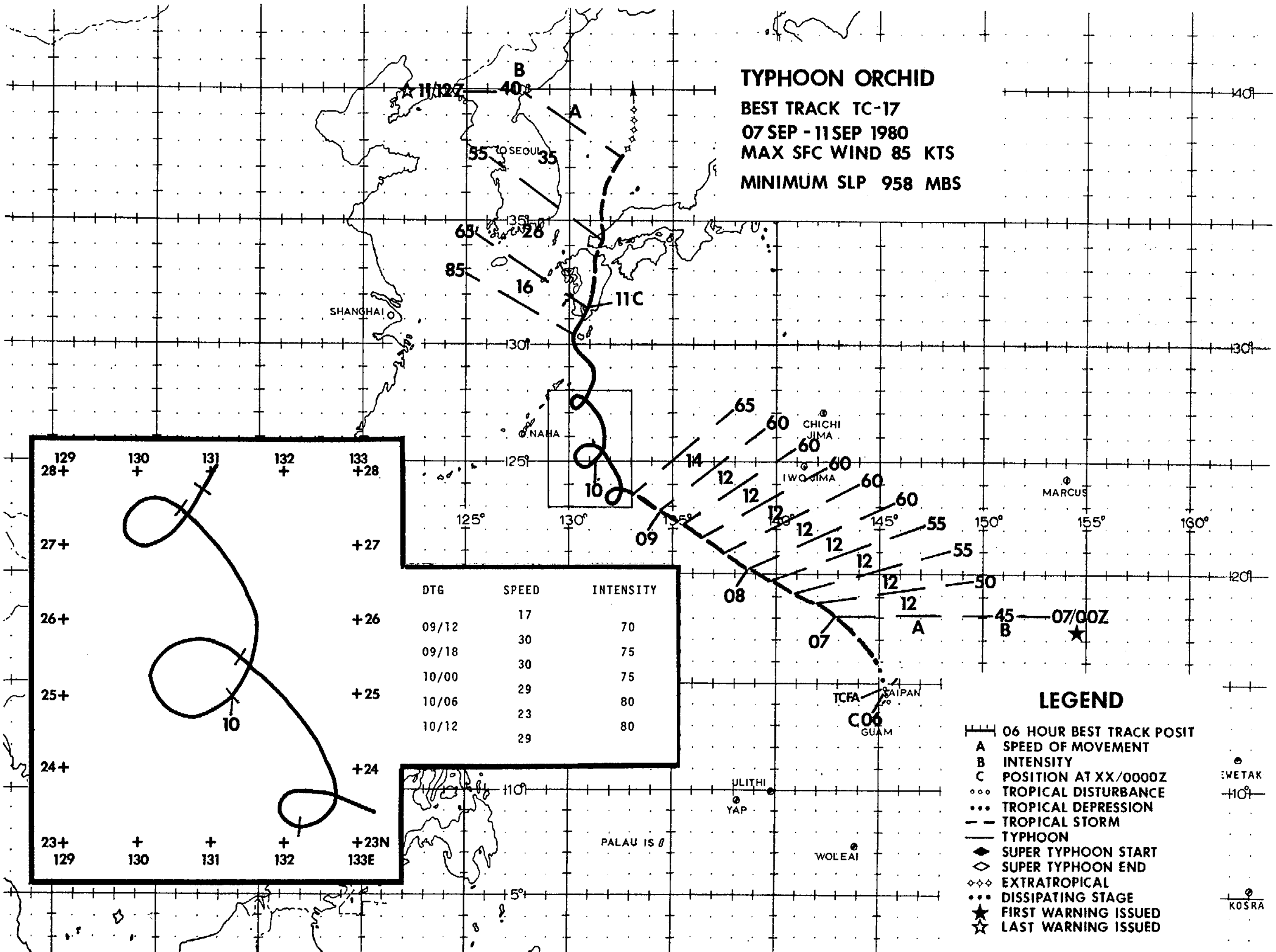
Norris passed 10 nm (18 km) southwest of Yonagunijima at 271200Z. At that time, the island reported southeast winds of 80 kt (41 m/sec) and a sea-level pressure of 952.2 mb (Fig. 3-15-3). Norris then turned to a more westward track toward northern Taiwan. Excellent radar coverage from the island stations of Ishigakijima and Miyakojima and from Hua-Lien on Taiwan permitted JTWC to follow Norris as he tracked across Taiwan and into the Formosa Strait just north of Hsin-Chu. Strongest surface winds of 39 kt (20 m/sec) with gusts to 64 kt (33 m/sec) on northern Taiwan were reported by Taipei at 271600Z (Fig. 3-15-4).

Norris' track across Taiwan, change in speed, and observed weakening were classic examples of the effects of Taiwan on tropical cyclones (Brand and Blelloch, 1973). The mountainous terrain of Taiwan apparently produced an induced surface low on the lee side of the mountain range, resulting in the marked increase in speed and the westward bend in Norris' track.

Landfall just south of Fu-Chou on mainland China occurred about 280900Z and, although penetrating deeper inland than forecast, Norris eventually recurved northeastward and the remnants linked with a frontal system that moved out over the Yellow Sea and Sea of Japan.

FIGURE 3-15-4. Typhoon Norris as seen by radar at Hua-Lien, 27 August 1980, 0800Z. (Photograph courtesy of the Central Weather Bureau, Taipei, Taiwan)


TROPICAL DEPRESSION 16
BEST TRACK TC-16
04 SEP-06 SEP 1980
MAX SFC WIND 25 KTS
MINIMUM SLP 1002 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
06/00Z
05
04
03
C02
TCFA
ABEH
25
17
10
A
25
25
25 B
04/06Z
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
TAIPEI
MANILA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
KOSRAE
KWAJALEIN
MAJURO
TARAWA
120° 125° 130° 135° 140° 145° 150° 155° 160° 165° 170° 175°
30° 25° 10° 5° 10°



TYPHOON ORCHID
BEST TRACK TC-17
07 SEP - 11 SEP 1980
MAX SFC WIND 85 KTS
MINIMUM SLP 958 MBS
DTG SPEED INTENSITY
17
09/12 70
30
09/18 75
30
10/00 75
29
10/06 80
23
10/12 80
29
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
11/12Z
07/00Z
TCFA
SEOUL
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
KOSRA

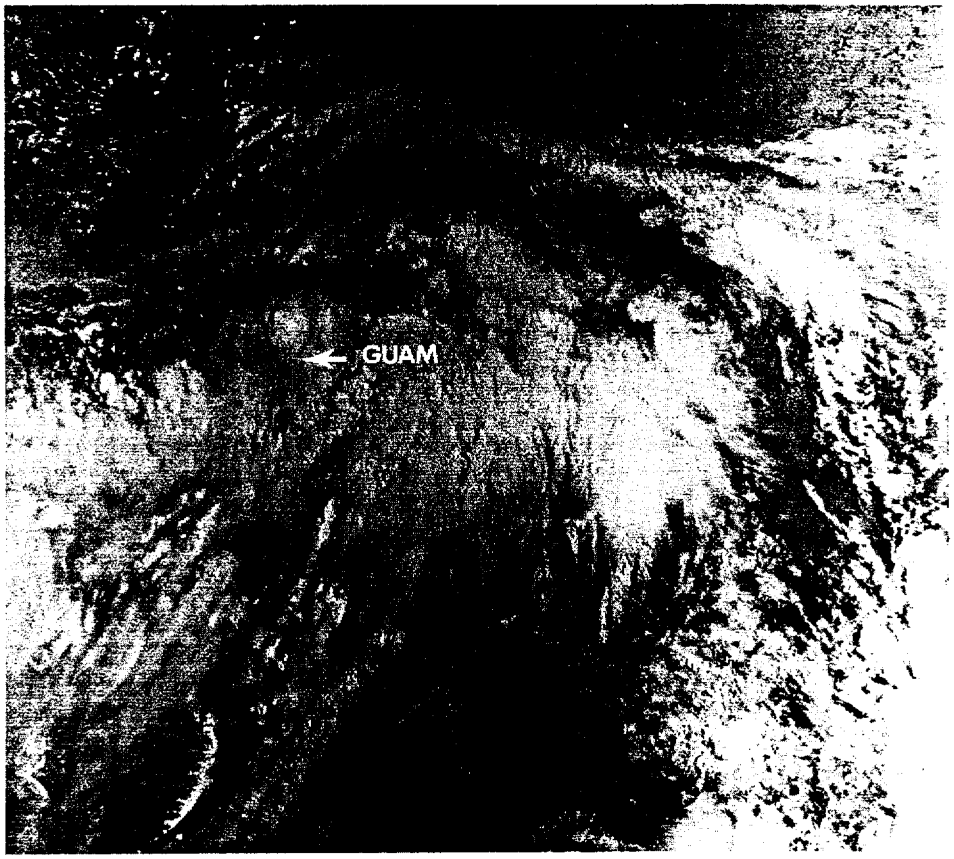


FIGURE 3-17-1. Satellite imagery showing extensive unorganized convection in the vicinity of Guam, 05 September 1980, 2150Z (NOAA6 imagery). This imagery along with synoptic and aircraft data supported the decision to issue the final warning on TD 16. The increased convection east and north of Guam also alerted JTWC to the possible development of another disturbance between Guam and Saipan.

The discussion for Typhoon Orchid would be incomplete without reviewing the brief life of Tropical Depression 16. Both systems developed near the eastern extension of the monsoon trough. The dissipation of Tropical Depression 16 was followed by the subsequent development of the disturbance which became Typhoon Orchid. The influence of the monsoon trough was investigated to explain the structure of these cyclones and, ultimately, to offer an explanation for Orchid's most unusual behavior south of Japan.

During the first few days of September, the monsoon trough was evident as far east as 160E along 05N. Satellite imagery on 2 September indicated increased convection near a weak circulation at the eastern end of the trough. The first of two formation alerts was issued at 021400Z. Further development was not observed on satellite imagery during the next 36 hours. A reconnaissance aircraft at 040155Z located a closed surface circulation with 25 kt (13 m/sec) maximum winds and a minimum sea-level pressure of 1002 mb. The first warning on TD 16 followed at 040600Z, and during the next 42 hours, JTWC tracked the depression as it moved west-northwestward. Aircraft investigations during this period showed a largely unorganized system. Unlike the investigation at 040155Z, these investigations repeatedly suggested that multiple centers existed in the area. Post-analysis indicated that sometime during the 42-hour period, the surface center associated with TD 16 weakened within the trough while JTWC continued to follow a persistent convective center to the west.

Although TD 16 continued to weaken, warnings were still issued because the potential for significant tropical cyclone development remained high in the region. Another disturbance eventually developed northeast of TD 16 as TD 16 weakened. Satellite imagery received at 052150Z (Fig. 3-17-1) showed that the entire area near Guam was under extensive, but apparently unorganized convection. The final warning was issued for TD 16 when aircraft reconnaissance at 060050Z failed to locate a significant surface circulation.

By 060000Z, satellite imagery indicated that a tropical cyclone formation alert was required for a rapidly developing disturbance just north of Guam. A reconnaissance aircraft investigated the disturbance at 060120Z but was unable to close a surface circulation. The aircraft and synoptic data showed an extensive light and variable wind area extending more than 100 nm (185 km) west of the disturbance. Synoptic data, nevertheless, indicated that gale force winds (greater than 33 kt (17 m/sec)) existed in the eastern semicircle of the disturbance. After coordination with forecasters at Naval Oceanography Command Center, Guam[1], a gale warning was issued for the area. The first warning for Tropical Storm Orchid was issued at 070200Z. This warning was based on aircraft reconnaissance at 070005Z which observed 45 kt (23 m/sec) surface winds in the northeast quadrant of the storm. The same aircraft observed only 10-15 kt (5-8 m/sec) northwest winds in the western quadrant, indicating that a closed surface circulation existed only for 6-12 hours before the first warning on Tropical Storm Orchid.

During the five-day period from 02 to 08 September, the axis of the monsoon trough moved from 05N to 18N. A near equatorial or buffer ridge developed at low latitudes and extended from the Philippines to the east of Orchid. The pre-existing subtropical ridge north of Orchid and the presence of the near equatorial ridge provided a broad wind band, which extended counter-clockwise from the south-southwest of Orchid to the northwest at distances as far as 800 nm (1482 km) from Orchid's center. A composite surface streamline analysis from 07000Z to 091200Z indicates that this pattern maintained itself, virtually unchanged, during a 60-hour period during which Orchid moved west-northwest at 12 kt (22 km/hr). Figure 3-17-2 shows this pattern with the 081200Z surface wind field around Orchid superimposed. After 091200Z, the northwest wind component strengthened around Orchid as the monsoon trough began interacting with a mid-latitude trough in the east China Sea.

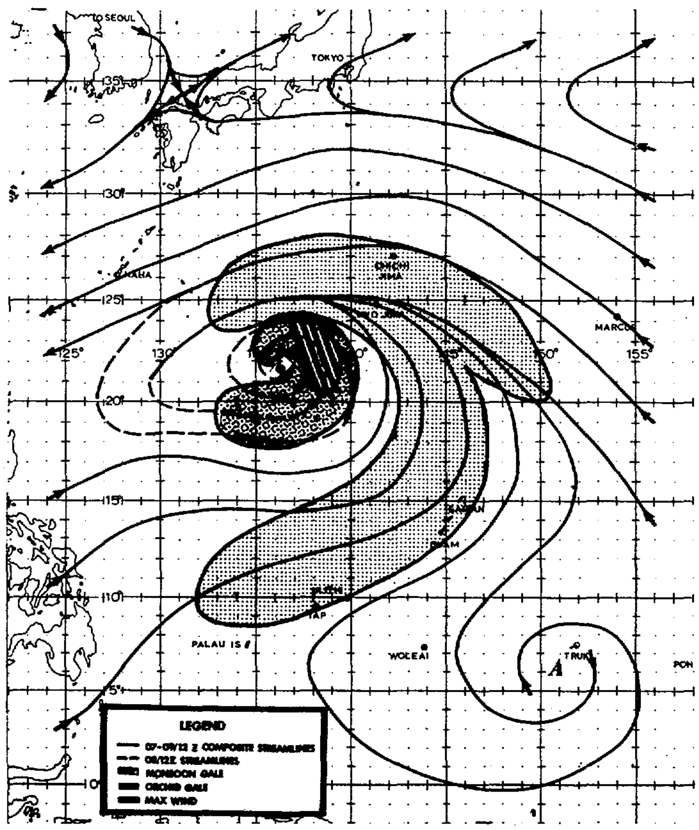


FIGURE 3-17-2. A composite surface streamline analysis of the monsoon trough based on data collected from 070000Z to 091200Z. The 081200Z streamlines (dashed lines) are superimposed for Orchid.

[1] The Joint Typhoon Warning Center functions operationally under the command of the Naval Oceanography Command Center, Guam (NOCC). Destructive wind warnings for the western North Pacific are included in the services provided by NOCC.

Although TD 16 and Typhoon Orchid developed in the same location with respect to the monsoon trough, it now appears that TD 16 failed to intensify because it could not sustain its own circulation pattern independent of the enhanced flow around the trough. Orchid sustained an independent circulation beginning on 7 September.

By 090000Z, Orchid had moved north into the subtropical latitudes near 23N. The monsoon trough was showing signs of weakening, and by 091800Z, the eastern portion of the wind band collapsed into Orchids's circulation pattern. Interaction with the mid-latitude trough moving east from Asia signalled the beginning of a change in Orchid's trajectory towards recurvature. The numerical prognostic series indicated a further eastward progression of the mid-latitude trough, but the series did not reflect the presence of Orchid at the middle and upper levels of the troposphere. Initial recurvature tracks anticipated a deepening of the trough and eventual recurvature southeast of Japan. The trough stalled near 130E, however, and the opportunity for recurvature was delayed until Orchid approached the Ryukyu Islands about 12 to 18 hours later.

In post-analysis, JTWC often finds some phenomenon that is not evident to the forecaster in real-time but which explains the motion or character of a tropical cyclone. In Orchid's case, JTWC was well aware of her circulation pattern; what wasn't known was the effect of this circulation pattern on Orchid's trajectory. Once formed, Orchid moved to the west-northwest at a nearly constant speed. During this portion of her track, Orchid was well behaved and there was no known "rule of thumb" which would have provided JTWC with a prior warning of the motion that the cyclone would undergo in the 36-hour period beginning at 090600Z. Beginning at 090600Z, Orchid executed three high speed cyclonic loops while maintaining an overall forward speed of 14 kt (26 km/hr) toward the north. Satellite, aircraft, and radar surveillance provided dense reconnaissance coverage of Orchid during these loops (Fig. 3-17-3). Orchid finally stabilized on her northward track just prior to landfall on Kyushu, Japan. Figure 3-17-4 illustrates an expanded surface best track, a partial 700 mb track based on aircraft data, and the overall smoothed track, which may have been followed by Orchid at some level above 700 mb. An analogy which may offer some insight into Orchid's unexplained motion is given next.

Before offering the analogy, some conjecture is required based on the assumption that Orchid's circulation pattern relative to the broad-scale circulation was "conditionally" unstable, i.e., all the forces acting on Orchid were only in balance as long as she maintained a constant heading. As Orchid approached the mid-latitude trough, this balance was interrupted and the potential unstable character of the cyclone, embedded in this particular synoptic pattern, was realized. One analogy that can be used to explain the trajectory involves a child's toy top. The top, inherently unstable because of its small base and wide body, will spin uniformly about its axis as long as it maintains equilibrium. A loss of rotational speed or a tap along the side will cause the top to stumble and the base will appear to accelerate along a predictable looping pattern until the top's stability is either restored or it comes to rest.

It is suggested here that the effect was virtually the same when Orchid began interacting with the mid-latitude trough. The best track shows that Orchid regained her equilibrium within the mid-latitude trough prior to making landfall in southern Japan. Orchid did not loop again and she returned to a slower speed of 18 kt (33 km/hr) prior to accelerating during the extratropical transition period.

Orchid caused considerable damage and loss of life in Japan and Korea. High winds and torrential rains associated with Orchid were blamed for six deaths, numerous injuries, and considerable damage to crops in southern Japan. At least three deaths were reported in South Korea as Orchid moved east of Korea into the Sea of Japan. Another 112 fishermen were reported missing in the Korea Straits following Orchid's passage.

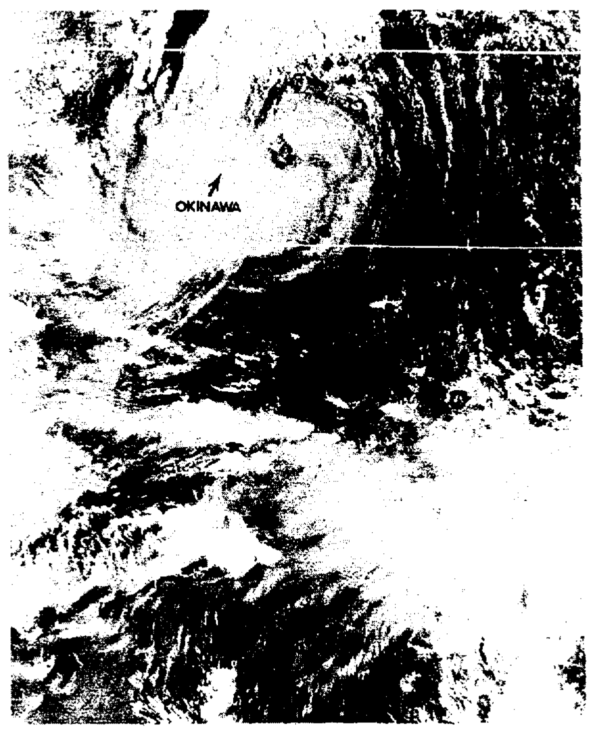


FIGURE 3-17-3. Typhoon Orchid, near maximum intensity, completing the second of three cyclonic loops, 10 September 1980, 0625Z. (TIROSN imagery)

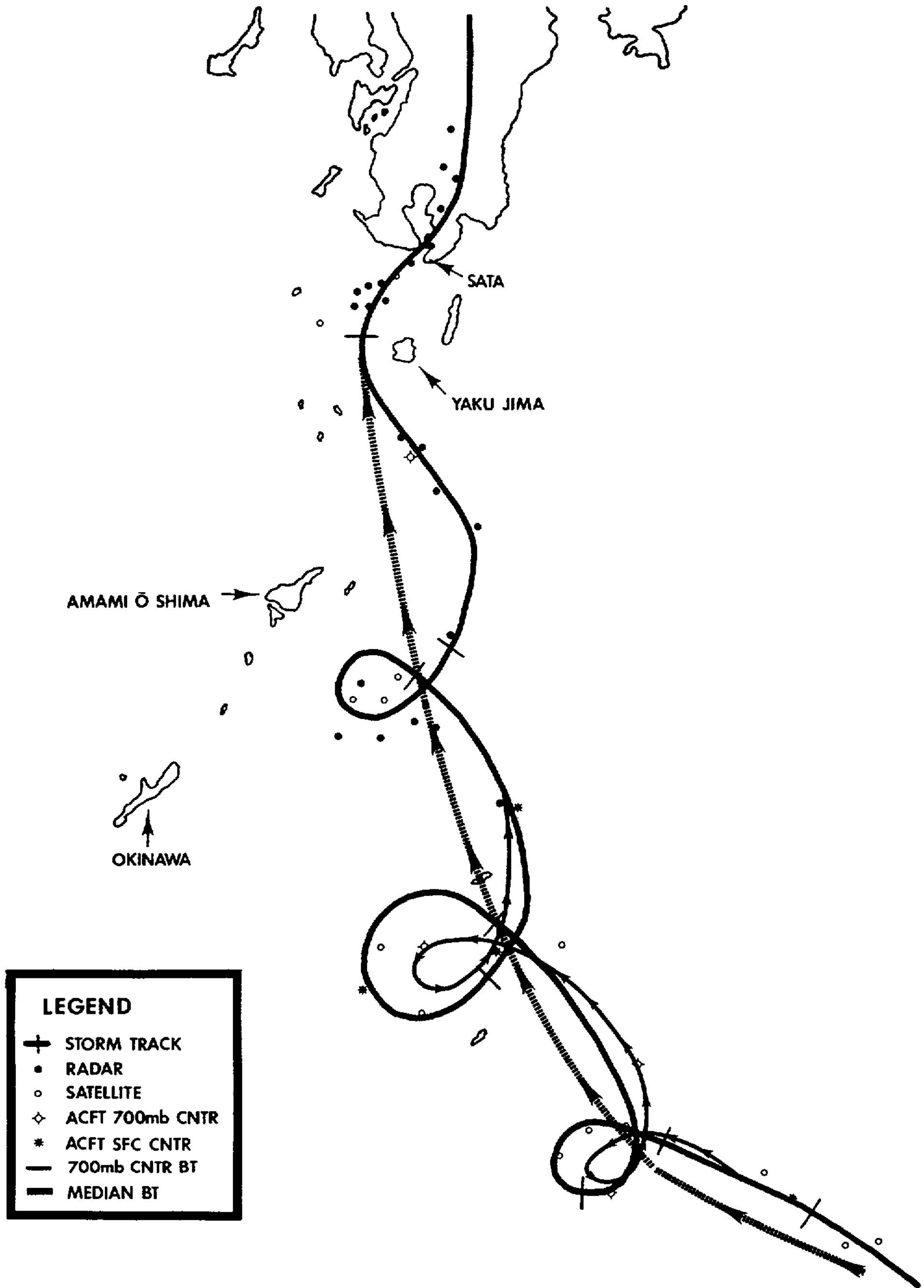
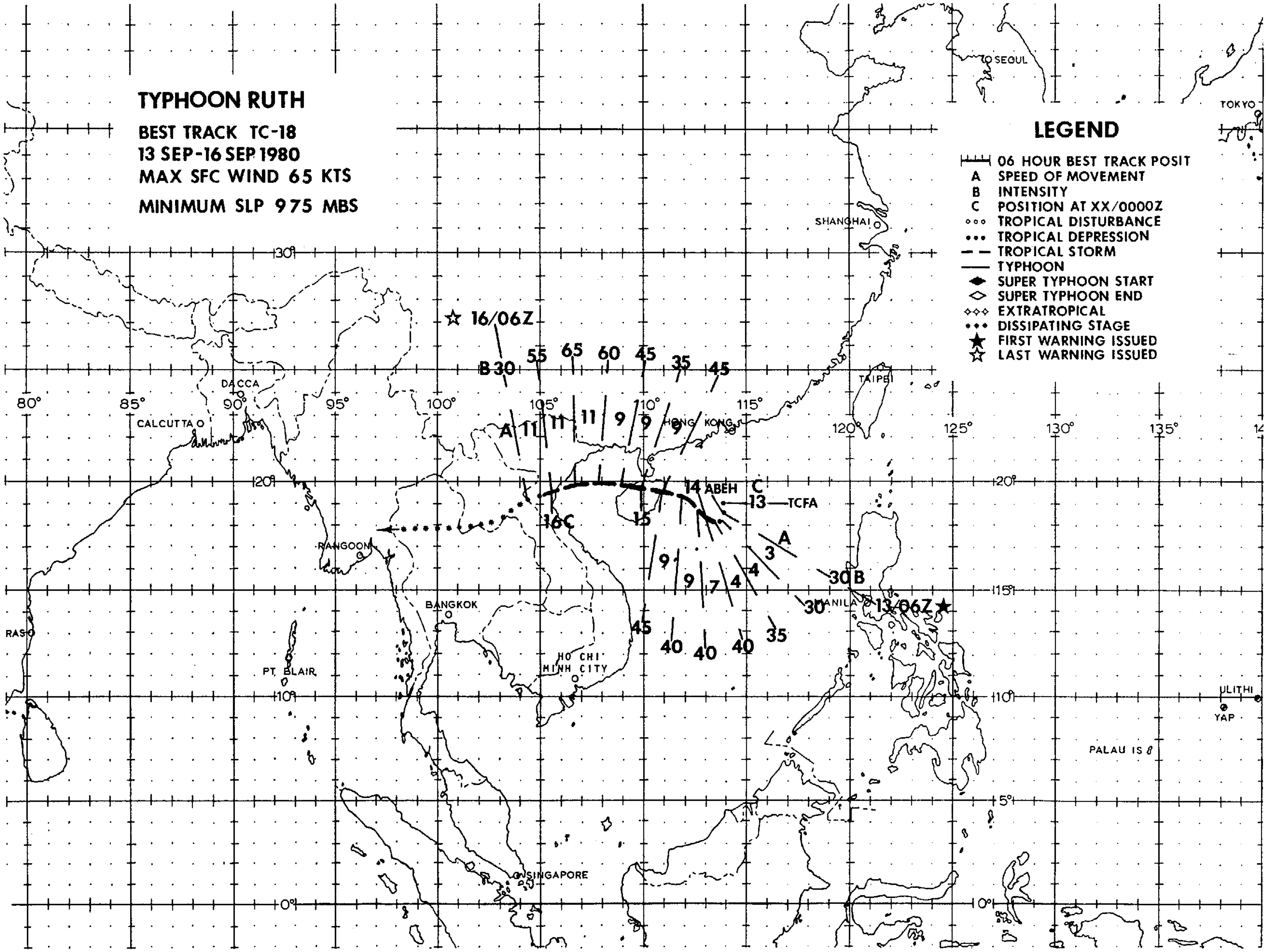


FIGURE 3-17-4. An expanded best track from 090600Z to 110000Z. The figure shows the distribution of fix positions, a partial 700 mb track, and an overall smoothed track.


TYPHOON RUTH
BEST TRACK TC-18
13 SEP-16 SEP 1980
MAX SFC WIND 65 KTS
MINIMUM SLP 975 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
16/06Z
13/06Z
TCFA
ABEH
SEOUL
TOKYO
SHANGHAI
TAIPEI
DACCA
CALCUTTA
HONG KONG
RANGOON
BANGKOK
MANILA
HO CHI MINH CITY
PT. BLAIR
SINGAPORE
ULITHI
YAP
PALAU IS

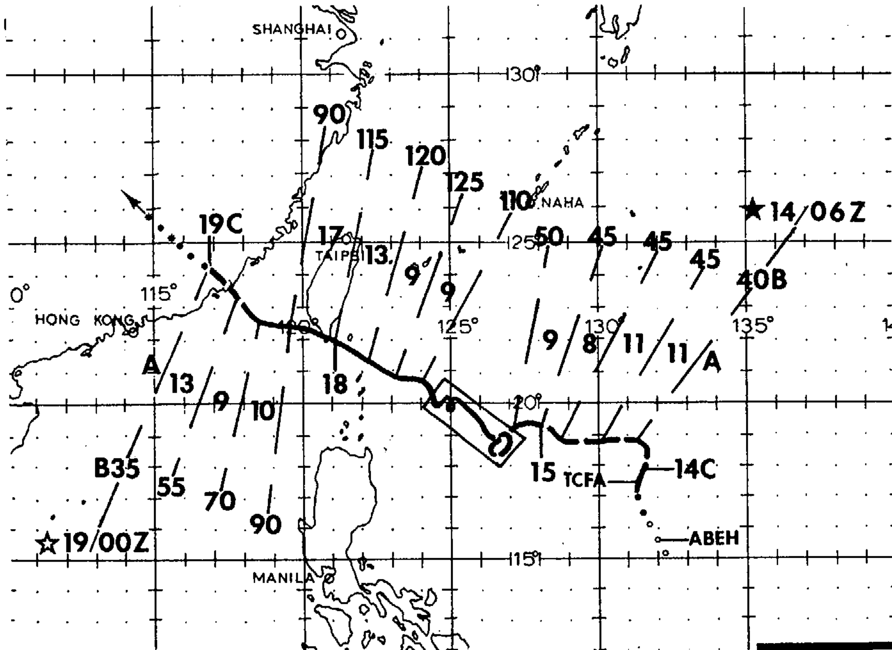
Typhoon Ruth was the second of five typhoons occurring in September. Unlike the other typhoons, Ruth began as a monsoon depression in the South China Sea on 11 September. For two days, the depression remained quasi-stationary with the weak surface circulation embedded in the monsoonal trough. Synoptic data on the 13th, however, indicated that the circulation was intensifying; also, satellite imagery showed that it was forming its own outflow center. As a result, JTWC issued a formation alert at 130047Z. Later satellite data showed that further development had occurred which prompted the first warning to be issued on TD 18 at 130600Z.

During the early phase of development, TD 18 tracked slowly southward, steered by the near surface wind flow. By 131800Z, TD 18 had intensified to Tropical Storm Ruth and had started to track northwestward at an accelerated forward speed of movement. For the rest of her existence, Ruth tracked along the southern periphery of the 500 mb ridge which was centered over southern Mainland China. She reached her first maximum intensity of 45 kt (23 m/sec) prior to landfall over Hai-Nan Tao, but quickly weakened to minimal tropical storm strength while over the island. Ruth entered the Gulf of Tonkin on the 15th and, during her transit, rapidly intensified to typhoon strength with a maximum surface wind of 65 kt (33 m/sec) and a minimum sea-level pressure of 975 mb.

Brand (1970) summarized the finding of Perlroth (1969) who showed that vertical temperature differences between the ocean surface and the 200 ft (61 m) water depth have an important effect on development of tropical cyclones. Perlroth reported that approximately 90% of the tropical cyclones that reached hurricane intensity in the equatorial Atlantic from 1901-1965 occurred where the climatological difference between the ocean surface temperature and the 200 ft (61 m) temperature was 3.9C degrees or less. Climatology for September shows that the Gulf of Tonkin has warm sea surface temperatures (29C) and a vertical temperature gradient along Ruth's track which is within the constraints reported by Perlroth for intensification to typhoon strength. Thus, the northern portion of the Gulf of Tonkin can serve as a sufficient heat source for tropical cyclones, such as Ruth, to intensify when conditions are favorable. This apparently is true despite the fact that the Gulf is surrounded on three sides by land.

Ruth made landfall at 160000Z south of Thanh Hou, Vietnam. Nearly half a million people were left homeless with 106 persons known dead or missing in Vietnam. Ruth also caused massive crop damages and interrupted communications in the area.

After landfall, Ruth again weakened and dissipated as a significant tropical cyclone. The remnants of Ruth tracked west-southwestward for the next two days and dissipated over the Bilauk-taung Range along the border of Burma and Thailand.



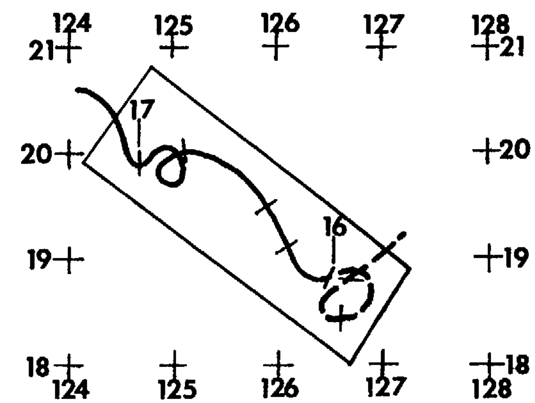
LEGEND

- 06 HOUR BEST TRACK POSIT
- A SPEED OF MOVEMENT
- B INTENSITY
- C POSITION AT XX/0000Z
- ○○○ TROPICAL DISTURBANCE
- ••• TROPICAL DEPRESSION
- — — TROPICAL STORM
- —— TYPHOON
- ◆ SUPER TYPHOON START
- ◇ SUPER TYPHOON END
- ◇◇◇ EXTRATROPICAL
- *** DISSIPATING STAGE
- ★ FIRST WARNING ISSUED
- ☆ LAST WARNING ISSUED

TYPHOON PERCY

BEST TRACK TC-19
14 SEP-19 SEP 1980
MAX SFC WIND 125 KTS
MINIMUM SLP 919 MBS

DTG	SPEED	INTENSITY
	06	
15/12Z		50
	07	
15/18Z		55
	10	
16/00Z		65
	04	
16/06Z		70
	04	
16/12Z		75
	09	
16/18Z		85
	11	
17/00Z		95
	10	



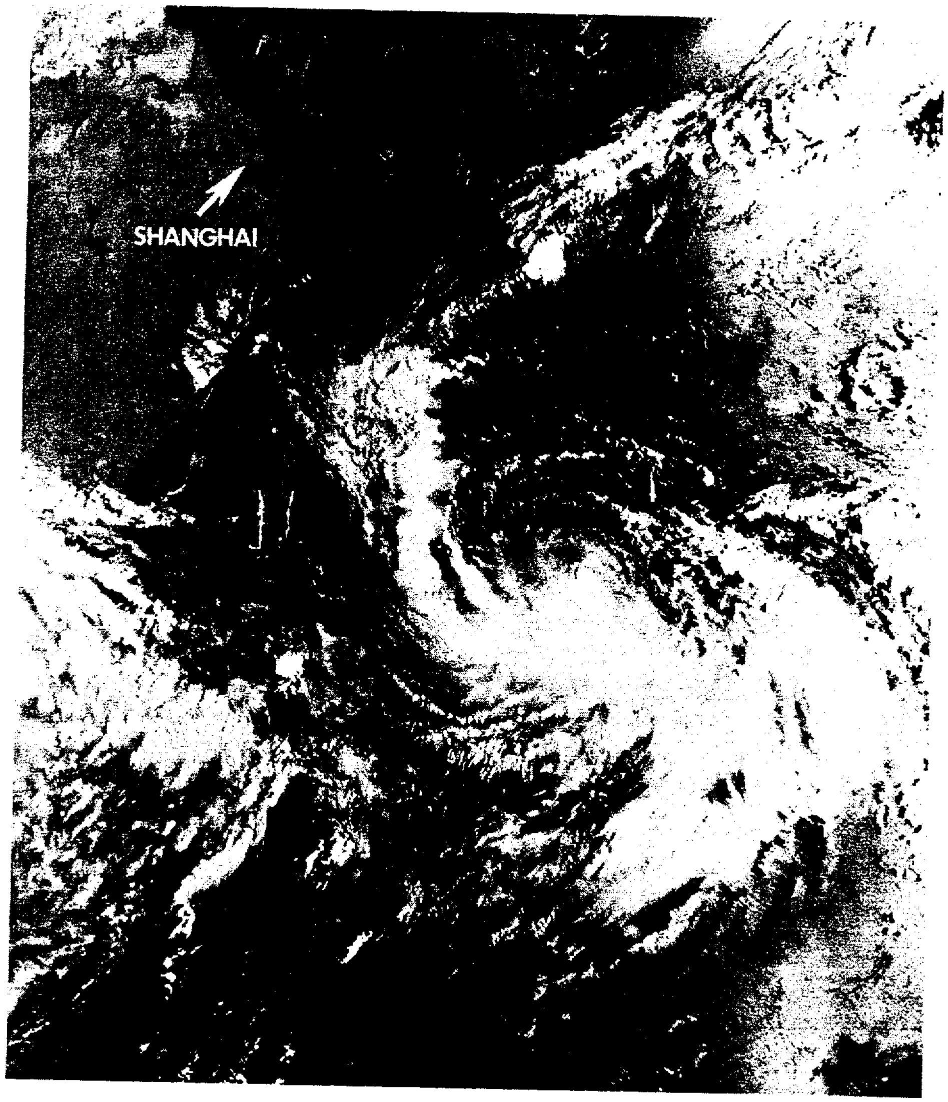


FIGURE 3-19-1. Tropical Storm Percy at 45 kt (23 m/sec) intensity, 14 September 1980, 2333Z. Percy's low-level center at that time was partially exposed on the northeastern edge of a large convective area. (NOAA6 imagery)

The disturbance which eventually developed into the tenth typhoon of 1980 became evident on satellite imagery at 130600Z September 1980 as a focal point of cumulus banding. This development occurred at the base of a mid-tropospheric, mid-latitude trough that extended south of Japan along 133E. With further intensification likely, a Tropical Cyclone Formation Alert (TCFA) was issued by JTWC at 132157Z. An aircraft reconnaissance investigation soon afterwards found a well-developed closed circulation with 1500 ft (457m) flight level winds of 35 kt (16 m/sec) and a minimum sea-level pressure of 992 mb. Upon receipt of these data, the first warning on Tropical Storm Percy was issued at 140600Z.

By that time, the mid-tropospheric, mid-latitude trough had moved eastward, and Percy's track, in response, shifted from a northward to a more climatological west-northwestward track as the ridge extended eastward north of Percy. Percy's development was slow at this time as scatter between satellite and aircraft fixes indicated that he was poorly aligned in the vertical. Convection on satellite imagery was displaced southwest of Percy's surface center. The mid-tropospheric ridge that extended from eastern China along 28N to a position north of Okinawa also restricted upper-level outflow in Percy's northern semicircle. Between 15 and 17 July, however, reconnaissance aircraft consistently reported decreasing heights and increasing temperatures near the 700 mb center (Fig. 3-19-1).

During this time, Percy decelerated and began moving erratically. He eventually completed two tight cyclonic loops while intensifying to typhoon strength by 160000Z. Shortly afterward, at 160248Z, aircraft reconnaissance reported that the eyewall had become fully enclosed. JTWC forecast significant intensification for the next 36 hours as the Theta E (Θ_e) forecast intensity aid (see summary for Super Typhoon Kim) indicated an approaching intersection of the 700 mb equivalent potential temperature trace and the minimum sea-level pressure trace.

During Percy's period of erratic movement, there was speculation that a Fujiwhara interaction might develop between Percy and then Tropical Storm Sperry, which was located 800 nm (1575 km) to the east. A comparison of the post-analysis best tracks for Percy and Sperry shows that the two vorticies were never close enough for an interaction to occur (Brand, 1968).

Interestingly, Percy did dominate much of the low-level circulation pattern between the two systems. A reconnaissance aircraft mission flown between Percy and Sperry indicated that the wind shift from southerly to northerly flow did not occur until about 100 nm (185 km) west of Sperry's surface center (Fig. 3-19-2).

Significant intensification did take place as Percy's eye gradually grew tighter and sea-level pressure continued to fall (Fig 3-19-3). An aircraft fix at 171306Z indicated a 700 mb height of 2387 m, which extrapolates to a sea-level pressure of 919 mb and supports maximum sustained surface winds of 125 kt (64 m/sec). Due to his proximity to Taiwan, however, Percy then began to slowly weaken. A subsequent reconnaissance aircraft at 171616Z reported a rise in the 700 mb level to 2407 m. Approximately 9 hours later, Typhoon Percy made landfall on the extreme southern tip of Taiwan.

FIGURE 3-19-2. Percy just prior to reaching typhoon strength during the period of erratic movement over the Philippine Sea, 15 September 1980, 2311Z. (NOAA6 infrared imagery)

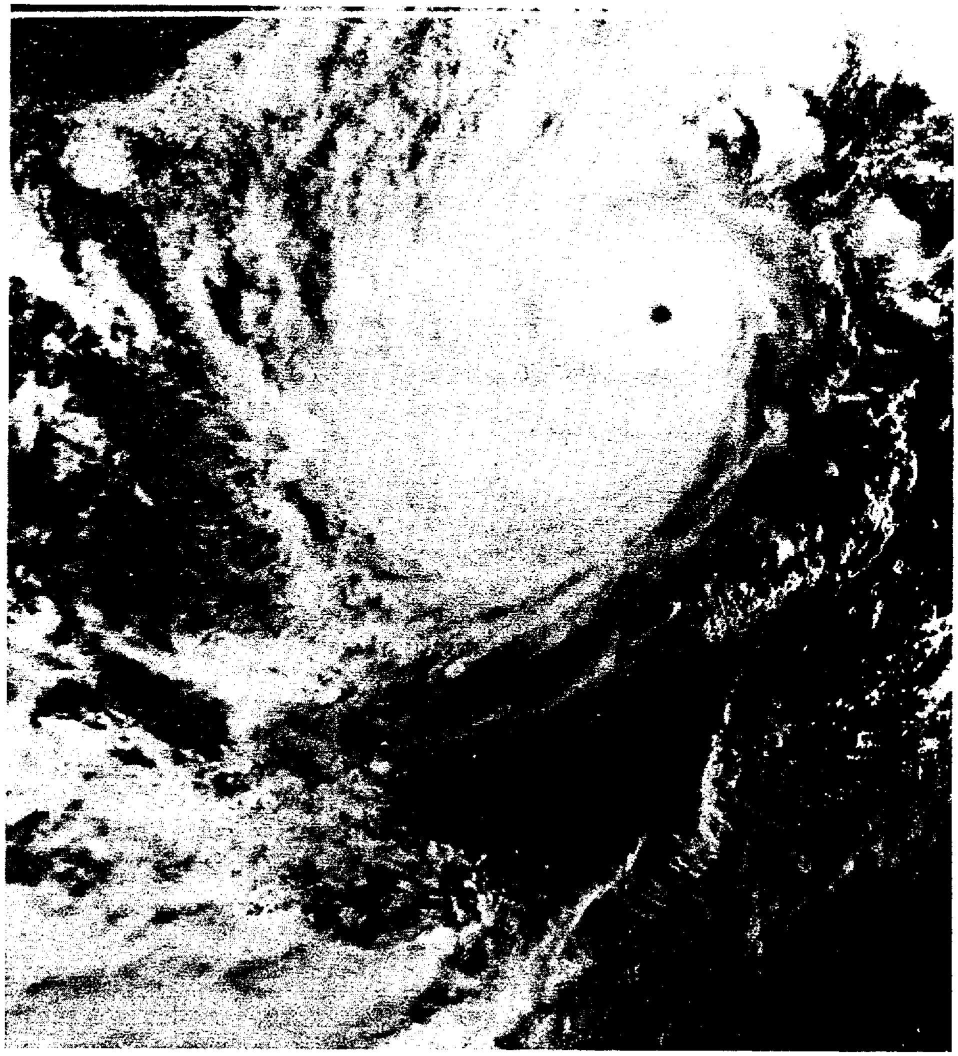
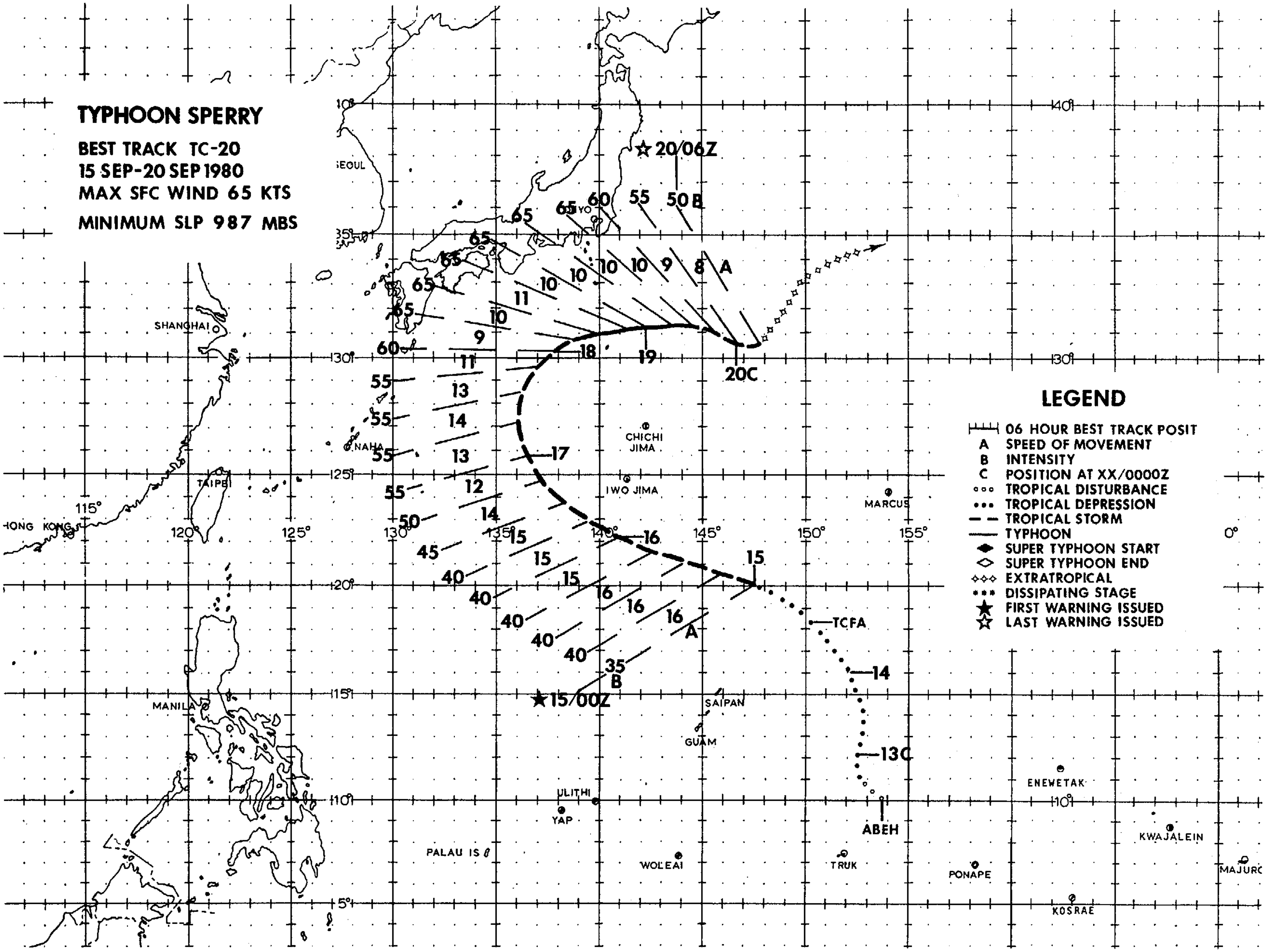


FIGURE 3-19-3. Typhoon Percy with 110 kt (57 m/sec) maximum surface winds while intensifying rapidly to near super typhoon strength, 17 September 1980, 0645Z. (Tiros N imagery)

The island of Taiwan so disrupted Percy's low-level inflow that he was never able to significantly re-intensify. Despite emerging over the Formosa Strait, Percy continued to weaken almost as quickly as he had intensified only a day earlier. By 182100Z he had made landfall on the coast of China about 240 nm (444 km) east of Canton, with estimated maximum sustained surface winds of 45 kt (23 m/sec). Percy continued to track inland and dissipated several hours later over the mountains of Mainland China.

Newspaper accounts of Typhoon Percy's landfall over southern Taiwan indicated 7 dead and 16 injured. Heavy rain accompanying Percy damaged 140 homes, flooded rice fields, and destroyed banana crops.


TYPHOON SPERRY
BEST TRACK TC-20
15 SEP-20 SEP 1980
MAX SFC WIND 65 KTS
MINIMUM SLP 987 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
20/06Z
15/00Z
TCFA
ABEH
SEOUL
SHANGHAI
NAHA
TAIPEI
HONG KONG
MANILA
CHICHI JIMA
IWO JIMA
MARCUS
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
ENEWETAK
KWAJALEIN
KOSRAE
MAJURO

Typhoon Sperry developed in the monsoon trough east-southeast of Guam. The disturbance was first reported in the Significant Tropical Weather Advisory (ABEH PGTW) on 12 September as an area of showers and thundershowers. Sparse synoptic data did not indicate that a surface circulation existed at that time. However, the upper-air pattern was favorable for continued development. Sperry developed slowly and was described in the ABEH PGTW on 14 September as a large surface circulation with little organized convection. A well-defined upper-level anticyclone, which provided a good outflow mechanism for continued development, existed over Sperry.

The initial warning for Tropical Depression 20 was issued at 150000Z. Post-analysis indicates that Sperry had actually attained tropical storm strength of 35 kt (18 m/sec) by that time. The 141200Z 500 mb analysis (Fig. 3-20-1) and the 72-hour numerical forecast series (see Fig. 3-20-2) suggested that a straight forecast track toward Kyushu, Japan was most likely because the forecast series built the subtropical ridge northwestward toward Japan. Thus, on the initial warning, Sperry was forecast to track along the southern periphery of the 500 mb subtropical ridge. An early recurvature track was not considered likely due to the forecast intensification of the subtropical ridge.

By 160000Z, it was evident that the subtropical ridge was not building as forecast. Southerly steering flow was evident south and east of Japan. Sperry was being steered by the mid-level southeasterly flow and was ex-

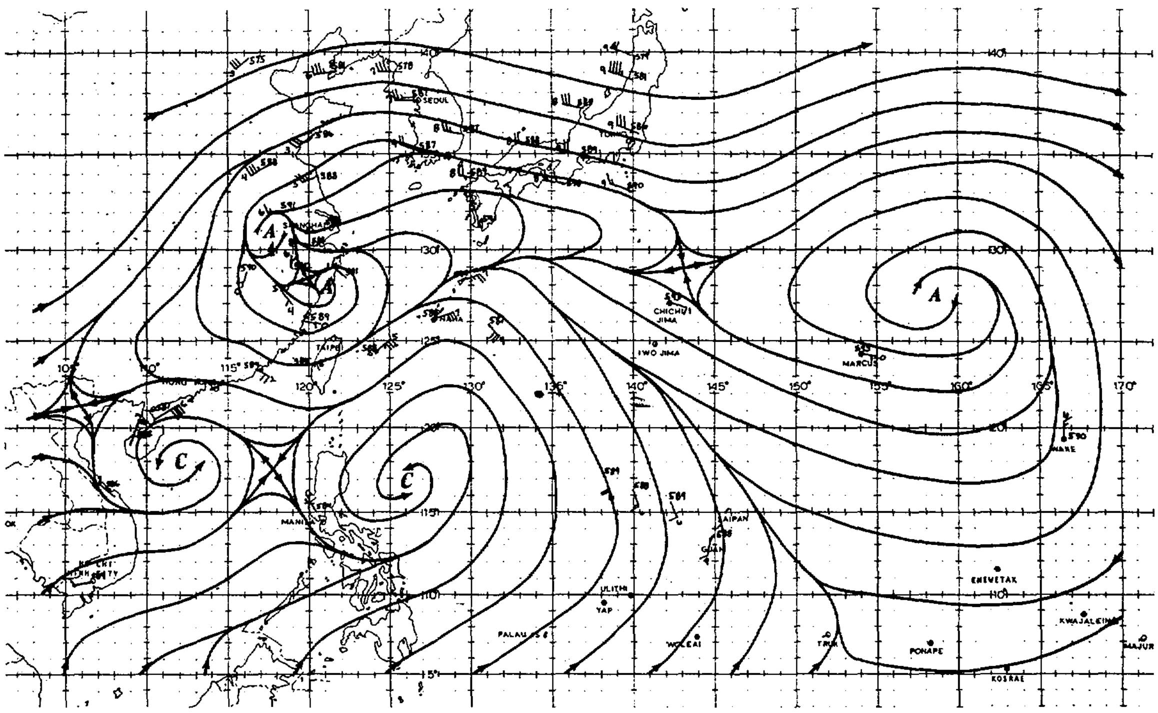


FIGURE 3-20-1. The 141200Z September 1980 500 mb streamline analysis. Wind speeds are in knots.

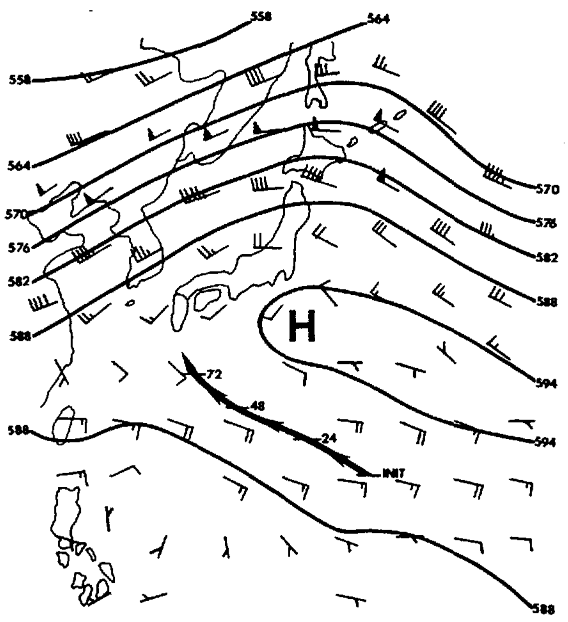


FIGURE 3-20-2. The 48 hour 500 mb numerical forecast chart based on the 141200Z September 1980 computer analysis. JTWC's 150000Z September forecast track for Sperry is also indicated, from initial position to the 72 hour forecast position.

pected to continue to follow a north-northwestward track until he moved north of the ridge axis. Then strong mid-level westerlies were expected to dominate. Therefore, a recurvature track and a weakening tendency over Japan was forecast. This change to a recurvature track was supported by the 161200Z 500 mb analysis (Fig. 3-20-3) and the 72-hour numerical forecast series (see Fig. 3-20-4). Sperry did, in fact, recurve, but significantly south of Japan as the subtropical ridge retreated to the southeast. This discussion of the forecast tracks for Sperry illustrates the difficulties that JTWC encounters both in analyzing the axis of the subtropical ridge in data sparse regions and interpreting the guidance from numerical forecasts for the same region.

As Sperry began to recurve on the 17th, the estimated maximum surface wind speeds were consistently higher than supported by the maximum wind/minimum sea-level pressure (MSLP) relationship of Atkinson and Holliday (1977). Maximum winds of 65 kt (33 m/sec) and MSLPs of 992 mb were observed by aircraft reconnaissance. A MSLP of 992 mb corresponds to a maximum wind of 45 kt (see Fig. 3-05-2). These stronger winds were probably due to an increased pressure gradient resulting from the higher environmental pressures at subtropical latitudes.

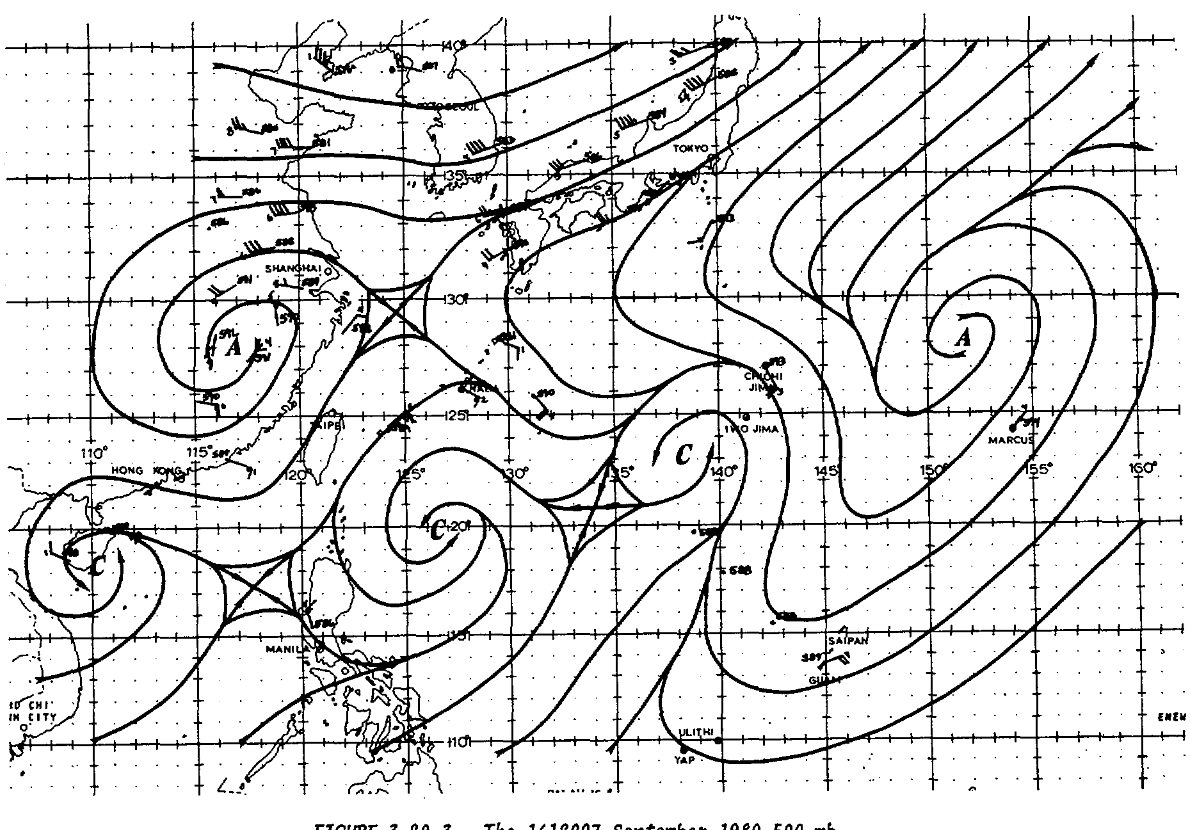


FIGURE 3-20-3. The 161200Z September 1980 500 mb streamline analysis. Wind speeds are in knots.

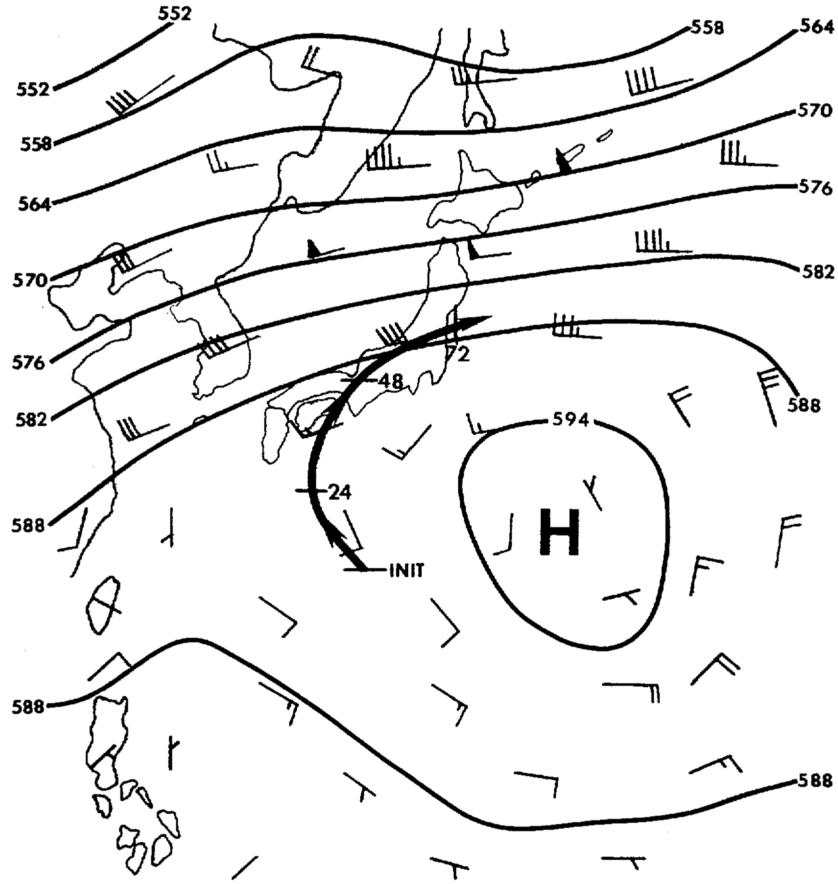
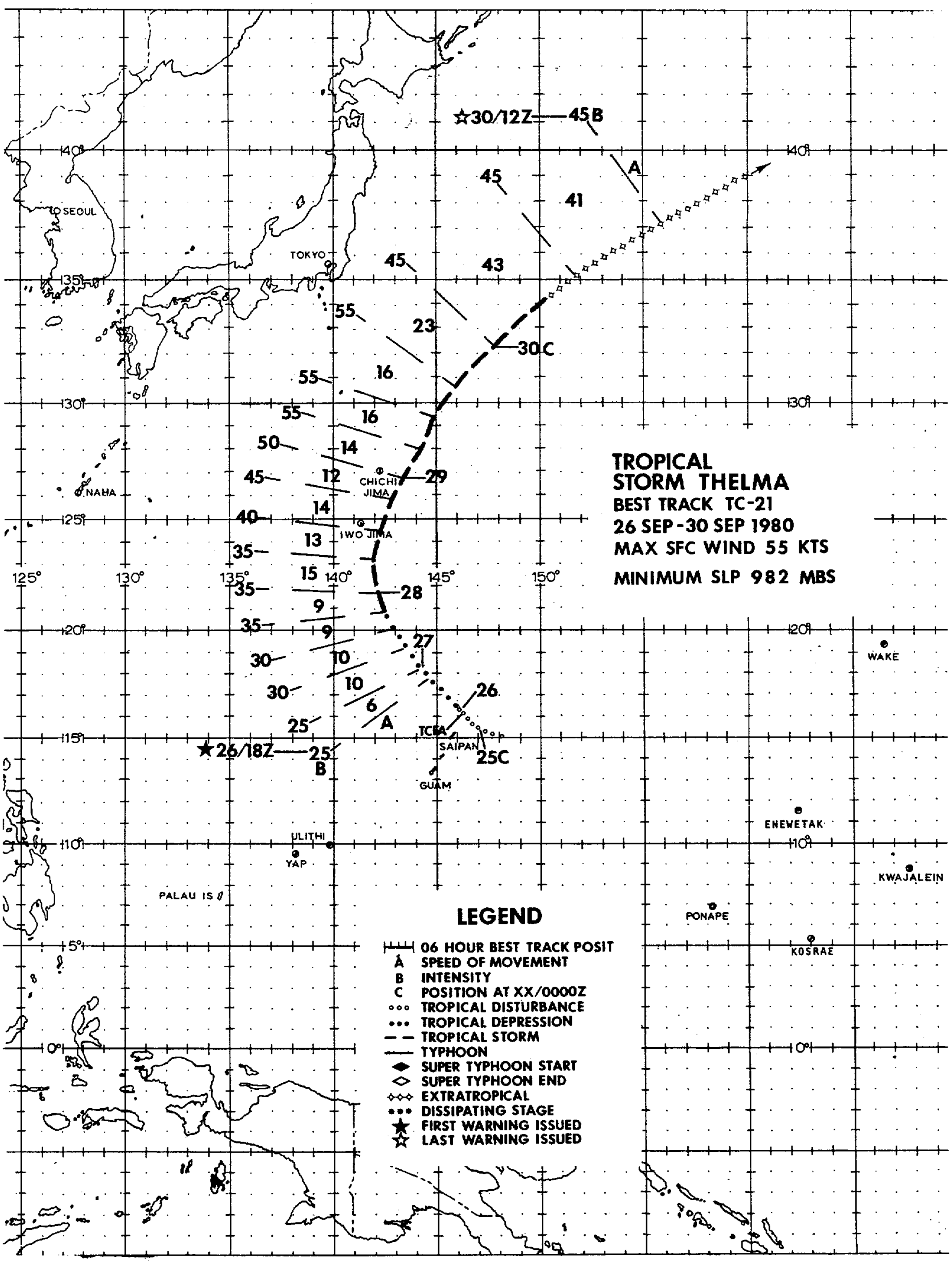


FIGURE 3-20-4. The 48 hour 500 mb numerical forecast chart based on the 161200Z September computer analysis. JTWC's 170000Z September forecast track for Sperry is also indicated, from initial position to the 72 hour forecast position.

Sperry did not begin to weaken significantly until the 19th because his eastward movement kept him over warmer water for a longer period of time and also kept him south of the strong mid- to upper-level westerlies which would have weakened him due to strong vertical wind shear.


TROPICAL STORM THELMA
BEST TRACK TC-21
26 SEP-30 SEP 1980
MAX SFC WIND 55 KTS
MINIMUM SLP 982 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
26/18Z
30/12Z
SEOUL
TOKYO
NAHA
CHICHI JIMA
IWO JIMA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WAKE
ENEWETAK
KWAJALEIN
PONAPE
KOSRAE

Thelma began as a disturbance in the monsoon trough approximately 100 nm (185 km) north-northeast of Saipan. Although Thelma's forecast track posed few problems, analysis of the cyclone was not straight-forward due to Thelma's very abnormal structure.

Satellite imagery began showing intensification of convective activity in the maximum cloud zone within the latitude belt 10-20N from southeast Asia to the Marshall Islands on 23 September. A significant area of convection developed in the monsoon trough near Saipan early on 23 September. This convection gradually increased and began suggesting the presence of a surface circulation later on the same day. Data at 500 mb, however, indicated that the curvature in the cloud signature was associated with a mid-tropospheric circulation as no circulation was apparent from surface/gradient wind data. A circulation at the surface/gradient level was finally analyzed just south of Guam at 241200Z. By 250000Z, the low-level circulation that produced Thelma was analyzed over the Northern Mariana Islands (Fig. 3-22-1). This circulation continued to develop and drift westward while becoming the dominant circulation in that portion of the monsoon trough.

From her onset, Thelma did not display classical tropical storm characteristics. Height gradients observed by reconnaissance aircraft at 700 mb were very flat; thus maximum winds near the center were significantly lower than suggested by the central sea-level pressure (Atkinson and Holliday, 1977). Also, the maximum wind band was some distance away from the center, and the 700 mb temperature field showed higher temperatures outside the cyclone center for most of the early aircraft penetrations. Table 3-3 presents a summation of aircraft data and highlights the points presented above. Figure 3-22-2 shows the reconnaissance data plot for the last daylight penetration of Thelma. The wind field and other data presented on the plot are fairly representative of Thelma's entire life.

As previously stated, Thelma's track presented no real problems. Streamline analyses at 500 mb showed that Thelma developed just south of a break in the subtropical ridge. After following a northwesterly course, Thelma first turned northward and almost immediately thereafter began to track northeastward. Cyclogenesis/frontogenesis occurred simultaneously in a baroclinic zone that persisted throughout Thelma's life in the area from Okinawa northeastward to a point off the coast of Japan. The continual presence of this surface trough appears to be one of the factors that directed Thelma's northeastward movement (Fig. 3-22-3). Upper-level steering was provided by relatively strong westerlies which reached south to the Bonin and Volcano Islands. During her northeast trek, Thelma reached maximum intensity of 55 kt (28 m/sec). Further intensification was probably suppressed by restrictions on her upper-level outflow. Thelma continued to accelerate toward the northeast and transitioned into an extratropical low pressure system by 0400Z on 30 September.

TABLE 3-3

	Maximum Temperature		Central Pressure	Intensity (KT)		Bearing/Range
Date/Time	Inside Center °C	Outside Center °C	MB	Observed Surface	Atkinson/ Holliday	of MAX FLT LVL Wind (DEG/NM)
27/1530Z	12	13	993	N/O	42	100/136
28/0308Z	14	15	989*	30	48	320/110
28/1418Z	12	15	987	N/O	50	240/40
29/0258Z	14	10	982	55	57	350/110
29/1500Z	13	13	981*	N/O	59	120/150

Aircraft data extracted from detailed vortex messages and peripheral data observations. (Asterisks indicate central pressure extrapolated from 700 mb data.)

FIGURES 3-22-1, 3-22-2 and 3-22-3 are on the following pages.

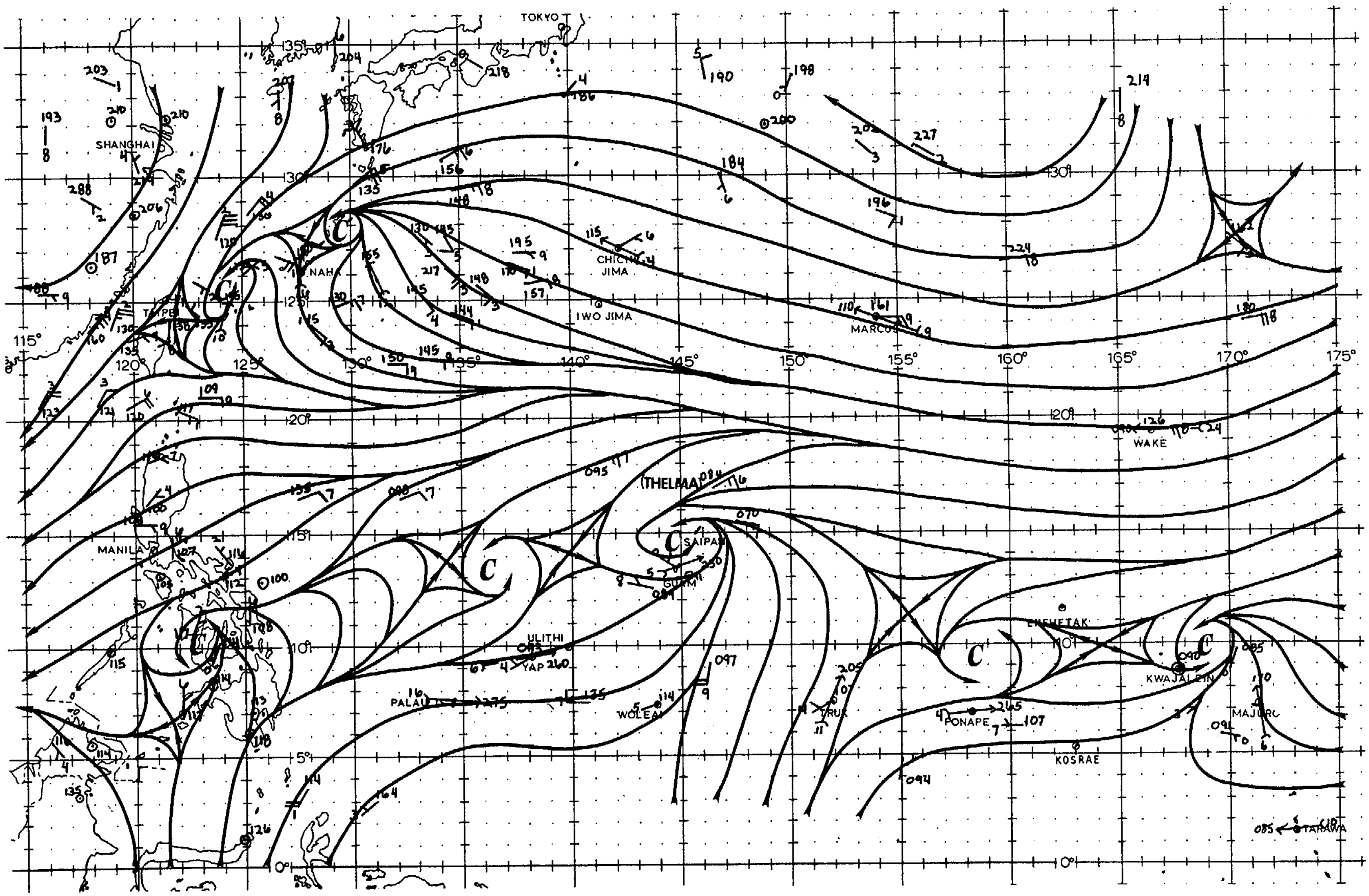


FIGURE 3-22-1. The 250000Z September 1980 surface (←—— ll)/gradient-level (ddd←— ₍ff) wind data and streamline analysis. Wind speeds are in knots.

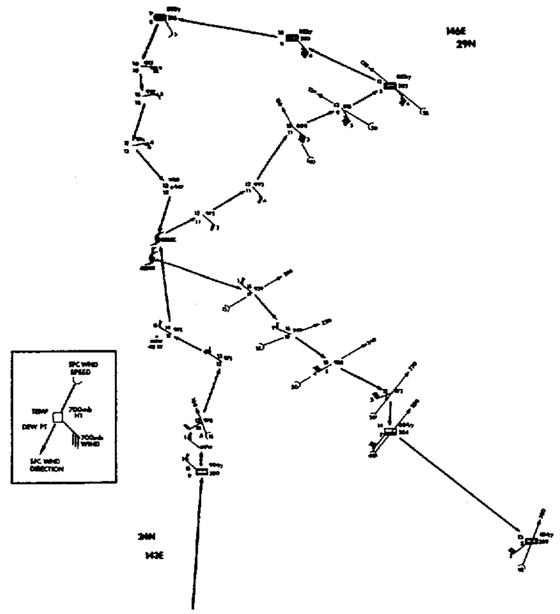


FIGURE 3-22-2. Plot of aircraft reconnaissance data for the 290300Z September 1980 fix of Tropical Storm Thelma.

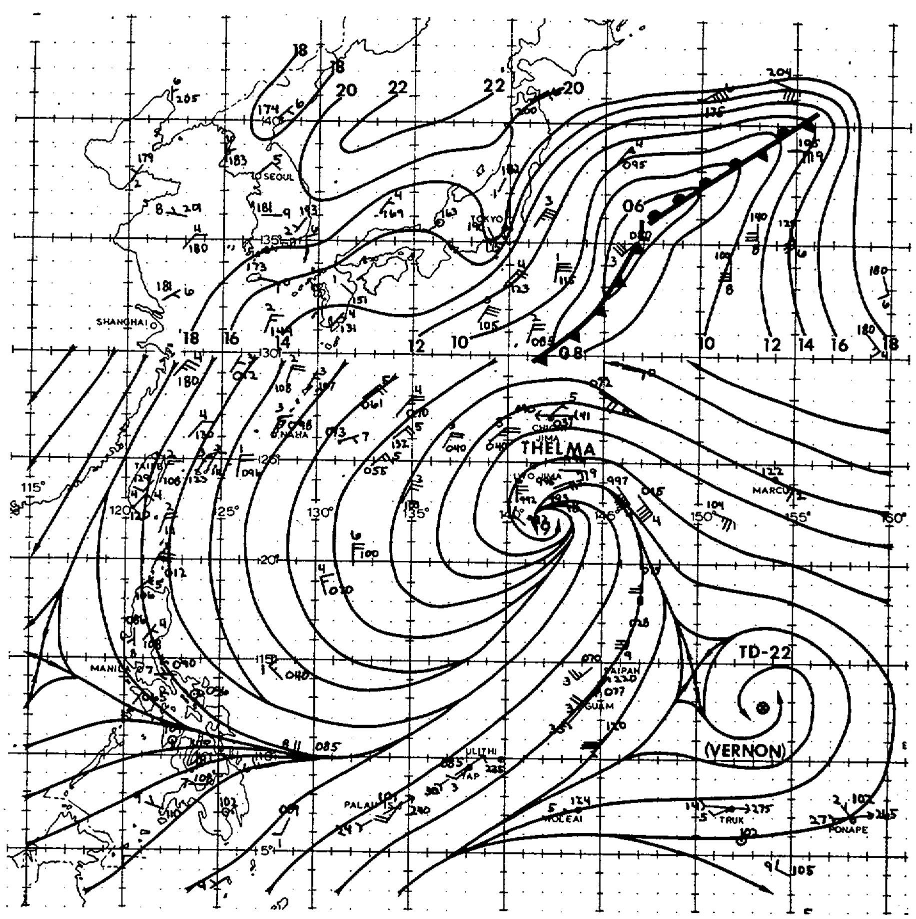
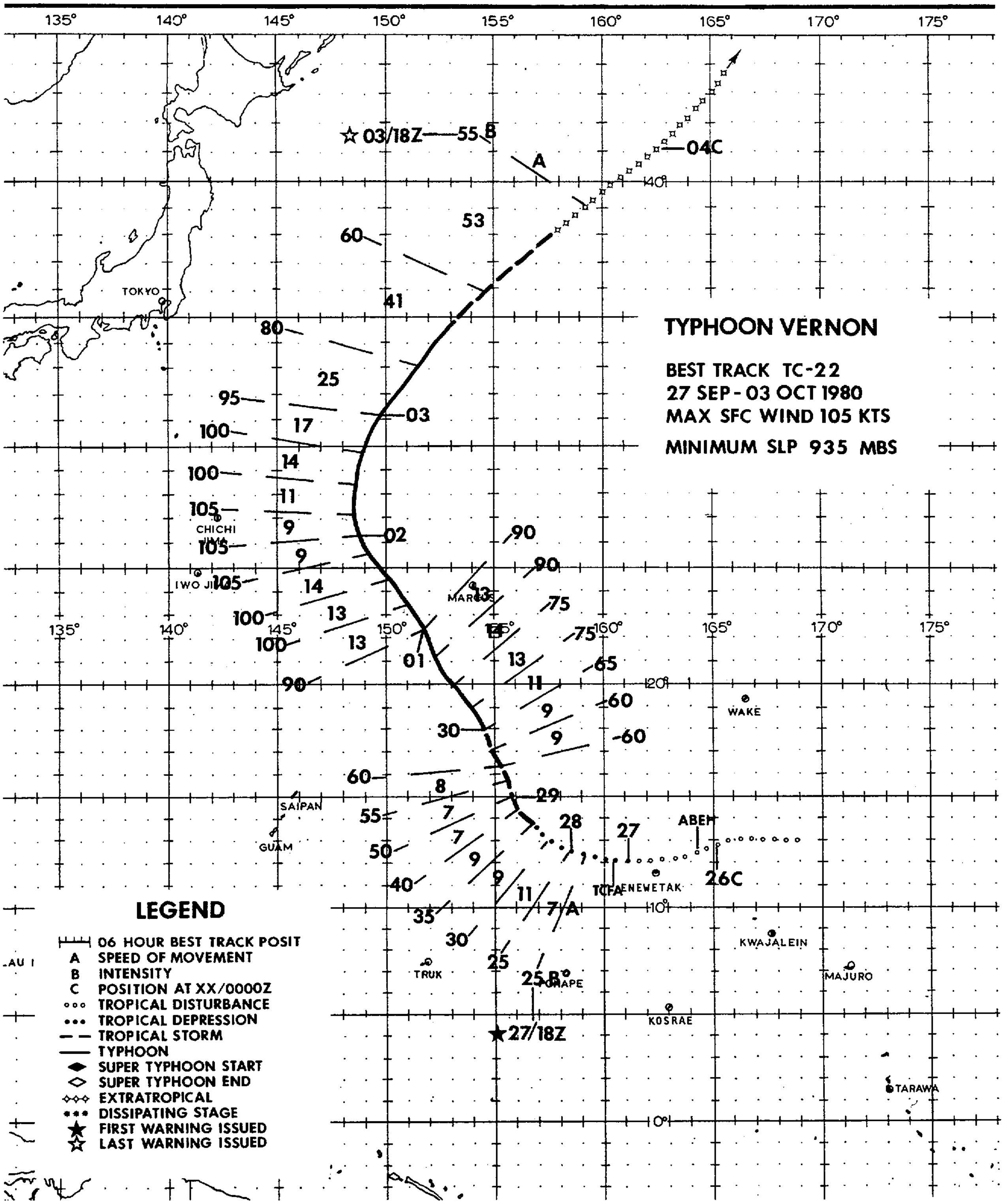


FIGURE 3-22-3. The 280000Z September 1980 surface (———) /gradient-level (ddd←——(ff) wind data and streamline analysis. Winds speeds are in knots.


TYPHOON VERNON
BEST TRACK TC-22
27 SEP - 03 OCT 1980
MAX SFC WIND 105 KTS
MINIMUM SLP 935 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
03/18Z
27/18Z
TOKYO
CHICHI JIMA
IWO JIMA
MARCUS
SAIPAN
GUAM
TRUK
PONAPE
KOSRAE
ENEWETAK
WAKE
KWAJALEIN
MAJURO
TARAWA
TCFA
ABEH
04C
03
02
01
30
29
28
27
26C

During the latter part of summer, tropical cyclone activity in the northwestern Pacific reaches its peak. Multiple circulations develop within the Near Equatorial Trough and two (or more) cyclones of tropical storm or typhoon strength often exist at the same time.

If one tropical cyclone is located to the northwest of another developing circulation, it usually dominates and prevents the system to the southeast from intensifying as rapidly as it normally would. This is due primarily to the upper-level outflow from the system to the northwest which enhances the climatological northwesterlies and restricts the outflow channels of the cyclone located to the southeast. The cyclone to the northwest is also, generally, the older of the two and has the opportunity to establish control of the low-level inflow. The development of the system to the southeast is, therefore, delayed until the other cyclone either weakens or moves far enough away from the tropics that its influence becomes insignificant (see Roger and Tip, 1979 and Lex and Marge, 1980). Typhoon Vernon and Tropical Storm Thelma engaged in just such an interaction during the end of September and beginning of October.

Vernon was first observed, as an area of increased thunderstorm activity, about 200 nm (370 km) northeast of Eniwetok Atoll on 26 September. Initial movement was westward at about 7 kt (13 km/hr). As the convection continued to organize, a Tropical Cyclone Formation Alert (TCFA) was issued at 270600Z, and the first warning followed 12 hours later.

During that period, Tropical Storm Thelma was developing north of Guam. Thelma, although never more than tropical storm strength, nonetheless had a huge associated cyclonic circulation pattern which extended to at least the 500 mb level and covered most of the area between the Philippine Islands and Guam, and as far north as southern Japan. Because Thelma covered such a large area and was located to the northwest of Vernon, she robbed him of strong low-level inflow and restricted the upper-level outflow in his northwest semicircle in the manner described above.

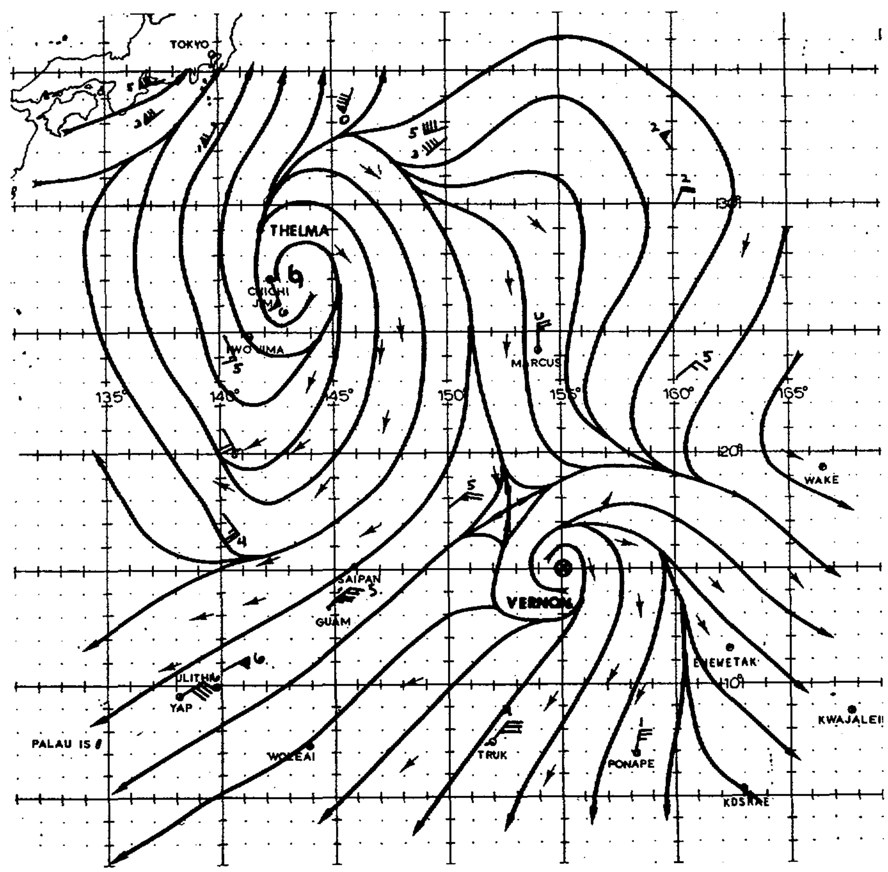


FIGURE 3-22-1. Upper-level streamline analysis (near the 200 mb level) at 290000Z September 1980. Data are rawinsonde and aircraft winds ( ) and satellite-derived wind vectors ( ). Winds are in knots.

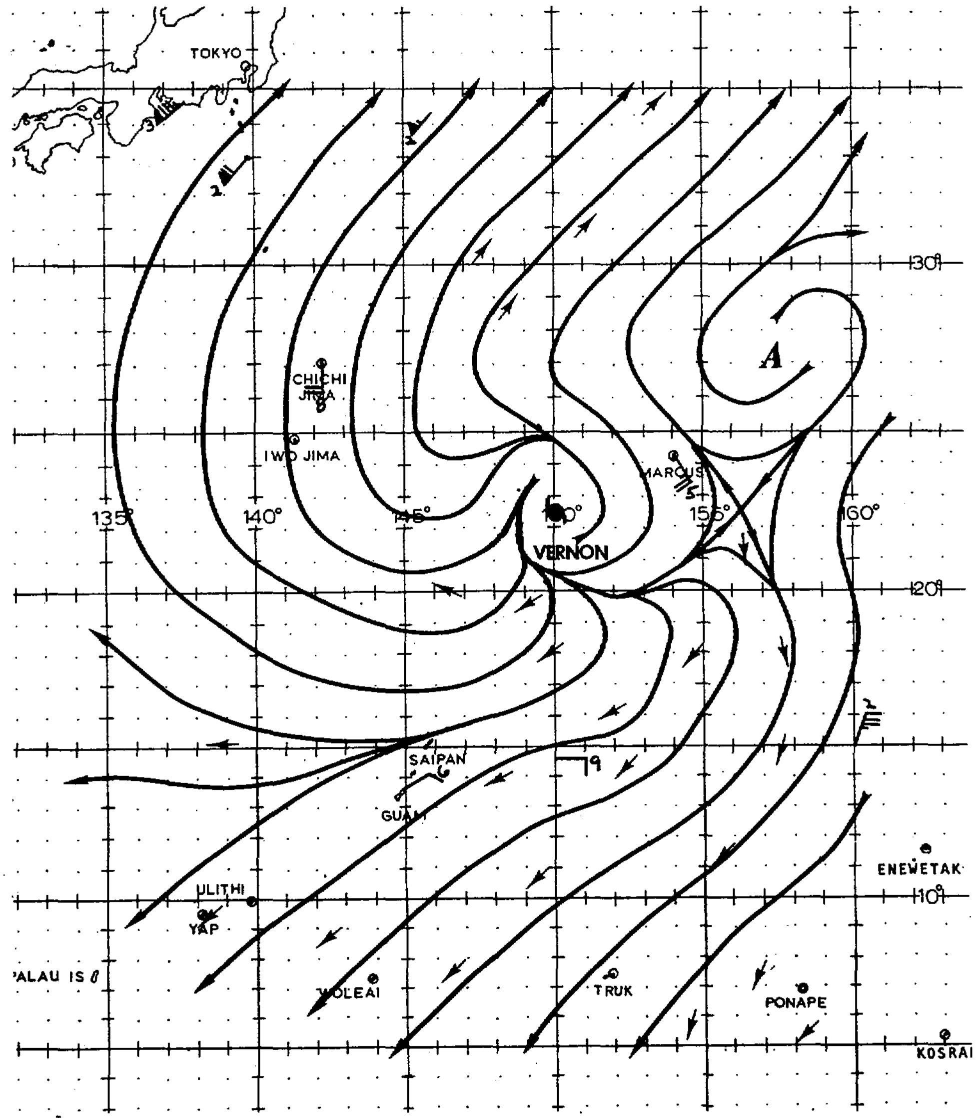


FIGURE 3-22-2. Upper-level streamline analysis (near the 200 mb level) at 011200Z October 1980. Data are rawinsonde and aircraft winds ( ) and satellite-derived wind vectors ( ). Winds are in knots.

After the 28th, Vernon began tracking more northwestward as he moved into the mid-level trough which was associated with Thelma. Thelma helped to maintain this trough throughout her lifetime as indicated by reconnaissance aircraft and the few island reporting stations in the vicinity. Vernon was steered by the southeasterly winds on the east side of this trough until 021200Z October. At that time, he came in contact with the southern extension of the mid-latitude jet-stream which accelerated him to the north-east, eventually to 53 kt (98 km/hr).

Figures 3-22-1 and 3-22-2 show a dramatic change which took place in the upper-level flow pattern; the outflow from Thelma initially restricted Vernon's outflow in his northwest semicircle (Fig. 3-22-1), but by 011200Z, Thelma had moved off to the northeast. This opened up an outflow channel to the north and northwest (Fig. 3-22-2) which enabled Vernon to reach his maximum intensity of 105 kt (54 m/sec) (Fig. 3-22-3). Without the influence of TS Thelma, Vernon most probably would have reached maximum intensity earlier and maintained it longer.

Vernon made the transition to an extra-tropical system quite rapidly. Satellite imagery showed that he lost almost all of his heavy convection between 031200Z and 031800Z. Ship reports off the coast of Japan indicated that the remnants of Vernon continued to maintain gale force winds until 5 October.

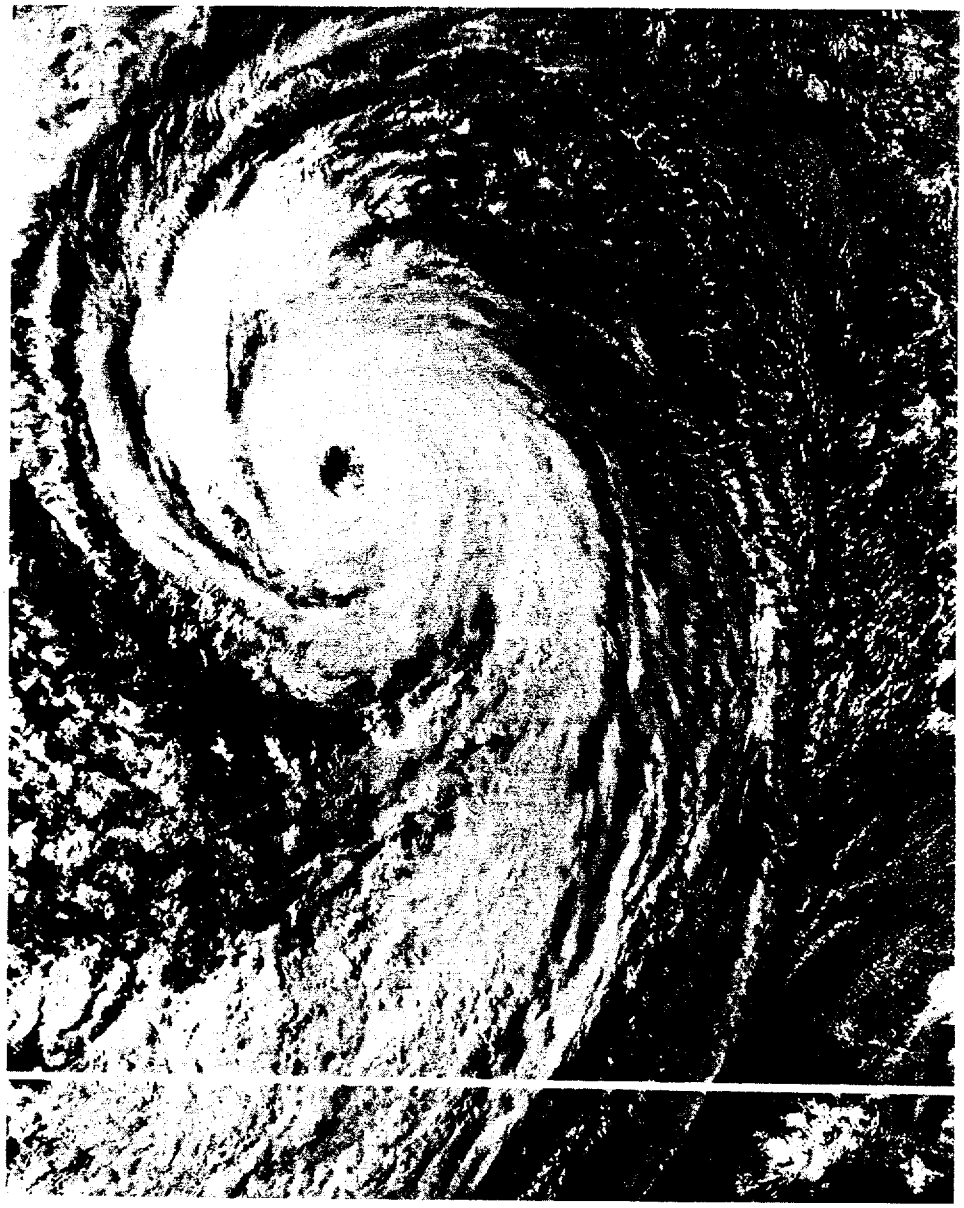
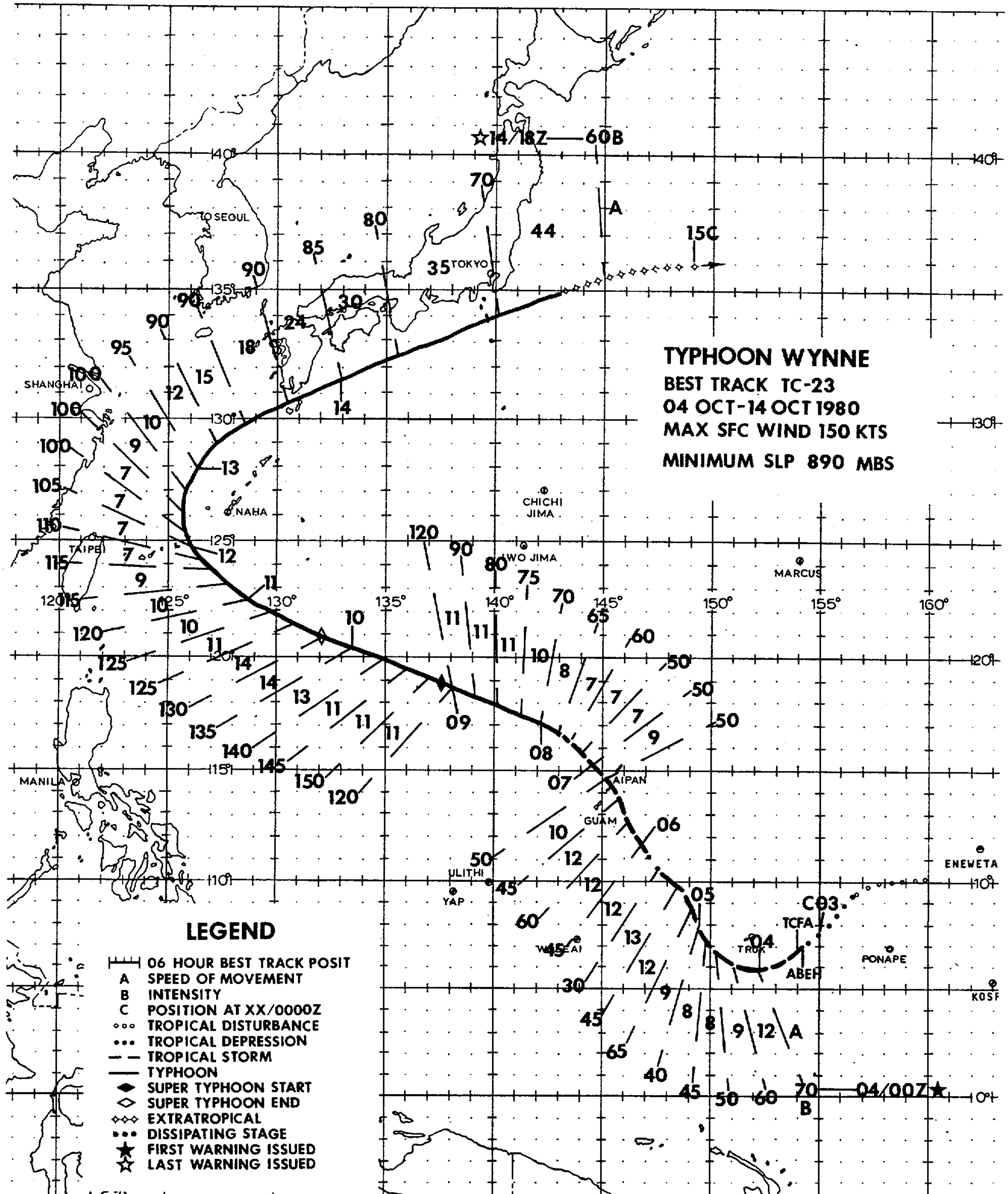


FIGURE 3-22-3. *Typhoon Vernon, with a 40 nm (74 km) ragged eye, at maximum intensity of 105 kt (54 m/sec), 01 October 1980, 2216Z. (NOAA6 imagery)*


TYPHOON WYNNE
BEST TRACK TC-23
04 OCT-14 OCT 1980
MAX SFC WIND 150 KTS
MINIMUM SLP 890 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
14/18Z
04/00Z
SEOUL
TOKYO
SHANGHAI
NAHA
TAIPEI
CHICHI JIMA
IWO JIMA
MARCUS
MANILA
SAIPAN
GUAM
ULITHI
YAP
TRUK
PONAPE
ENEWETAK
TCFA

The disturbance that eventually developed into Super Typhoon Wynne was evident on satellite imagery as early as 1800Z on 30 September, although at that time, it appeared to be simply enhanced convection embedded in the convergent inflow into Typhoon Vernon located 1000 nm (1852 km) to the northwest. By 020000Z October, however, the disturbance had separated from the inflow into Vernon, and by 021200Z, the convective activity had increased in organization with good curvature and upper-level outflow evident from satellite data.

The small scale of the disturbance and the tightness of the circulation that characterized Wynne during most of her life prevented the circulation from appearing on synoptic analyses and led to an underestimation of severity during her formative stage. These facts heavily influenced the decision to delay the issuance of a Tropical Cyclone Formation Alert for 21 hours, although post-analysis indicates that tropical storm strength was achieved as early as 030600Z.

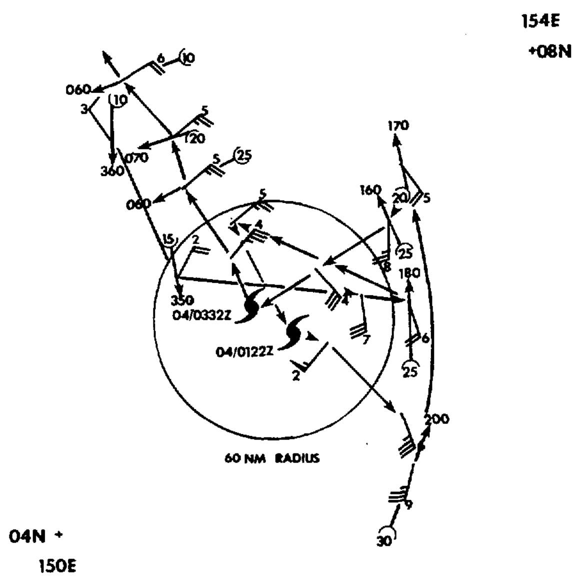


FIGURE 3-23-1. Plot of aircraft reconnaissance data for the 040122Z and 040332Z October 1980 fixes of Tropical Storm Wynne.

Because of her proximity to Guam, numerous aircraft reconnaissance missions were flown into the developing tropical cyclone. This extensive coverage confirmed Wynne's small circulation (Fig. 3-23-1 and 3-23-2) and permitted JTWC to monitor her development very closely.

Although Typhoon Lex may have been the most interesting cyclone of the year in terms of movement, Super Typhoon Wynne proved to be the most unusual in terms of intensity oscillations. As shown in Figure 3-23-3, Wynne's early stage of development was characterized by short periods of rapid intensification and weakening, rather than by a typical smooth, gradual intensification. From 3 October to 7 October, Wynne's intensity and convective activity fluctuated significantly, as she attained typhoon or near typhoon strength only to weaken to near tropical depression intensity three times following a diurnal cycle. Although not as marked as the oscillations in the observed maximum winds, the minimum sea level pressure also exhibited a cyclical oscillation that closely approximated the periodicity of the maximum winds.

There have been documented cases of tropical cyclones exhibiting intensity variations (Holliday, 1976). However, these occurrences were limited to well-developed typhoons with minimum sea level pressures below 970 mb and with a single weakening-reintensifying cycle.

An examination of the satellite imagery during this period of large short-term changes in intensity (Figure 3-23-4) reveals that maximum activity in deep convection occurred in the early morning hours (0700 to 0800 local time) with a minimum in the evening hours (1900 to 2000 local time). An increase in cirrus toward the late afternoon (1600 local time) was also evident. These observations agree with the findings of Arnold (1977). Although Arnold found no evidence of intensity change accompanying the change in cirrus or deep convection, significant intensity change was observed in the case of Wynne with a lag of 6 to 8 hours between maximum convective activity and maximum observed winds.

Wynne's third and final period of weakening occurred as she tracked 45 nm (83 km) northeast of Guam. This weakening, combined with her small circulation, resulted in Wynne having virtually no effect on Guam. Wynne continued to intensify rapidly following her third reintensification cycle at 071800Z, attaining super typhoon strength just 30 hours later and a peak intensity of 150 kt (77 m/sec) in another 6 hours. Figure 3-23-5 depicts Wynne near maximum intensity about 490 nm (908 km) southeast of Okinawa. Minimum sea level pressure (MSLP) during this 35-hour period dropped from 982 mb to 890 mb - a 31 mb/12 hr fall.

JTWC's Theta E (θ_e)/MSLP study once again accurately predicted this explosive deepening as the θ_e and MSLP trace intersected at 081400Z. Wynne's intensity peaked 16 hours after the time of intersection with the surface winds increasing by 85 kt (44 m/sec) and the MSLP falling another 62 mb.

As Wynne tracked north-northwestward past Guam, she was expected to move through

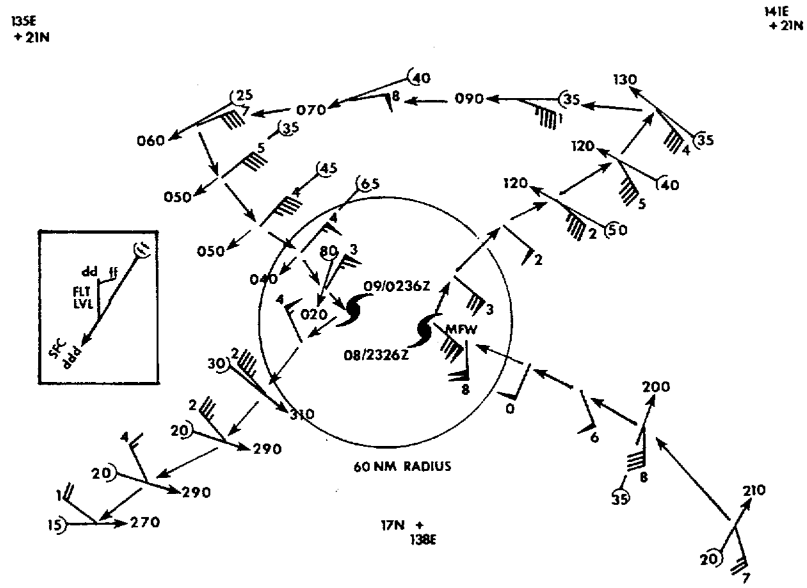


FIGURE 3-23-2. Plot of aircraft reconnaissance data for the 082326Z and 090236Z October 1980 fixes of Typhoon Wynne.

an apparent weakness in the subtropical ridge north of Guam. However, the weakness between 25N and 30N was evidently too far north to permit her to break through the ridge, and she eventually came under the influence of the strong anticyclone located between Okinawa and the Bonin Islands. Post-analysis of available 500-mb data indicates that a change to a more westward forecast track around the southern periphery of the anticyclone could have been made 24 hr earlier. Once the forecast track was changed to reflect the shift in the synoptic flow pattern to a more definitive easterly steering current, JTWC was consistent in accurately predicting recurvature just west of Okinawa. Wynne actually recurved 100 nm (185 km) west of Okinawa, and her slow 7 kt (4 m/sec) bend around the island brought over two days of torrential rain and winds gusting to more than 65 kt (33 m/sec). Very few injuries were reported with farm crops receiving the major wind damage. A small island 30 nm (56 km) northwest of Okinawa and closer to Wynne's path, however, reported winds of 100 kt (51 m/sec) and severe damage.

Once north of the ridge axis, Wynne tracked virtually straight east-northeastward on a heading of 070 degrees. This course kept her approximately 80 nm (148 km) from the coast of Japan. Thirty (15 m/sec) to forty-five kt (23 m/sec) winds were reported by Japanese coastal stations as Wynne accelerated northeastward. Heavy rains claimed several lives and flooded over a thousand homes.

As Wynne accelerated past Japan at speeds exceeding 40 kt (74 km/hr), the vertical wind shear and the influx of cooler, drier air resulted in rapid extratropical transition. A reconnaissance aircraft at 141500Z was unable to find a circulation at 700 mb and satellite imagery at 141800Z revealed no active convection. The remnants of Wynne eventually were absorbed by a developing low pressure system east of Japan.

FIGURES 3-23-3 and 3-23-4 are on the following pages

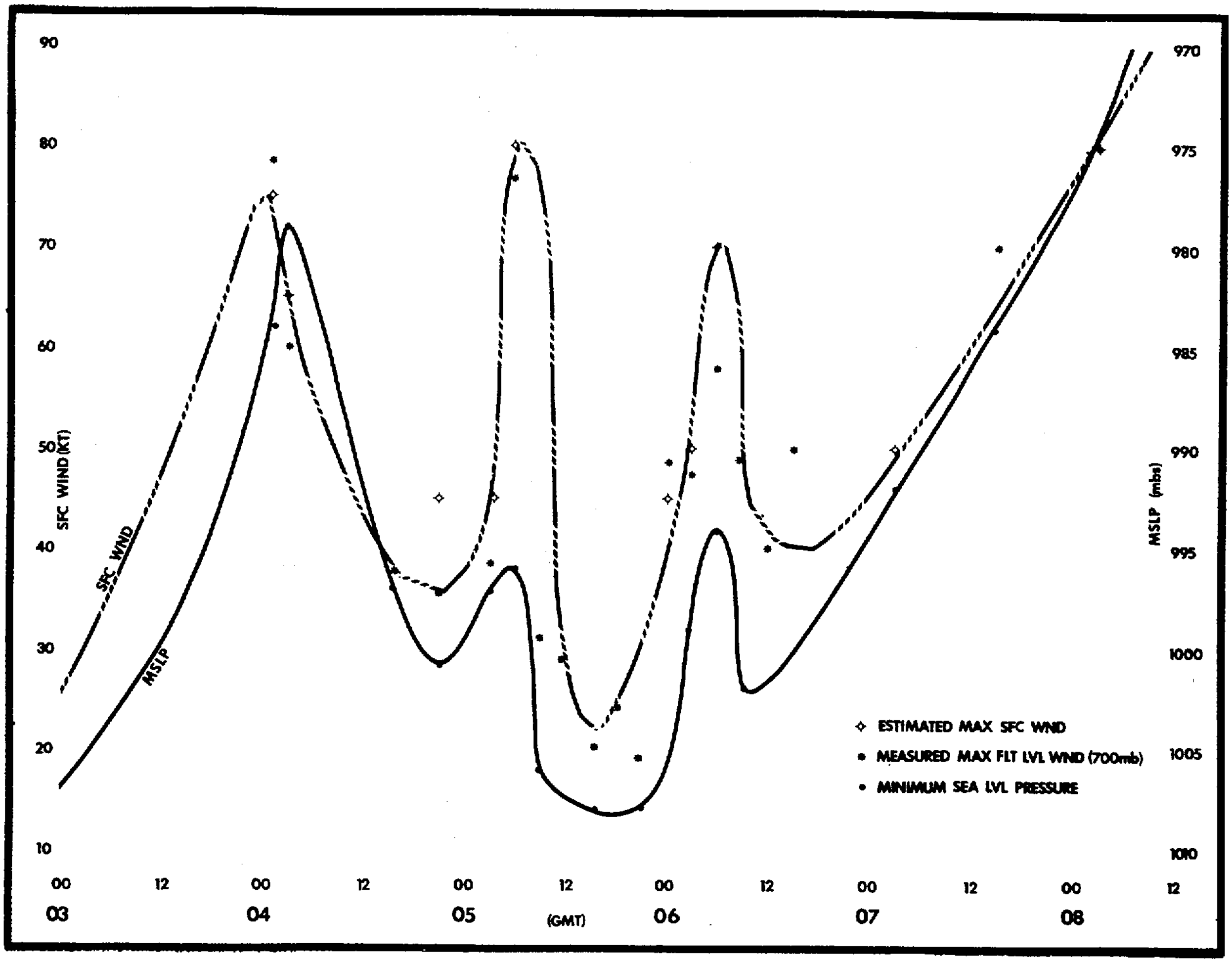


FIGURE 3-23-3. *Smoothed traces of Wynne's maximum wind speed and minimum sea level pressure versus time for the period 3-9 October 1980.*

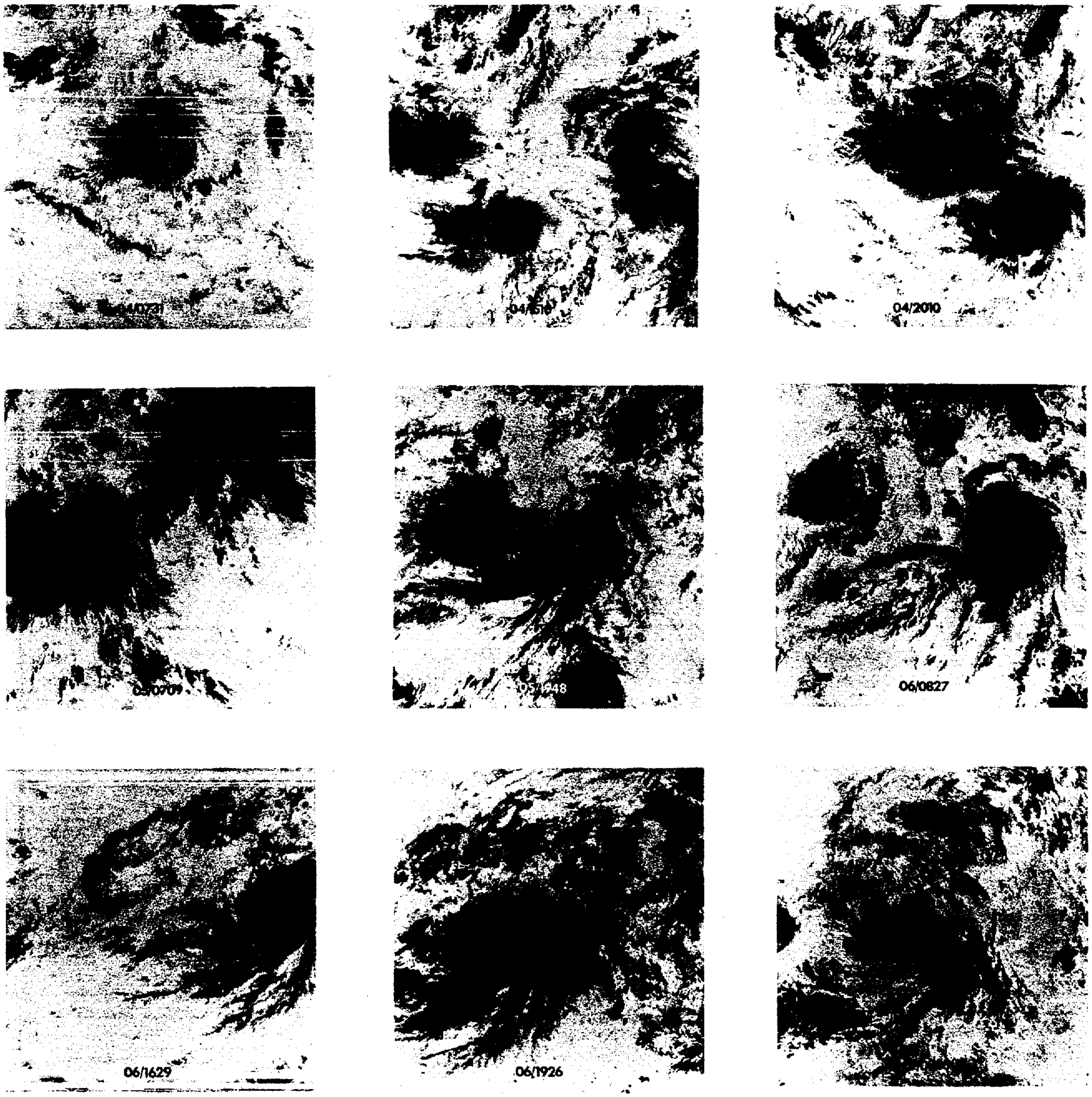


FIGURE 3-23-4. Series of infrared imageries of Wynne during the period of intensity oscillations. The sequence shows definite weakening of the deep convection during the late evening (particularly the 042011 local and 061916 local imageries), followed by a noticeable increase in the convection on the morning satellite imagery. All times local. (NOAA6 and TIROS N imagery)

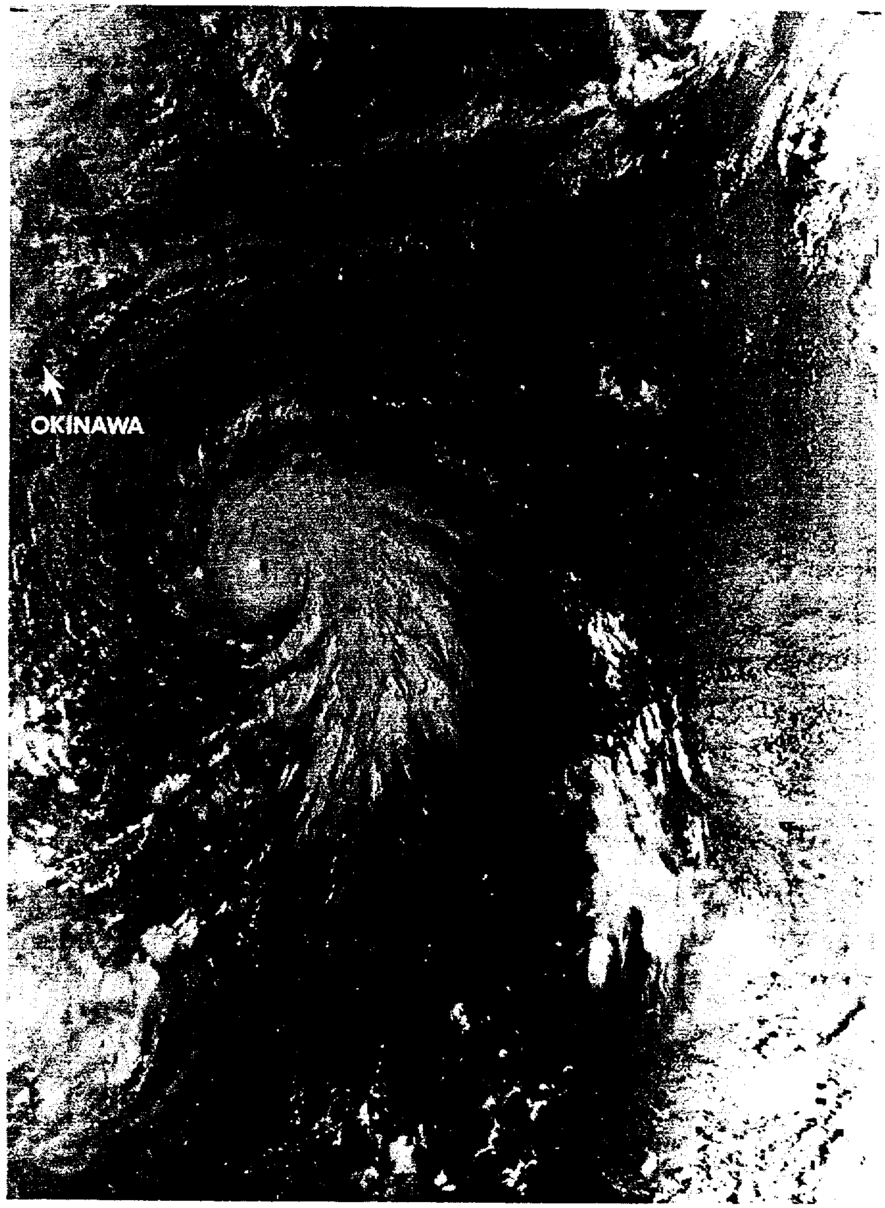
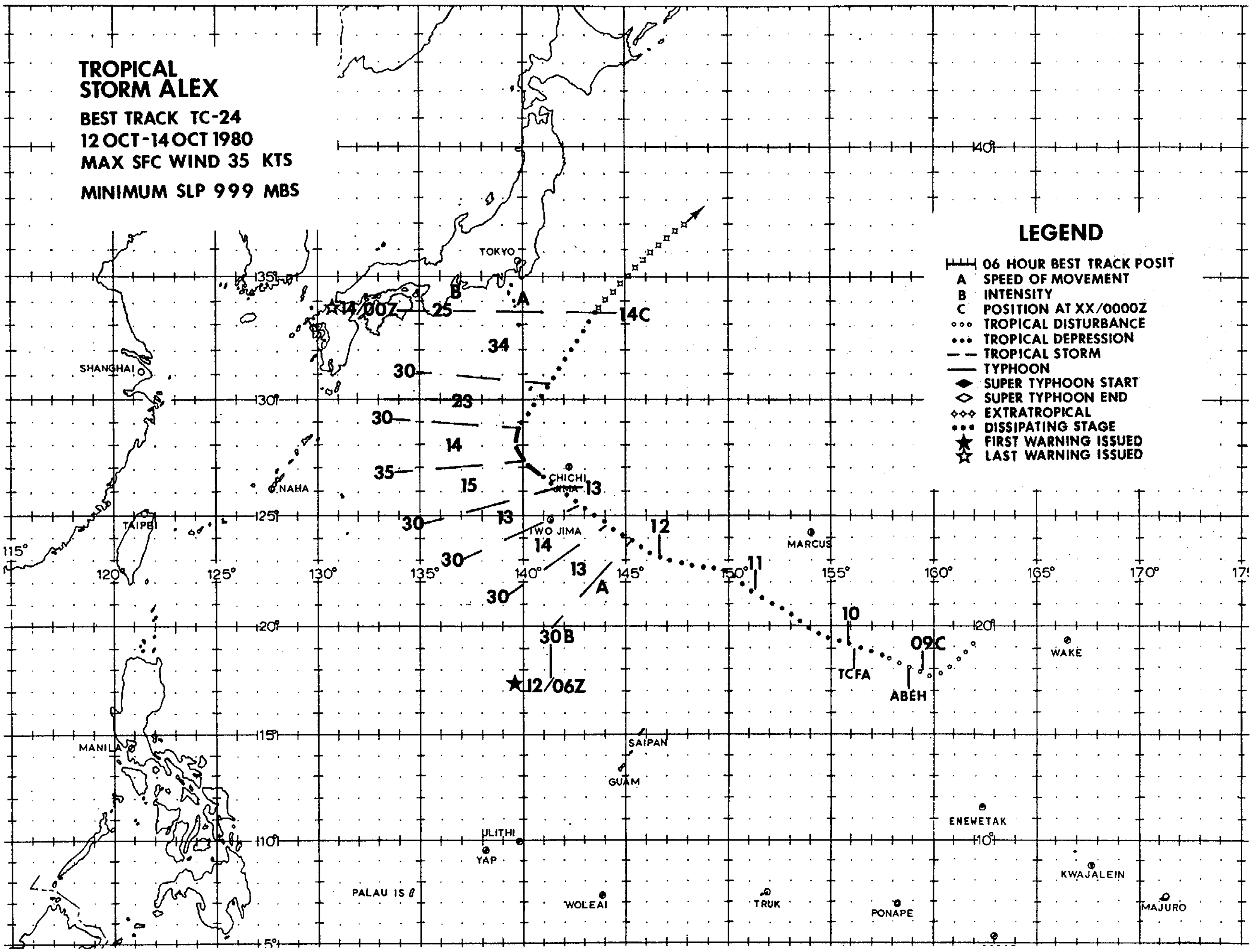


FIGURE 3-23-5. *Super Typhoon Wynne near maximum intensity 490 nm (907 km) southeast of Okinawa and 730 nm (1352 km) northwest of Guam, 9 October 1980, 2240Z. (NOAA6 imagery)*



TROPICAL STORM ALEX
BEST TRACK TC-24
12 OCT-14 OCT 1980
MAX SFC WIND 35 KTS
MINIMUM SLP 999 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
14/00Z
14C
12/06Z
09C
TCFA
ABEH
TOKYO
SHANGHAI
NAHA
TAIPEI
CHICHI JIMA
IWO JIMA
MARCUS
WAKE
MANILA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
ENEWETAK
KWAJALEIN
MAJURO

Tropical Storm Alex, the 24th tropical cyclone of 1980, was induced by a Tropical Upper Tropospheric Trough (TUTT) in a manner similar to that described by Sadler (1976). A small disturbed area of convection drifting westward from near 170E was observed on satellite imagery on 7 October. By the 8th, this area had come under the influence of a relatively strong upper-level divergent area generated by a dissipating TUTT cell. The convection increased noticeably with outflow to the north, but little outflow was evident in the southern and western quadrants.

The restricted outflow pattern was characteristic throughout Alex's existence and is attributed to the proximity of Super Typhoon Wynne, which was located west of Alex. Alex's coexistence with Wynne was significant in light of Wynne's overall dominance of the western Pacific region. Wynne absorbed much of the energy that otherwise would have been available to Alex (Fig. 3-24-1).

Satellite imagery showed that the convective area continued to persist until late on the 9th when JTWC issued a Tropical Cyclone Formation Alert (TCFA). Aircraft reconnaissance on the 10th found a weak surface circulation with the associated convection located north and east of the surface center. Nearly the entire western half of the circulation was exposed at that time.

For the following 48 hours, the disturbance intensified gradually and tracked northwestward at 12 kt (22 km/hr). At 120600Z, the first warning was issued for TD 24. Within 24 hours, TD 24 intensified to Tropical Storm Alex with maximum surface winds of 35 kt (18 m/sec). At that time, Alex's low-level circulation center was not exposed to the west, but aircraft reconnaissance encountered only weak convective activity around the circulation's center. During the next 6 hours, Alex recurved to the northeast and weakened to 30 kt (15 m/sec) intensity.

After Alex had recurved, Alex and Wynne were within 800 nm (1482 km) of each other such that a Fujiwhara effect was possible. This was not observed, however, because Alex and Wynne were both beginning to interact with the jet stream which became the dominant steering mechanism over both cyclones. Due to this jet stream, Alex rapidly accelerated northeastward. At 140000Z, JTWC issued the final warning on Alex as he was beginning to transition into an extratropical system. Satellite imagery received after the final warning showed that the transition was very rapid.

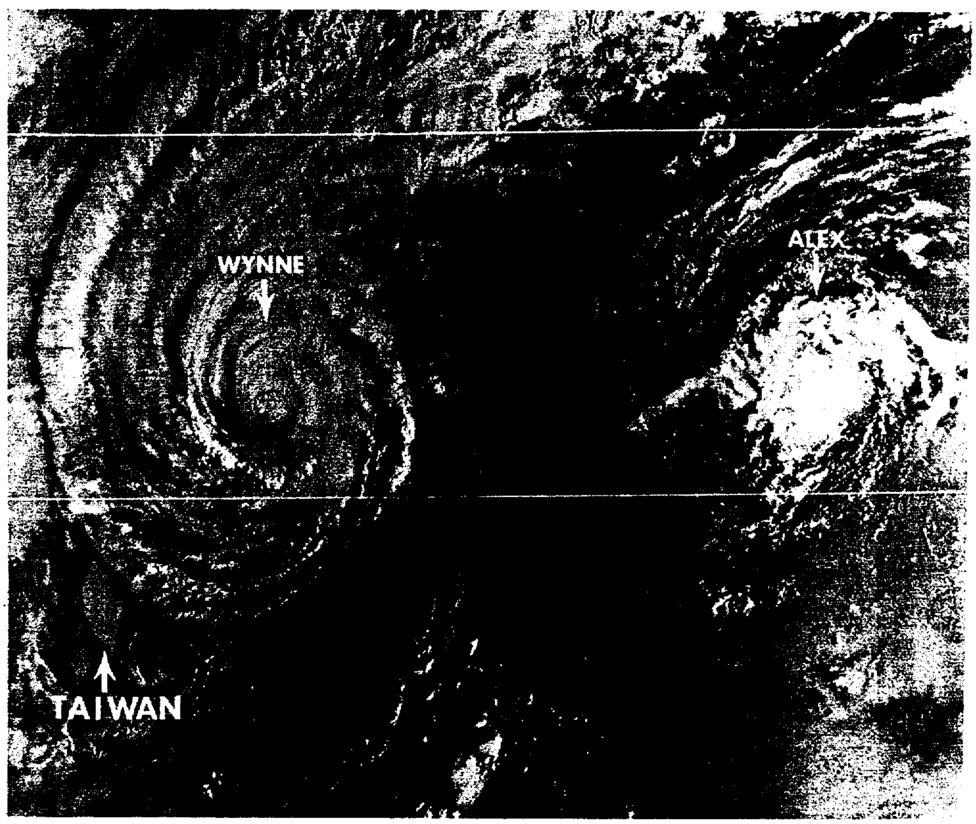
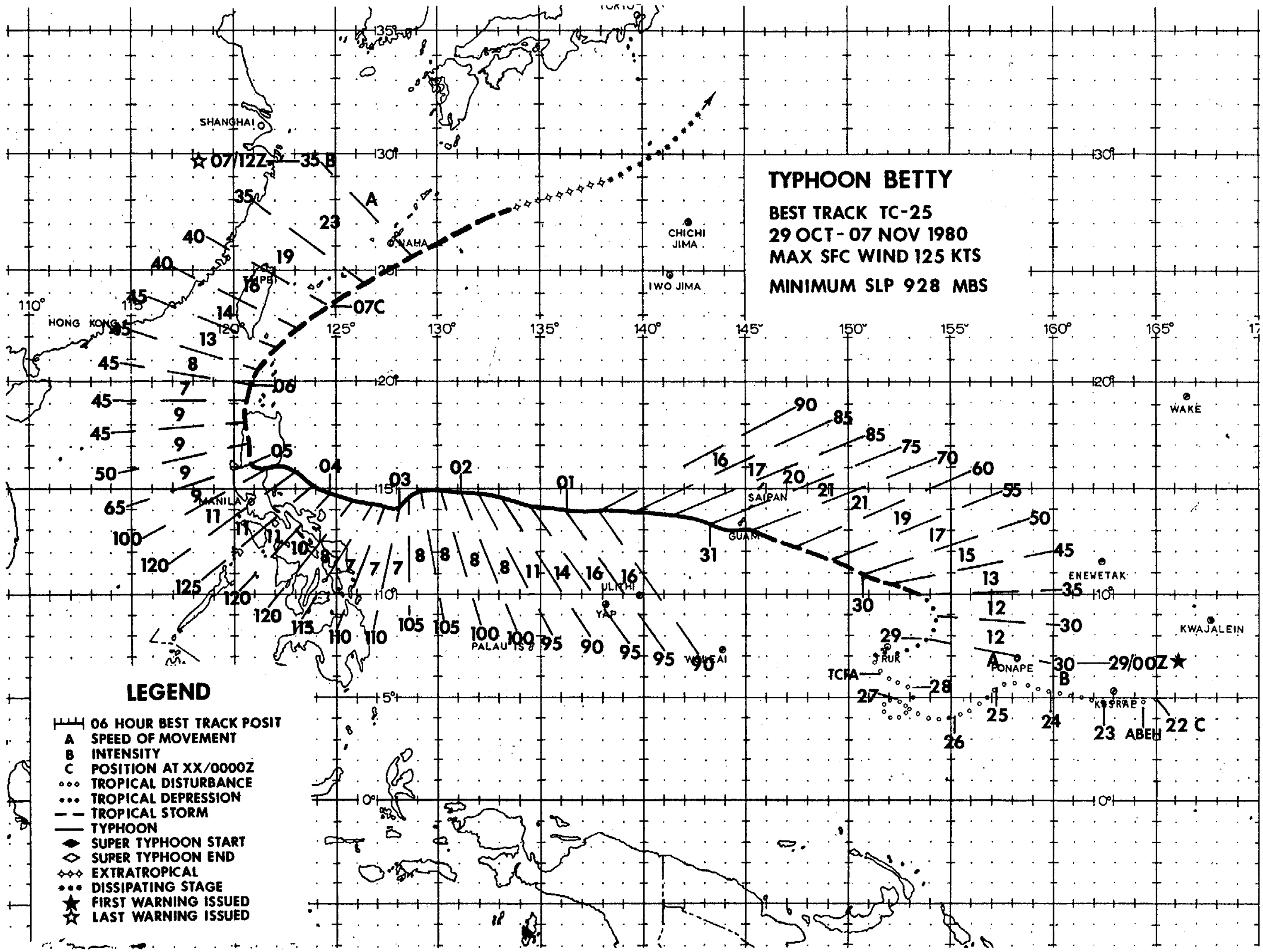


FIGURE 3-24-1. Visual imagery of Tropical Storm Alex at 30 kt (15 m/sec) intensity, 12 October 1980, 2314Z. (NOAA6 imagery)



TYPHOON BETTY
BEST TRACK TC-25
29 OCT - 07 NOV 1980
MAX SFC WIND 125 KTS
MINIMUM SLP 928 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
HONG KONG
TAIPEI
NAHA
MANILA
CHICHI JIMA
IWO JIMA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
PONAPE
KOSRAE
ENEWETAK
KWAJALEIN
WAKE
TCFA
07/12Z
07C
29/00Z
22 C
ABEH

Betty, the 25th significant tropical cyclone of 1980, developed east of Truk Atoll from a weak disturbance which had been monitored for almost a week in the eastern Caroline Islands. Just prior to passing south of Guam, Betty attained typhoon strength and then continued to intensify as she tracked into the Philippine Sea. About 12 hours prior to landfall on Luzon, Betty reached her peak intensity of 125 kt (64 m/sec). During the 18 hours that Betty tracked over north central Luzon, she weakened considerably, but in the process caused extensive damage and loss of life. Downgraded to tropical storm strength, Betty moved northeastward through the Bashi Channel and eventually dissipated as a weak extratropical low southeast of Japan.

Betty had her origin in a weak disturbance south of Truk which showed increased potential for development on the 27th and 28th of October. The 280000Z gradient level winds from Truk and Ponape, as well as low-level winds from a weather reconnaissance flight west of the disturbance, indicated a closed circulation with 20 to 25 kt (10 to 13 m/sec) winds. At 280800Z, a Tropical Cyclone Formation Alert (TCFA) was issued and, during the alert period, the disturbance veered sharply to the east as it approached Truk on an erratic course from the south. After veering, hourly reports from Truk and satellite imagery indicated increased organization, and the first warning for TD 25 was issued at 290000Z with maximum surface winds of 30 kt (15 m/sec).

Despite the erratic movement shown during its formative stages and the apparent northeastward trajectory TD 25 had assumed by the first warning, the initial and subsequent warnings correctly identified a west-northwest track which indicated passage just south of Guam. However, due to limited mid-level (700 mb or 500 mb) steering data north of Betty, the first six warnings failed to adequately forecast her acceleration which resulted in a speed of movement of 21 kt (39 km/hr) as she passed Guam. As a result, although the 72 hour forecast position of the second warning predicted Betty's exact position as she passed south of Guam, the average vector error during this period was very high. The 72 hour forecast position mentioned above, had a total error of 585 nm (1083 km) due to the acceleration which caused Betty to reach the 72 hour point in just 34 hours! Such errors resulting from under forecasting speed of movement highlight the importance of adequate mid-level data in the steering current. When available, especially from reconnaissance aircraft, such data usually increase the ability of forecasters to evaluate the potential for changes in the short-term, as well as long-term motion of tropical cyclones.

After passing south of Guam, Typhoon Betty turned west and continued to intensify, reaching 100 kt (61 m/sec) 48 hours later. During this period, the 500 mb analyses began to show a short wave trough moving east through mainland China. JTWC forecasts keyed on this feature and, based on computer-derived prognostic charts, recurvature was expected to begin near 125E by 030000Z. The probability of this forecast verifying increased when, at 020000Z, the short wave trough deepened as it moved off Asia. However, by 021200Z, Betty unexpectedly turned southwestward. By 030000Z, the trough had moved quickly eastward north of Betty and the opportunity for recurvature had passed. Shortly afterwards, attention focused on another short wave moving through China and recurvature was again forecast to occur, this time just east of Luzon. By 040000Z, however, available 500 mb data did not show any significant amplitude to this trough and the recurvature track was abandoned in favor of a northwestward track over Luzon into southern China.

Although Betty continued to intensify after passing Guam, the data normally used to evaluate a tropical cyclone's intensity showed considerable scatter. Figure 3-25-1 graphically depicts these data as well as the final best track intensities. In searching for an explanation of the scatter, the comments from mission ARWOs on 01 November and 04 November may offer some insight concerning the character of Betty during this period. On the 011594Z mission: "Although this storm (Betty) was strong, it had no eyewall. The most fascinating feature was the rotating feeder band of convection that was spiralling inward at an enormous rate."[1] On this single fix mission, a maximum 700 mb flight level wind of 95 kt (49 m/sec) was observed. However, daylight missions before and after this mission estimated surface winds in excess of 100 kt (51 m/sec) (Fig. 3-25-2). On the 032200Z, 040150Z and 040340Z fix missions, it was observed that "Typhoon Betty.....was a textbook typhoon. Everything was aligned perfectly."[2] At 040600Z, Betty reached her peak intensity of 125 kt (64 m/sec). The development of a textbook typhoon correlated closely with the reduction in scatter between maximum surface wind estimates shown in Figure 3-25-1.

[1]Candis L. Weatherford, 1 LT, USAF, Mission ARWO.

[2]James B. Near, 1 LT, USAF, Mission ARWO.

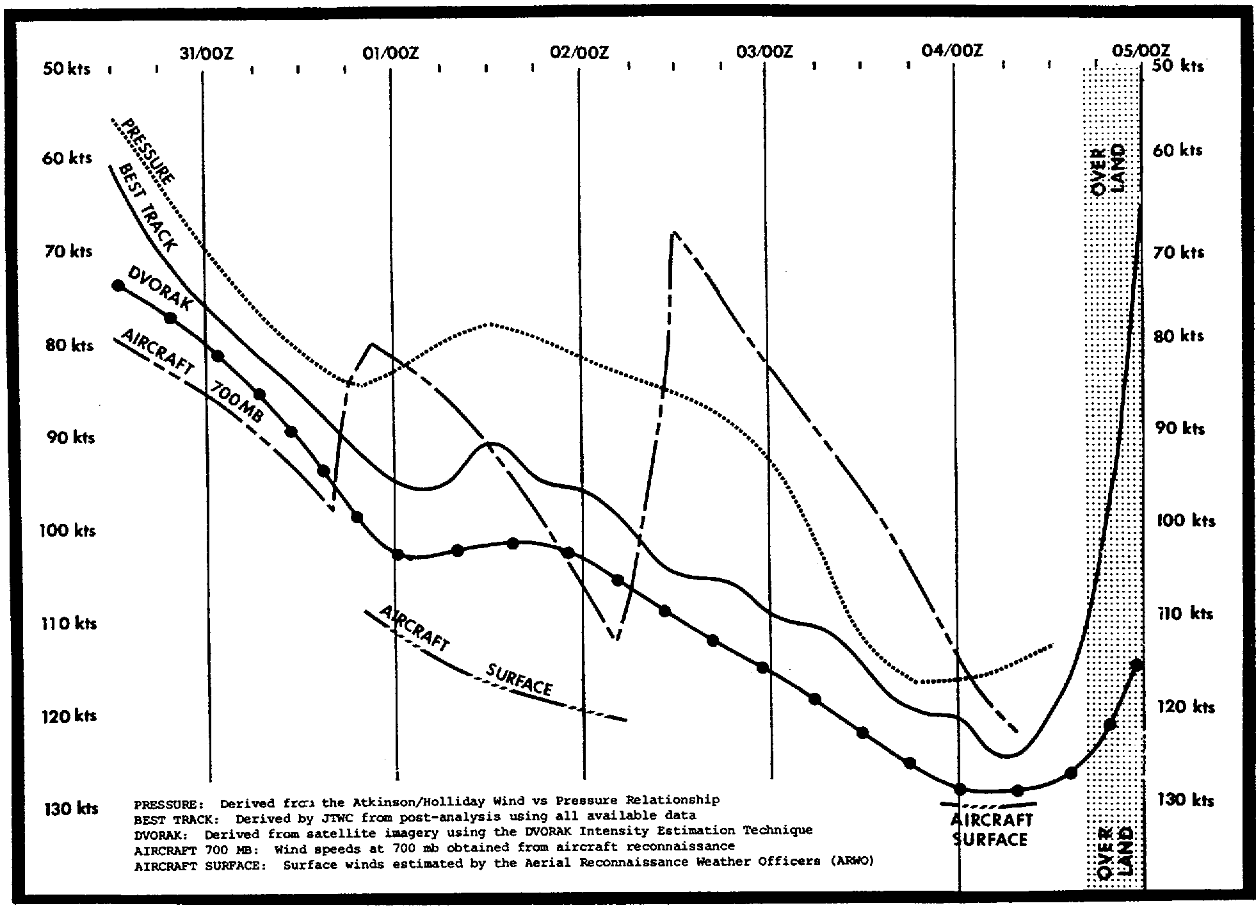
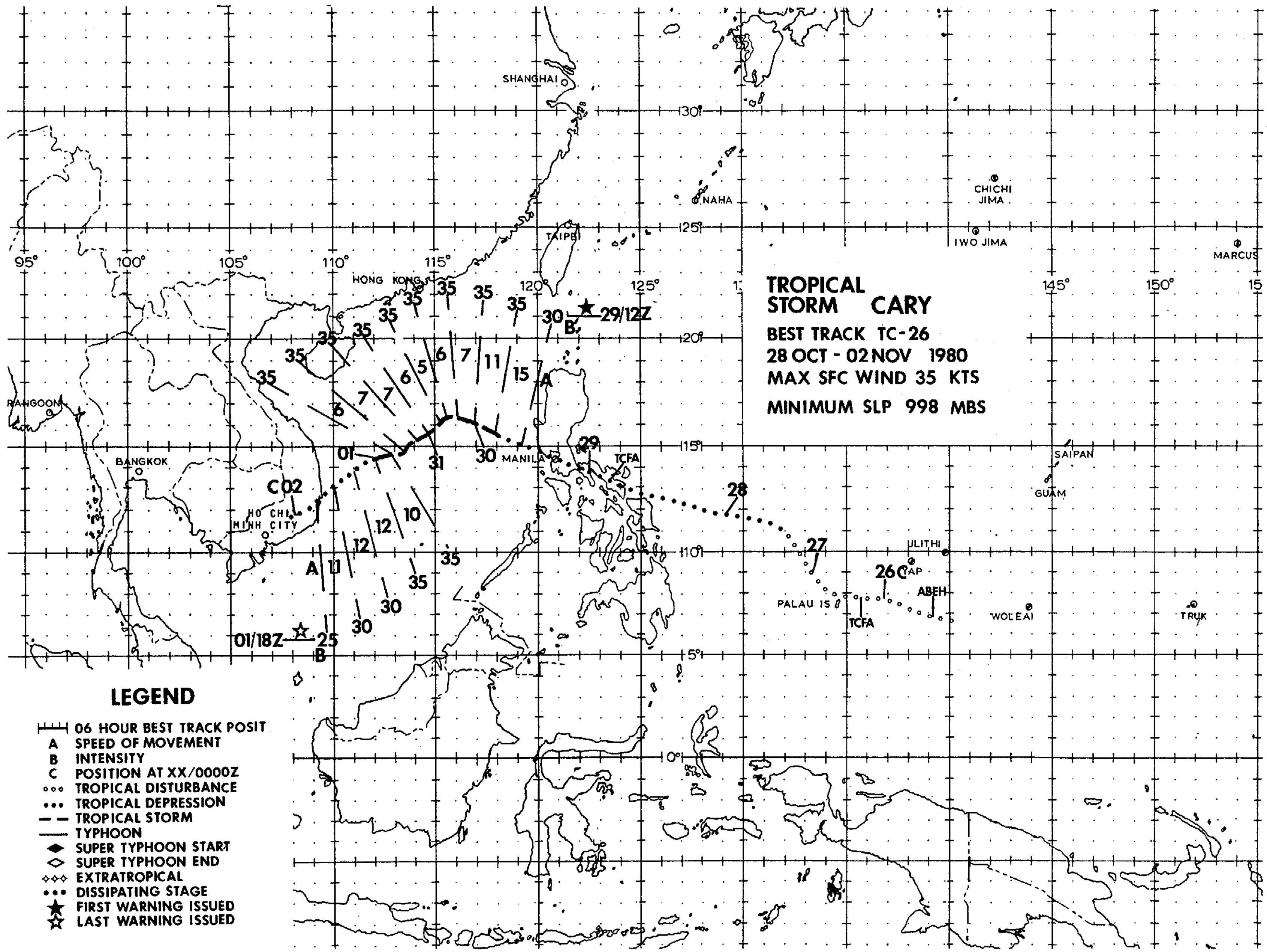


FIGURE 3-25-1. Time series of various intensity parameters evaluated by JTWC while Betty was at typhoon strength. Note the large scatter in the traces until 040600Z.

The decision to abandon a forecast recurve track east of Luzon put the central and northern provinces of Luzon on alert. At 041600Z, Typhon Betty, packing 120 kt (62 m/sec) winds, slammed into central Luzon south of Cape San Ildefonso. Most weather observing stations stopped reporting prior to Betty's approach, so her actual intensity as she crossed Luzon can only be inferred from a JTWC study of prior tropical cyclones crossing the Philippines (Sikora, 1976), satellite imagery (Dvorak intensity estimates), and aircraft reconnaissance reports just prior to and after Betty crossed Luzon. However intense Betty may have been, there is little doubt that she was one of the most destructive typhoons of recent history. Initial reports received several days after Betty crossed Luzon indicated at least 81 people dead, thousands homeless, and extensive crop damage from flooding and mudslides. The Cagayan Valley in northern Luzon, hard hit by Betty, lost most of its rice crop from floodwaters which rose to roof-top level in some areas. Philippine Defense Minister, Juan Ponce Enrile, described the Cagayan Valley from a helicopter survey, stating, "It looks like a sea from the air."

As Betty weakened over Luzon, the ridge that influenced her track into the Philippines broke down, thus, allowing her to drift northward along the Cordillera Central Mountains and eventually to be drawn into the weak short wave trough which had stalled off the coast of Taiwan. Emerging from Luzon as a 45 kt (23 m/sec) tropical storm, Betty never regained her earlier fury as she moved east of Taiwan and the Ryukyu Islands before undergoing an extratropical transition just prior to 080000Z. As an extratropical system, the remnants of Betty did not persist long. This once powerful typhoon was last observed 12 hours later dissipating southeast of Honshu, near 32N 143E.

FIGURE 3-25-2. During this stage of Typhoon Betty's development, intense banding was reported by aircraft reconnaissance along with areas of surface winds in excess of 100 kt (51 m/sec), 31 October 1980, 2252Z. (NOAA6 imagery)


TROPICAL
STORM CARY
BEST TRACK TC-26
28 OCT - 02 NOV 1980
MAX SFC WIND 35 KTS
MINIMUM SLP 998 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SHANGHAI
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
TAIPEI
HONG KONG
RANGOON
BANGKOK
HO CHI MINH CITY
MANILA
SAIPAN
GUAM
ULITHI
YAP
PALAU IS
WOLEAI
TRUK
TCFA
29/12Z
01/18Z
95°
100°
105°
110°
115°
120°
125°
145°
150°
30°
25°
20°
15°
10°
5°
0°

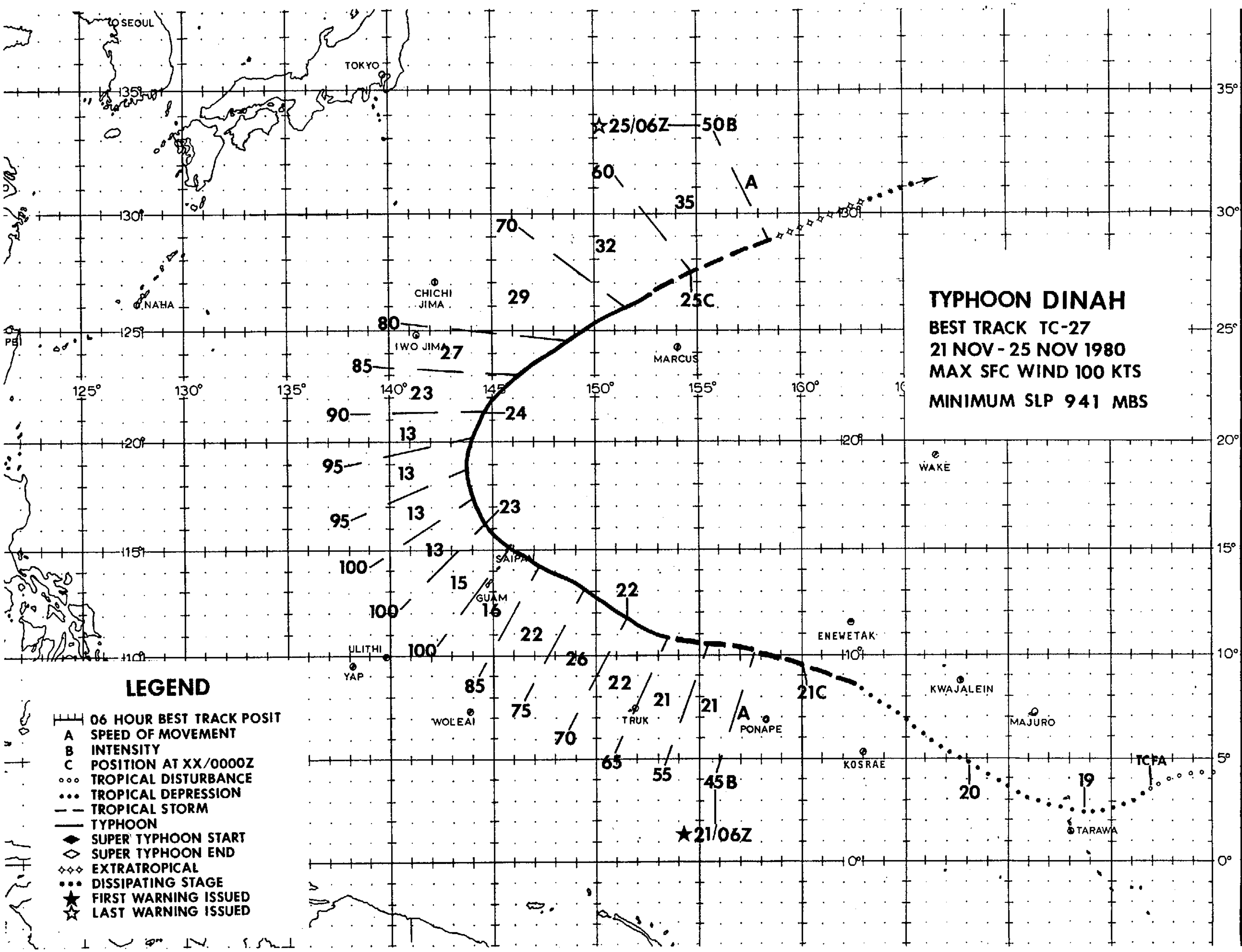
Tropical Storm Cary was first observed as a area of increased convective activity east of the Palau Islands on the 25th of October. A Tropical Cyclone Formation Alert (TCFA) was issued a day later when the disturbance had moved to a position about halfway between Yap and the Palau Islands.

All convection associated with this circulation dissipated shortly afterward, however, leaving only an exposed low-level circulation center. Over the next 24 hours, the system moved west-northwestward toward the Philippines. Just east of the Phillippines, the convection again developed and 30 kt (15 m/sec) winds were reported from coastal stations. Based on this information, a second TCFA was issued at 282200Z.

The circulation maintained its indentity as it passed over the Philippines just south of Clark AB and Subic Bay Naval Station. The first warning on TD 26 was issued at 291200Z as the disturbance was moving into the South China Sea. Tropical storm strength was reached 6 hours later.

Tropical Storm Cary moved west-northwestward and then west-southwestward in response to a low-level northeast monsoonal surge (Fig. 3-26-1) and eventually dissipated over Vietnam on 02 November.

FIGURE 3-26-1. Tropical Storm Cary near maximum intensity in the South China Sea. The surface center is partially exposed as indicated by the cumulus banding southeast of the main convection, 01 November 1980, 0033Z. (NOAA6 imagery)


TYPHOON DINAH
BEST TRACK TC-27
21 NOV - 25 NOV 1980
MAX SFC WIND 100 KTS
MINIMUM SLP 941 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0000Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
25/06Z
50B
21/06Z
45B
25C
21C
TCFA
SEOUL
TOKYO
NAHA
CHICHI JIMA
IWO JIMA
MARCUS
WAKE
SAIPAN
GUAM
ULITHI
YAP
WOLEAI
TRUK
PONAPE
ENEWETAK
KWAJALEIN
MAJURO
KOSRAE
TARAWA

Dinah, the final typhoon of the 1980 season and the third tropical cyclone this season to threaten Guam, began to develop in mid-November as a focal point of cumulus banding embedded in the monsoon trough oriented east-west near Kwajalein. Initial development of this system was slow and erratic, as four successive Tropical Cyclone Formation Alerts (TCFAs) were issued for this area between 18 and 20 November. On the 21st however, this system finally established a well-developed outflow pattern, and its heaviest associated convection, which was initially more evident along the periphery of the circulation, began to consolidate about the system's center. The first warning on Tropical Storm Dinah was issued at 210600Z. At that time, having established a well-developed outflow to all quadrants, Dinah intensified rapidly and subsequently reached typhoon strength at 211800Z, just 12 hours after the initial warning.

A post-analysis of Dinah's development reveals some unique properties. First, she exhibited a very compact circulation, which she maintained throughout her lifespan as a tropical cyclone. The 30 kt (15 m/sec) wind radius was significantly less than normal. Second, a persistent easterly flow occurred near the surface during Dinah's initial development and may have been a primary factor for her slow and erratic development. For example, the surface analysis at 200000Z (Fig. 3-27-1) indicated an associated surface circulation near 4N 168E and a brisk easterly gradient-level flow north of the surface circulation. This flow pattern resulted in both the abnormally rapid movement of the developing system and an unusually pronounced asymmetry in her wind field which displaced the maximum wind band to the north of the circulation center. A subsequent surface analysis, at 211200Z (Fig. 3-27-2), however, did not indicate a surface circulation, but rather weak easterly flow south of where the

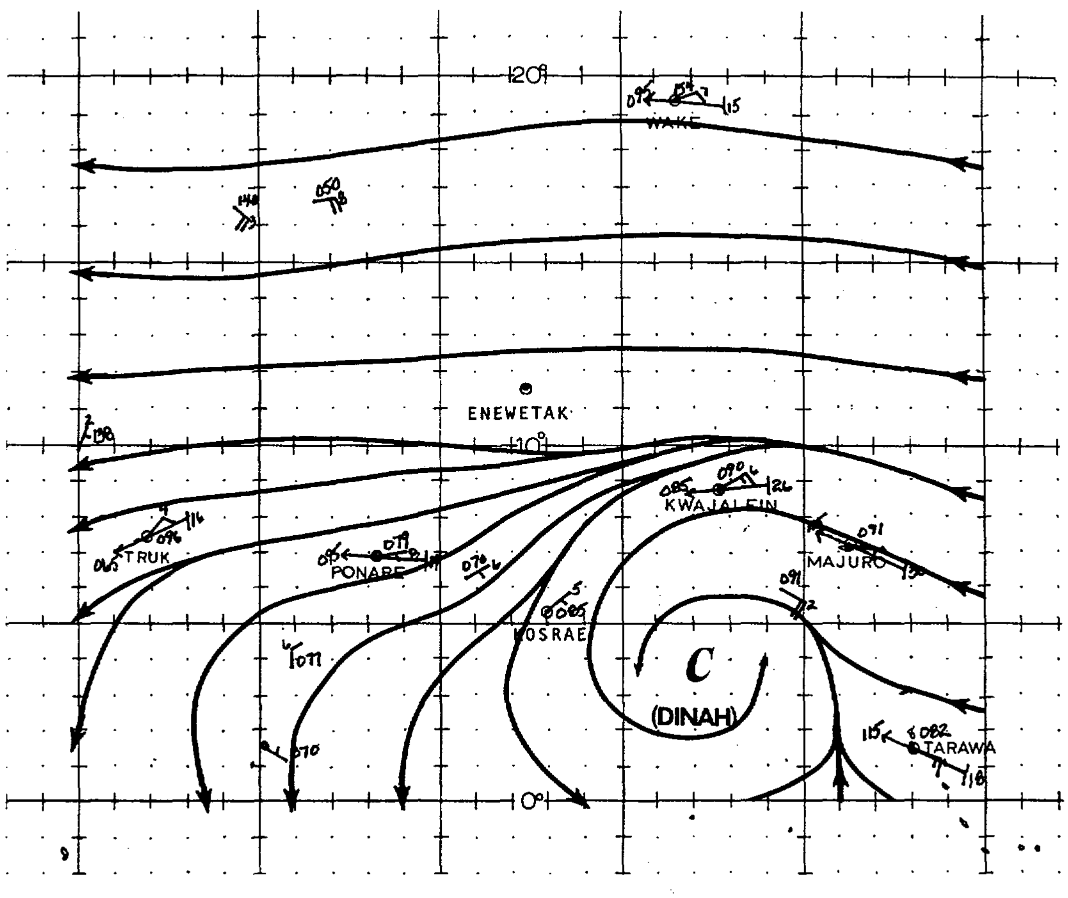


FIGURE 3-27-1. The 200000Z November 1980 surface (——) /gradient-level (ddd←——ff) wind data and streamline analysis. Wind speeds are in knots.

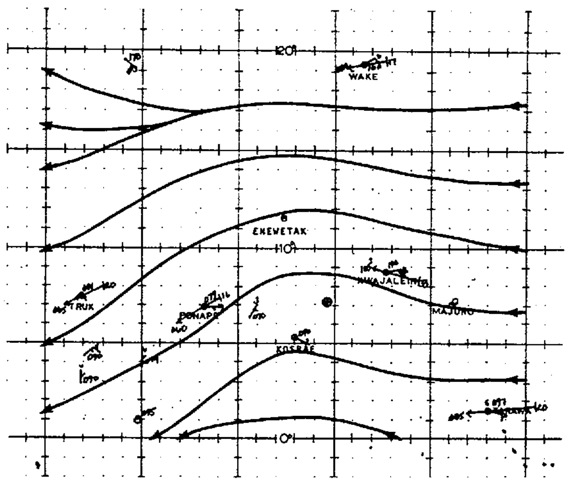


FIGURE 3-27-2. The 201200Z November 1980 surface (———)/gradient-level (ddd←——ff) wind data and streamline analysis. Wind speeds are in knots. ⊗ *indicates satellite position of Dinah at about the same time.*

circulation's position was indicated in satellite imagery. At the same time, cyclonic flow was present over the area at 500 mb and a closed cyclonic center existed just northwest of the disturbance at 200 mb. In view of the above data, it is probable that Dinah developed from a mid- or upper-level cyclone that subsequently generated its own surface circulation. The Aerial Reconnaissance Weather Officer (ARWO)[1] aboard the initial flight into what ultimately became Typhoon Dinah, stated "the storm was compact, with a very sharp pressure gradient and good banding....We had difficulty closing off the circulation to the north and northwest because it may just have actually closed [itself] off".

By the time Dinah intensified to a typhoon, she posed a definite threat to Guam within 48 hr; thus, the decision was made to evacuate military aircraft from the island. A comparison of the 500 mb analysis (which is generally considered the primary steering level for tropical cyclones) just prior to and subsequent to the aircraft evacuation, demonstrates the great importance of enroute aircraft reports of flight-level winds (AIREPS) and the significance they can make to a tropical cyclone forecast. The 500 mb streamline analysis at 211200Z (Fig. 3-28-3) shows a strong anticyclone near Marcus Island and strong ridging west-southwestward toward the Philippine Islands. In response, JTWC forecast Typhoon Dinah to pass just off the northeastern tip of Guam. The next 500

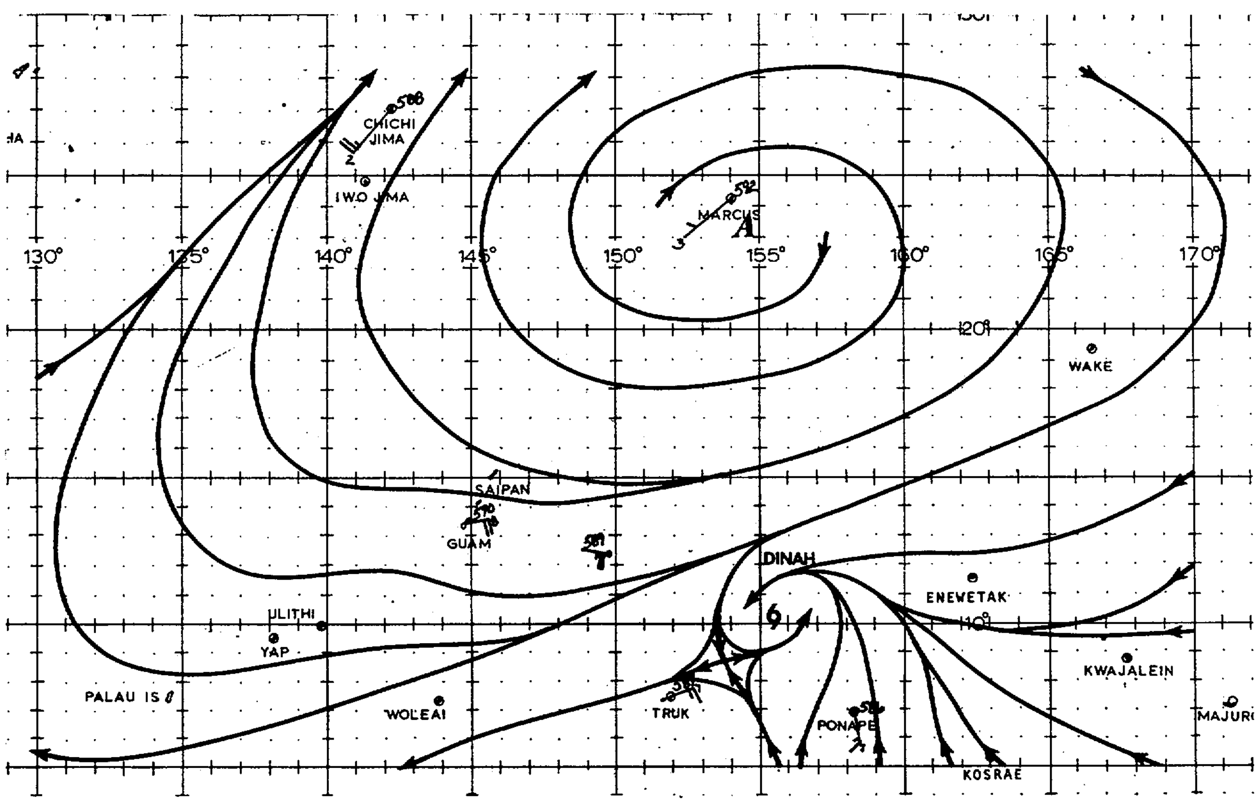


FIGURE 3-27-3. The 211200Z November 1980 500 mb wind data and streamline analysis. Wind speeds are in knots.

[1]Richard F. Ferris, 1 Lt, USAF: Mission ARWO

mb streamline analysis at 220000Z (Fig. 3-28-4), which was augmented by a series of AIREPS taken by an evacuation flight enroute from Guam to Okinawa, enabled JTWC to analyze a weakness in the ridge just north of Guam. In view of this new information, JTWC amended Dinah's forecast track to predict that Dinah would track near Saipan vice Guam. Because Dinah was so compact, this small change in track was enough that Guam received very little wind as Dinah passed to the northeast, but Saipan and nearby Tinian received typhoon-force winds and sustained extensive damage.

Dinah continued to intensify rapidly as she began to move into the weakness north of Guam toward the Northern Marianas Islands. Dinah subsequently crossed the northeastern portion of Saipan at 221845Z and reached maximum intensity at 222100Z, with maximum sustained winds of 100 kt (52 m/sec) and gusts to 130 kt (67 m/sec). After crossing Saipan, Dinah continued to move through the weakness in the ridge near 140E and began to recurve to the north on 23 November. She then weakened and accelerated to the northeast in response to a mid-tropospheric longwave trough which was moving eastward past Marcus Island on the 24th. Dinah transitioned to an extratropical cyclone by 251200Z.

Damage to the islands of Saipan and Tinian was massive, with 60 homes destroyed and another 214 homes suffering damages. Saipan, in the aftermath of Typhoon Dinah, was completely without power for several days and 85 percent of the water system was not functioning. Carlos S. Camacho, Governor of Saipan, estimated damages totalling 7 million dollars. Shortly after damages were assessed, President Carter declared the area a major disaster area, enabling the area to qualify for federal disaster fund relief.

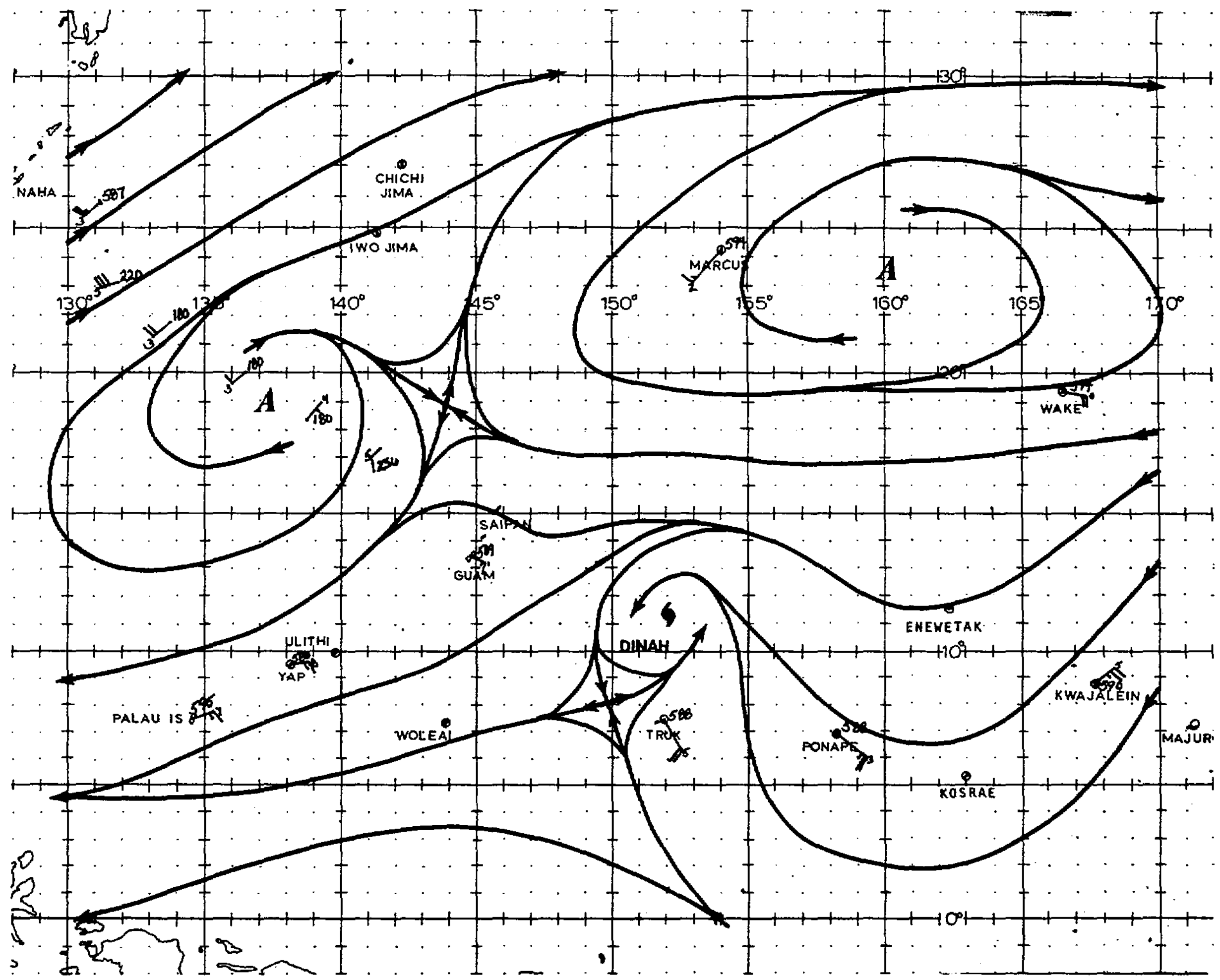
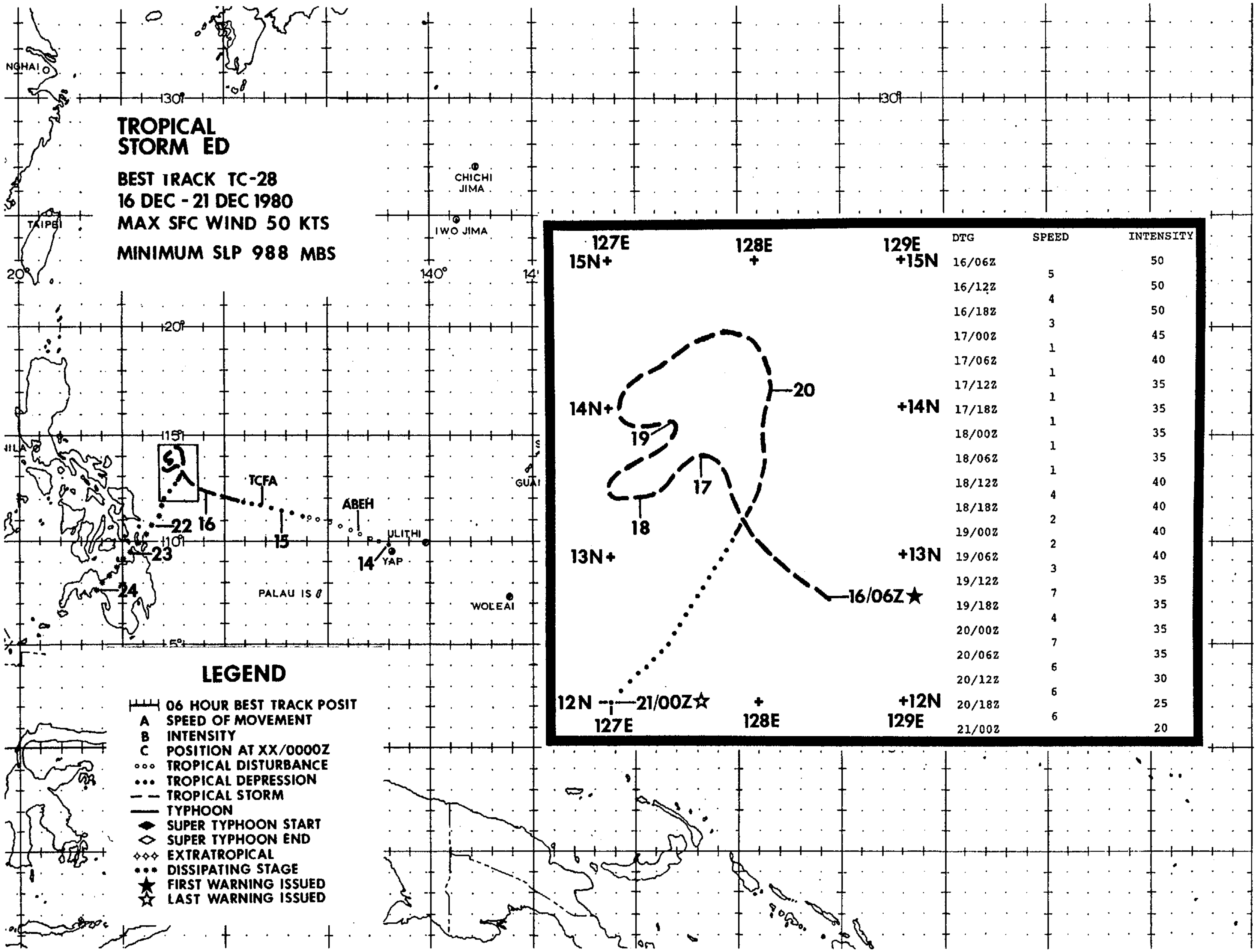


FIGURE 3-27-4. The 220000Z November 1980 500 mb wind data and streamline analysis. Wind speeds are in knots. Note the AIREPS northwest of Guam which were provided to JTWC by a Navy aircraft evacuating from Guam to Okinawa.



DTG	SPEED	INTENSITY
16/06Z		50
	5	
16/12Z		50
	4	
16/18Z		50
	3	
17/00Z		45
	1	
17/06Z		40
	1	
17/12Z		35
	1	
17/18Z		35
	1	
18/00Z		35
	1	
18/06Z		35
	1	
18/12Z		40
	4	
18/18Z		40
	2	
19/00Z		40
	2	
19/06Z		40
	3	
19/12Z		35
	7	
19/18Z		35
	4	
20/00Z		35
	7	
20/06Z		35
	6	
20/12Z		30
	6	
20/18Z		25
	6	
21/00Z		20

Tropical Storm Ed was the last significant tropical cyclone to develop in the western North Pacific in 1980. Ed was never forecast to reach typhoon strength due to the strong vertical wind shear which developed in the vicinity of the Philippine Islands during the last half of December.

Tropical Storm Ed was first observed as a disturbance near Yap on the 14th of December. The disturbance moved westward at between 12 and 15 kt (22 to 28 km/hr) as its convective activity and overall organization continued to improve. A Tropical Cyclone Formation Alert (TCFA) was issued when a reconnaissance aircraft observed a well-defined low-level circulation with a minimum sea-level pressure of 1004 mb. The first warning on Tropical Storm Ed was issued at 160000Z when 50 kt (25 m/sec) surface winds and a 991 mb pressure were reported. Maximum surface winds were consistently observed northeast of Ed in a region of enhanced pressure gradient between the cyclone's center and a strong surface ridge.

It became evident from synoptic analyses that Ed was moving into an area which was unfavorable for continued development. Figures 3-28-1 and 3-28-2 are representative of the

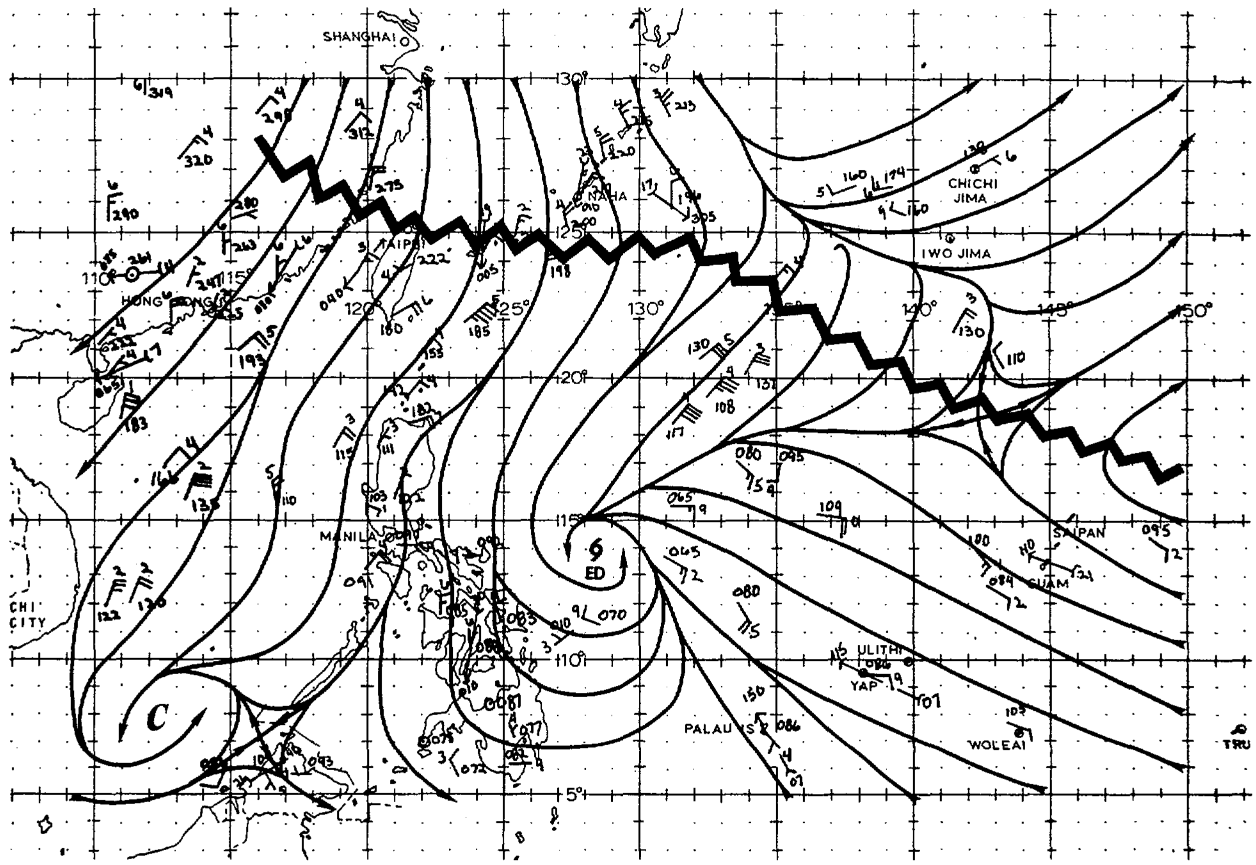


FIGURE 3-28-1. The 200000Z December 1980 surface (——\)/gradient level (<——() wind data and streamline analysis in the vicinity of Tropical Storm Ed. Wind speeds are in knots.

basic flow patterns which existed at the surface and 200 mb levels during most of Ed's existence. The strong surface ridge mentioned above extended from the Asian mainland into the Pacific Ocean north of Ed and maintained a strong northeasterly low-level flow in the vicinity of the Philippine Islands (Fig. 3-28-1). At the same time, strong southwesterly flow at the 200 mb level was present off the east coast of the Philippines (Fig. 3-28-2). The resultant strong vertical wind shear not only caused Ed to weaken as his convection moved off to the northeast, but it also helped to maintain a confused steering flow which induced Ed to follow an erratic course while he was northeast of Simar.

Eventually, after most of his convection had been sheared off, Ed's surface center began to track to the southwest under the influence of the strong surface ridge to the north. Dissipation as a significant tropical cyclone was completed on the 24th as the remnants of Ed moved into northern Mindanao.

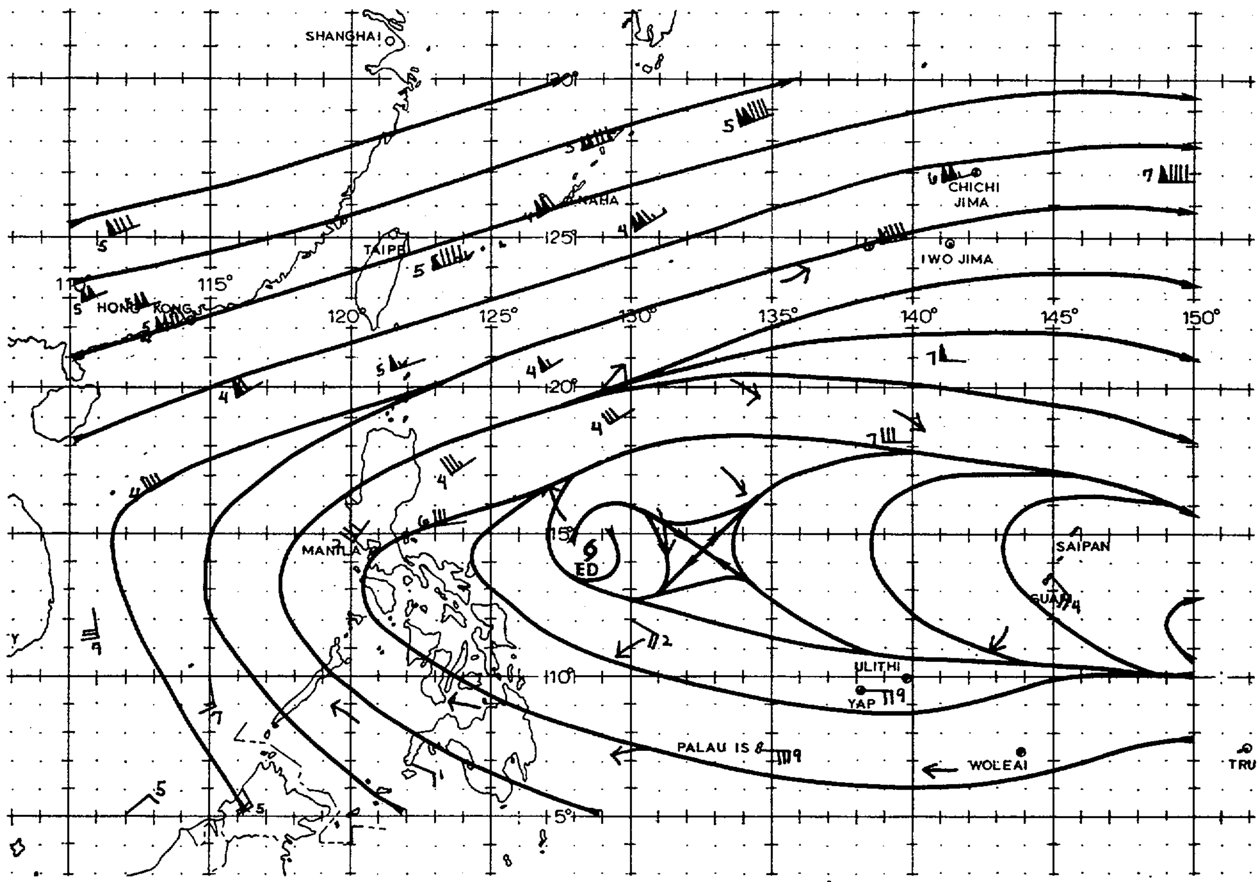


FIGURE 3-28-2. The 200000Z December 1980 200 mb streamline analysis. Wind speeds are in knots.

2. NORTH INDIAN OCEAN TROPICAL CYCLONES

During 1980, there was a notable lack of significant tropical cyclone activity in the North Indian Ocean area (Table 3-3). Two tropical cyclones developed near the end of the year: one in the Arabian Sea (TC 23-80) and one that began in the Bay of Bengal (TC 27-80) and tracked south of India into the Arabian Sea. This was a dramatic decrease from the 1979 total of seven which was the greatest number observed in the two areas since JTWC expanded its area of responsibility westward to include the Arabian Sea in 1975.

TABLE 3-3 NORTH INDIAN OCEAN

1980 SIGNIFICANT TROPICAL CYCLONES

CYCLONE	PERIOD OF WARNING	CALENDAR DAYS OF WARNING	MAX SFC WIND (KT)	EST MIN SLP	NUMBER OF WARNINGS	DISTANCE TRAVELLED (NM)
TC 23-80	17 NOV-19 NOV	3	35	995	8	940
TC 27-80	16 DEC-17 DEC	2	35	996	6	2122
	1980 TOTALS	5			14	

TABLE 3-4

1980 SIGNIFICANT TROPICAL CYCLONE STATISTICS

NORTH INDIAN OCEAN	JAN	FEB	MAR	APR	MAY	JUN	JUL	AUG	SEP	OCT	NOV	DEC	TOTAL
ALL CYCLONES	0	0	0	0	0	0	0	0	0	0	1	1	2
(1971-1979) AVERAGE*	0.1	0	0	0.3	0.5	0.4	0	0	0.5	0.8	1.5	0.3	4.4

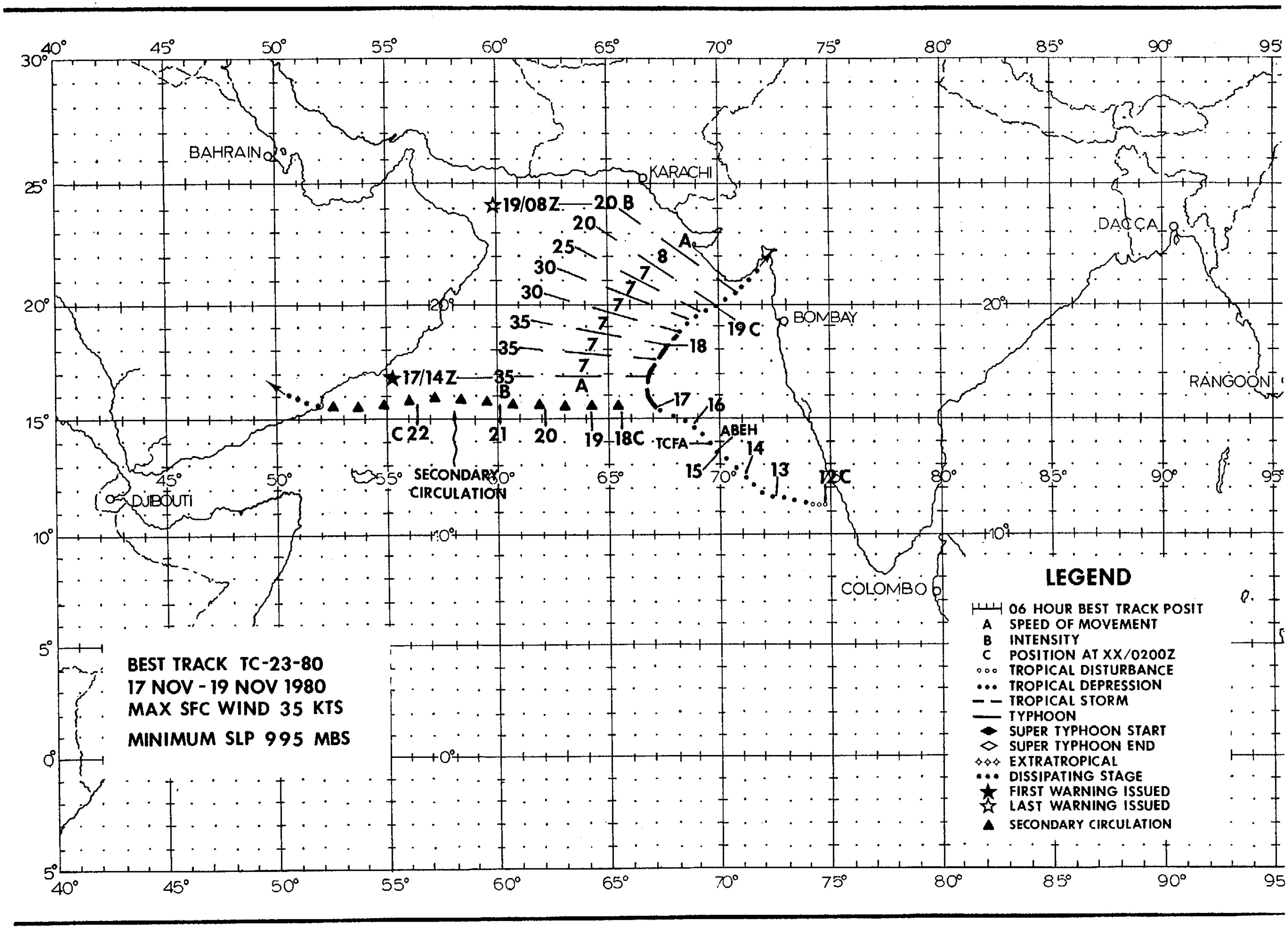
FORMATION ALERTS 2 of the 7 (28%) Formation Alert Events developed into numbered cyclones.

WARNINGS Number of warning days: 5

Number of warning days with 2 cyclones: 0

Number of warning days with 3 or more cyclones: 0

*From 1971 through 1979, only Bay of Bengal cyclones were considered; the JTWC area of responsibility was extended in 1975 to include Arabian Sea cyclones.

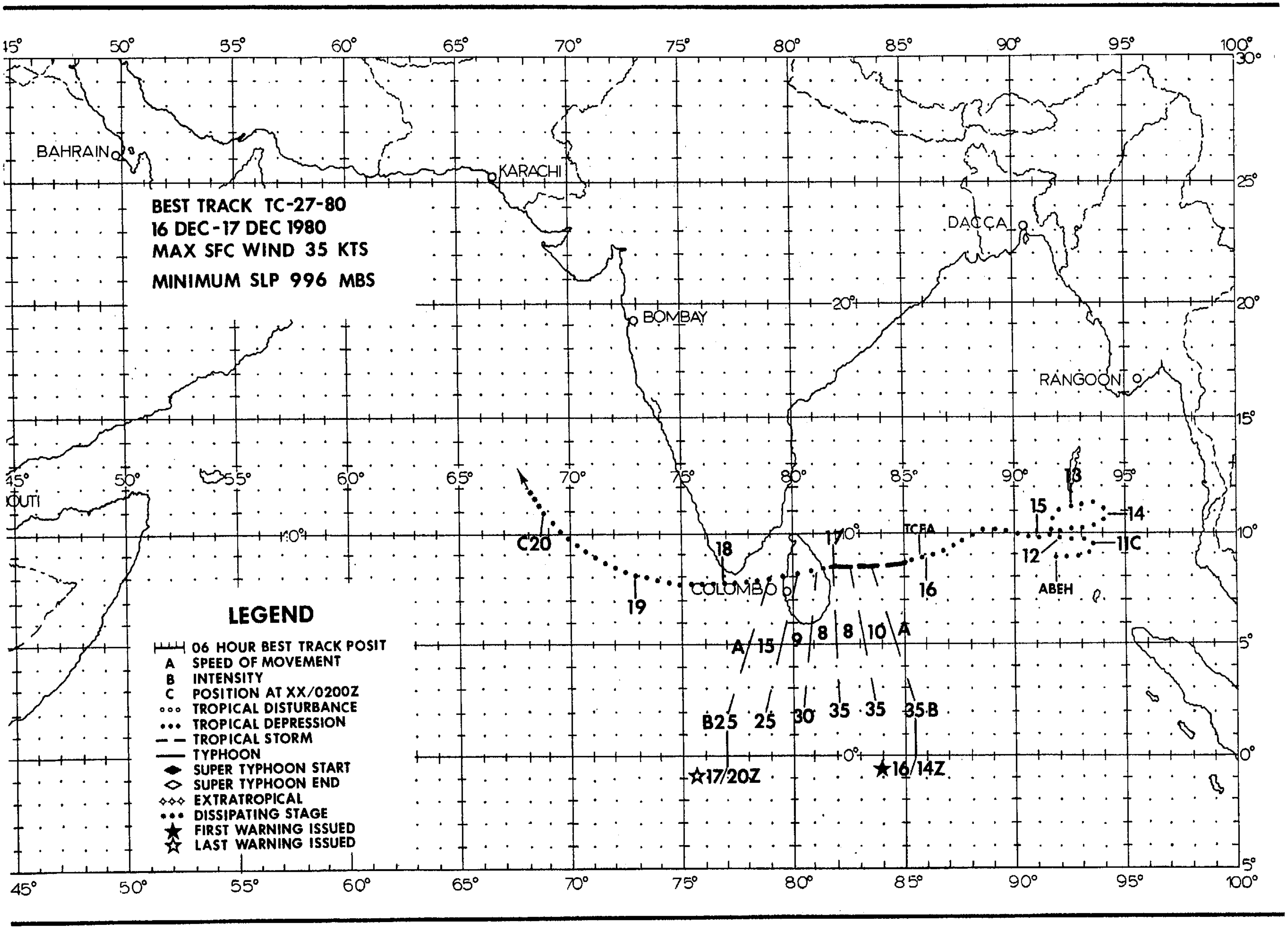

BEST TRACK TC-23-80
17 NOV - 19 NOV 1980
MAX SFC WIND 35 KTS
MINIMUM SLP 995 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0200Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
SECONDARY CIRCULATION
BAHRAIN
KARACHI
DACCA
BOMBAY
RANGOON
DJIBOUTI
COLOMBO
19/08Z
17/14Z
SECONDARY CIRCULATION
TCFA
ABEH

TC 23-80 was the first of only two significant tropical cyclones in 1980 to occur over the North Indian Ocean. It developed during the autumn transition season just prior to the northeast monsoon period. Originating as an area of enhanced convection in the monsoon trough off the southwest coast of India, TC 23-80 tracked steadily northwestward over the Arabian Sea between 12 and 17 November.

On 17 November, a mid-tropospheric trough tracking eastward toward India began to induce TC 23-80 to recurve to the northeast. The main area of convection and the associated low-level center did indeed begin to recurve. However, satellite imageries, which were evaluated during post-analysis (Fig. 3-29-1), indicate that a second low-level circulation formed southwest of TC 23-80 on 17 November. The first circulation continued northeastward and dissipated near the Indian coast, while the secondary circulation continued moving westward as a well-defined, exposed low-level circulation (Fig. 3-29-2). This secondary circulation eventually weakened and dissipated on the coast of Saudi Arabia, five days after cyclogenesis.

FIGURE 3-29-1. TC 23-80 just prior to recurvature over the Arabian Sea. The low-level secondary circulation (⊗) is just beginning to develop on the extreme southwest edge of the main mass of convection, 17 November 1980, 1539Z. (NOAA6 imagery from AFGWC, Offutt AFB, Nebraska)

FIGURE 3-29-2. The secondary low-level center as a well-developed circulation tracking westward towards Saudi Arabia, 19 November 1980, 0358Z. (NOAA6 imagery from AFGWC, Offutt AFB, Nebraska)


BEST TRACK TC-27-80
16 DEC-17 DEC 1980
MAX SFC WIND 35 KTS
MINIMUM SLP 996 MBS
LEGEND
06 HOUR BEST TRACK POSIT
A SPEED OF MOVEMENT
B INTENSITY
C POSITION AT XX/0200Z
TROPICAL DISTURBANCE
TROPICAL DEPRESSION
TROPICAL STORM
TYPHOON
SUPER TYPHOON START
SUPER TYPHOON END
EXTRATROPICAL
DISSIPATING STAGE
FIRST WARNING ISSUED
LAST WARNING ISSUED
BAHRAIN
KARACHI
DACCA
BOMBAY
RANGOON
COLOMBO
ABEH
TCFA
16/14Z
17/20Z
A 15 9 8 8 10 A
B25 25 30 35 35 35B
C20 19 18 17 16 15 14 13 12 11C
45° 50° 55° 60° 65° 70° 75° 80° 85° 90° 95° 100°
30° 25° 20° 15° 10° 5° 0° 5°