

TYPHOON NORRIS
BEST TRACK TC-15
24 AUG-28 AUG 1980
MAX SFC WIND 90 KTS
MINIMUM SLP 950 MBS

LEGEND

- 06 HOUR BEST TRACK POSIT
- A SPEED OF MOVEMENT
- B INTENSITY
- C POSITION AT XX/0000Z
- ... TROPICAL DISTURBANCE
- ... TROPICAL DEPRESSION
- TROPICAL STORM
- TYPHOON
- ◆ SUPER TYPHOON START
- ◇ SUPER TYPHOON END
- ◇◇ EXTRATROPICAL
- ... DISSIPATING STAGE
- ★ FIRST WARNING ISSUED
- ★ LAST WARNING ISSUED

The near equatorial trough was reestablished between Guam and Ponape as Typhoon Marge moved northward toward Marcus Island on 10 August. A weak surface circulation developed along the trough axis south of Guam and slowly drifted toward the Philippine Islands over the next two weeks. Although this disturbance never developed into a significant tropical cyclone, it played a major role in delaying the intensification of a disturbance that tracked westward from Wake Island and eventually became Typhoon Norris.

A deep Tropical Upper-Tropospheric Trough (TUTT) was first analyzed on the 151200Z 200 mb analysis over the Marshall Islands from Wake Island southwestward to Truk. Sparse surface data gave no indications of a perturbation in the low-level tradewind flow at that time.

For the next seven days, the TUTT and associated convective activity to the southeast migrated slowly westward. Figure 3-15-1 depicts the position of the TUTT in relation to the area of enhanced convection that eventually developed into Typhoon Norris. It was not until 211200Z that the upper-level disturbance was reflected at the surface as a weak circulation.

The increased convection and resultant heating of the upper troposphere, as the

disturbance approached a position north of Guam, is graphically illustrated in Figure 3-15-2. The streamline analysis reveals the development of a sharp ridge which built northeastward toward Marcus Island and eventually split the TUTT into two cells. By 230000Z, an upper-level anticyclone had formed over the surface disturbance, and, as the disturbance continued to organize, a Tropical Cyclone Formation Alert was issued at 230900Z. The preceding discussion illustrates the initiation of a tropical cyclone induced by upper-level divergence and enhanced convection southeast of the TUTT cell (Sadler, 1978). Norris tracked virtually straight west-northwestward at an average speed of 12 kt (22 km/hr) from the time of first warning as a tropical depression at 240200Z until landfall on northern Taiwan at 271600Z. This straight track was due to the strong mid-level subtropical ridge which extended along 27N from southern China eastward to the International Dateline during the latter part of August.

The circulation mentioned earlier near the Philippine Islands prevented Norris from developing and intensifying more rapidly. The surface flow pattern was split between the two circulations until 260000Z when the other circulation finally went ashore over Luzon and dissipated. With all the low-level inflow now available, Norris intensi-

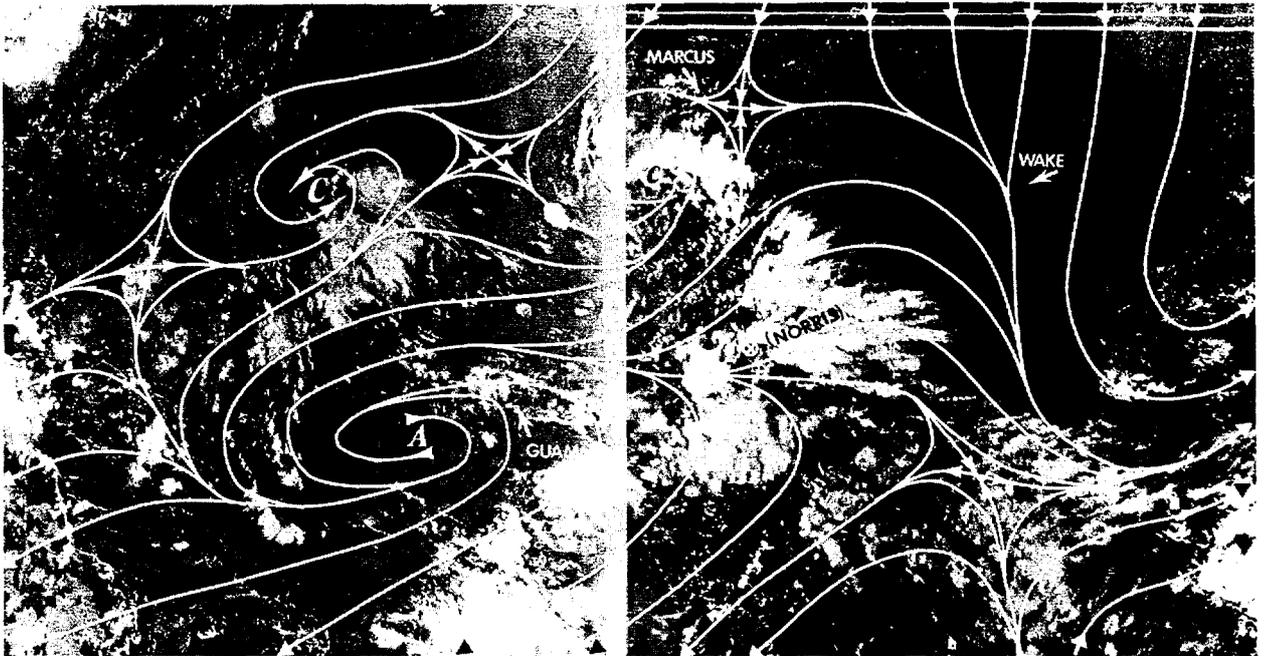


FIGURE 3-15-1. 210000Z August 1980 200 mb streamline analysis superimposed on satellite imagery at 202200Z. This figure depicts convective activity associated with upper-level cyclonic circulations and the enhanced convection southeast of the TUTT that eventually developed into Typhoon Norris. [NOAA6 imagery]

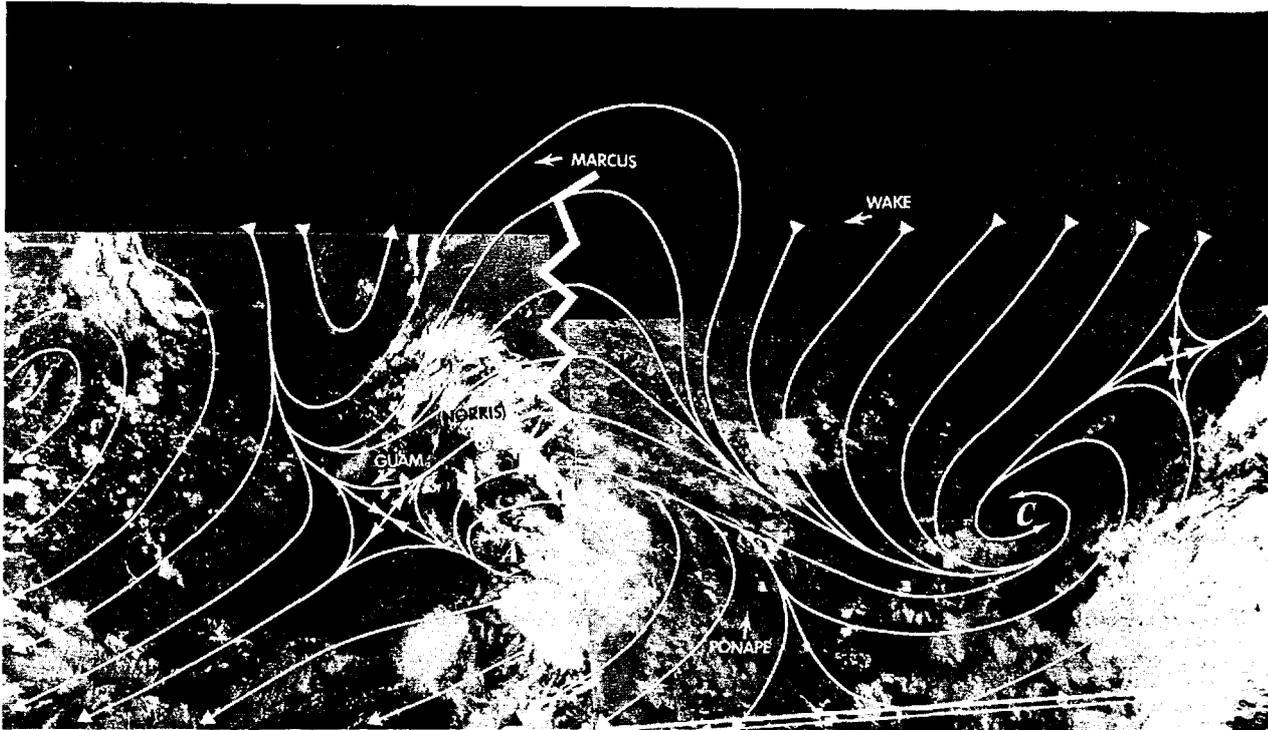


FIGURE 3-15-2. 220000Z August 1980 200 mb streamline analysis superimposed on satellite imagery at 212100Z. This figure illustrates convective heating which built a sharp ridge to the east of the TUTT cell. This ridge pushed the TUTT northward and eventually split it into two cells. (NOAA6 imagery)

fied quickly from 50 kt (26 m/sec) and 985 mb at 260008Z to a peak of 90 kt (46 m/sec) and 950 mb about 36 hours later.

Norris' equivalent potential temperature (θ_e) and minimum sea-level pressure (MSLP) curves intersected at 260000Z. Using JTWC's θ_e /MSLP study (see discussion on Super Typhoon Kim), Norris' sea level pressure was expected to fall 44 mb and maximum surface winds to increase 55 kt (28 m/sec) beyond that point. It seems very likely that this intensification would have occurred if land-fall on Taiwan had not taken place within 42 hours.

The well-established mid-level ridge north of Norris, with a strong high pressure cell located between Taiwan and Okinawa, was responsible for the climatological west-northwestward track with Norris skirting the northern tip of Taiwan. However, with Norris 500 nm (926 km) southeast of Taiwan, the 260000Z 500 mb analysis indicated that the ridge was beginning to weaken between Taiwan and Okinawa with the high pressure cell retreating northeastward to a position east of Okinawa. By 261200Z, a definite break in the ridge was evident with the high cell now over the Bonin Islands and a secondary cell located near 25N 112E over southern China. The numerical forecast series during this period also supported the persistence of this break. Thus, JTWC's warnings after 260000Z were consistent in forecasting recurvature



FIGURE 3-15-3. Typhoon Norris 20 nm (37 km) south-east of Vonagunijima (WMO station 47912) and 95 nm (176 km) southeast of Taipei near peak intensity of 90 kt (46 m/sec), 27 August 1980, 1051Z (NOAA6 imagery). Vonagunijima reported sustained winds of 80 kt (41 m/sec) one hour later.

north of Taiwan and along the coast of mainland China into the Korea Strait.

Norris passed 10 nm (18 km) southwest of Yonagunijima at 271200Z. At that time, the island reported southeast winds of 80 kt (41 m/sec) and a sea-level pressure of 952.2 mb (Fig. 3-15-3). Norris then turned to a more westward track toward northern Taiwan. Excellent radar coverage from the island stations of Ishigakijima and Miyakojima and from Hua-Lien on Taiwan permitted JTWC to follow Norris as he tracked across Taiwan and into the Formosa Strait just north of Hsin-Chu. Strongest surface winds of 39 kt (20 m/sec) with gusts to 64 kt (33 m/sec) on northern Taiwan were reported by Taipei at 271600Z (Fig. 3-15-4).

Norris' track across Taiwan, change in speed, and observed weakening were classic examples of the effects of Taiwan on tropical cyclones (Brand and Brelloch, 1973). The mountainous terrain of Taiwan apparently produced an induced surface low on the lee side of the mountain range, resulting in the marked increase in speed and the westward bend in Norris' track.

Landfall just south of Fu-Chou on mainland China occurred about 280900Z and, although penetrating deeper inland than forecast, Norris eventually recurved northeastward and the remnants linked with a frontal system that moved out over the Yellow Sea and Sea of Japan.



FIGURE 3-15-4. Typhoon Norris as seen by radar at Hua-Lien, 27 August 1980, 0800Z. (Photograph courtesy of the Central Weather Bureau, Taipei, Taiwan)